The Interfacial-Area-Based Relative Permeability Function

ZF Zhang
R Khaleel

September 2009
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ZF Zhang
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the U.S. Department of Energy
under Contract DE-AC05-76RL01830

Pacific Northwest National Laboratory
Richland, Washington 99352

1 Fluor Government Group, Richland, Washington.
Summary

CH2M Hill Plateau Remediation Company (CHPRC) requested the services of the Pacific Northwest National Laboratory (PNNL) to provide technical support for the Remediation Decision Support activity within the Soil and Groundwater Remediation Project. A portion of the support provided in fiscal year 2009 was used to develop an alternative approach to estimating the soil unsaturated hydraulic conductivity. This alternative approach uses the interfacial-area-based relative permeability (k_r) model presented by Embid\textsuperscript{1} rather than the more traditional permeability models of Burdine\textsuperscript{2} and Mualem.\textsuperscript{3} The expectation was that the Embid k_r model would improve the estimation of unsaturated conductivity for at least a subset of soil types. Three retention functions (Brooks and Corey,\textsuperscript{4} van Genuchten,\textsuperscript{5} and modified van Genuchten) were successfully combined with the Embid k_r model. The k_r relationship from the Brooks-Corey-Embid combination for the wetting phase is identical to that from the Brooks-Corey-Burdine combination. The general performance of the combined models is shown using typical hydraulic parameters. The relative permeability models for the wetting phase were further examined using two datasets from the literature. The results indicate that the interfacial-area-based model can describe the relative permeability of the wetting phase reasonably well. However, the comparison of the k_r relationship from the van-Genuchten-Embid combination with that from the van-Genuchten-Mualem combination shows mixed performance results. Further tests are needed with a larger data set.

\textsuperscript{1} Embid, DSM. 1997. “Modeling Capillary Pressure and Relative Permeability for Systems with Heterogeneous Wettability.” Ph.D Dissertation, University of Texas, Austin, Texas.
Acknowledgments

This work is a portion of the Pacific Northwest National Laboratory’s Remediation Decision Support project managed by George V. Last. Funding for this work was provided by the CH2M HILL Plateau Remediation Company.
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1.0 Introduction

The U.S. Department of Energy’s (DOE’s) “Report to the House and Senate Committees on Appropriations on Groundwater Vadose Zone Organization and Operations at the Hanford Site”\(^1\) indicates that future decisions on Hanford cleanup shall be based on an integrated understanding of how contaminants move through the environment. As the Plateau Remediation Contractor, CH2M HILL is required to develop a process to manage risk assessment activities across the Hanford Site, and maintain the key physical, chemical, and other parameters/assumptions associated with modeling the fate and transport of environmental contaminants for remediation decision support. To this end, CH2M HILL Plateau Remediation Company (CHPRC) has requested technical support from the Pacific Northwest National Laboratory (PNNL)\(^2\). A portion of the support requested in fiscal year 2009 was to develop an alternative approach to estimating the soil unsaturated hydraulic conductivity.

Accurate description of soil relative permeability ($k_r$) is necessary in modeling the unsaturated or multi-phase fluid flow in the vadose zone. In the past decades, the Burdine (1953) and Mualem (1976) relative permeability models, in combination with the different water-retention models, have been the primary models used to calculate the unsaturated hydraulic conductivity. Researchers have been searching for alternative models for a better description of the relative permeability. For example, Embid (1997) developed the relative permeability functions based the interfacial areas between the solid, the wetting, and the non-wetting phases, and the Carmen-Kozeny permeability equation (Carmen 1937). The expectation was that the Embid $k_r$ model would improve the estimation of unsaturated conductivity for at least a subset of soil types.

In this report, we document how we incorporated the Brooks and Corey model (1964), van Genuchten model (1980), and a modified van Genuchten water-retention model into the Embid (1997) relative permeability model. The models were examined using two datasets from the literature.

Section 2.0 of this report describes the relationships between soil relative permeability and saturation after incorporating the water retention functions into the Embid relative model and demonstrates the general performance of each model. Section 3.0 examines the interfacial-area-based relative-permeability relationships using data from the literature.

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\(^1\) Rispoli JA. 2006. Letter to the Honorable Thad Cochran (Chairman, Senate Appropriations Committee) from James A. Rispoli (Assistant Secretary for Environmental Management, U.S. Department of Energy), March 29, 2006.

\(^2\) Pacific Northwest National Laboratory is operated by Battelle for the U.S. Department of Energy under Contract DE-AC05-76RL01830.
2.0 Relative Permeability

This section presents the relative permeability relationships after incorporating the water retention functions into the Embid (1997) relative model and demonstrates their general performance. It is customary to assign non-positives value to the pressure head, h, while positive values are often given to the capillary, P_c. To prevent confusion and potential error, the absolute value of pressure head, |h|, is used below in the place of capillary pressure, P_c, after proper conversion.

2.1 The Relative Permeability Relationship

Building on the Carmen-Kozeny equation (Carmen 1937) of permeability, Embid (1997) developed the relative permeability functions for the wetting and the non-wetting phases based on the interfacial areas between the phases:

wetting phase:

\[ k_{rw} = \frac{\tau_s S_w^3}{\tau_w} \left[ \frac{1}{\beta(0,S_w)/\beta(0,1)} \right]^2 \]  

(2.1a)

non-wetting phase:

\[ k_{rn} = \frac{\tau_s (1-S_w)^3}{\tau_n} \left[ \frac{1}{1-\beta(0,S_w)/\beta(0,1)} \right]^2 \]  

(2.1b)

where

\[ \beta(a,b) = \int_a^b |h| dS_w \], \quad a \leq b \]  

(2.1c)

where \( S_w \) is the wetting phase saturation; \( k_{rw} \) and \( k_{rn} \) are the relative permeability for the wetting and non-wetting phases, respectively; \( \tau_s \) is the flow-path tortuosity at full saturation; and \( \tau_w \) and \( \tau_n \) are the tortuosity at saturation \( S_w \). We found that there is a minimum saturation \( S_w^{\min} \), below which \( k_{rn} \) is larger than 1. Hence, Equation (2.1b) is valid only when \( S_w \geq S_w^{\min} \). Soil water retention functions can be incorporated into the above relationships to produce the relative permeability functions.

2.2 Brooks and Corey Water-Retention Model

The Brooks and Corey (1964) model (BC model) expresses water retention as follows:

\[ S_w = \left( \frac{|h|}{h_c} \right)^{\lambda}, \quad \lambda > 0, \text{ or} \]  

(2.2a)

\[ |h| = |h_c| S_w^{-\lambda}, \quad \lambda > 0 \]  

(2.2b)
where \( h_e \) is the air-entry pressure-head and \( \lambda \) is a pore-size distribution parameter. Substituting Equation (2.2b) into (2.1c) derives the following:

\[
\beta(a,b) = \frac{|h_e|}{1-1/\lambda} (b^{1-1/\lambda} - a^{1-1/\lambda})
\]

\[
\beta(0,S_w) = \frac{|h_e|}{1-1/\lambda} (S_w^{1-1/\lambda})
\]

\[
\beta(0,1) = \frac{|h_e|}{1-1/\lambda}
\]  

(2.3)

The Burdine tortuosity model is assumed:

\[
\frac{\tau_s}{\tau_w} = S_w^2
\]

\[
\frac{\tau_s}{\tau_n} = S_n^2 = (1-S_w)^2
\]  

(2.4)

Substituting Equations (2.3) and (2.4) into Equation (2.1) yields the relative permeability for the wetting and non-wetting phases:

wetting phase:

\[
k_{rw} = S_w^{3+2/\lambda}
\]  

(2.5a)

non-wetting phase:

\[
k_{rn} = \frac{(1-S_w)^5}{(1-S_w^{1-1/\lambda})^2}, \quad S_w \geq S_w^{min}
\]  

(2.5b)

Equation (2.5) will be referred to as the Brooks and Corey-Embid (BC-E) model. Equation (2.5a) is the same as the formulation obtained with the Brooks and Corey (1964) retention function and the Burdine (1953) relative permeability model. The \( S_w^{min} \) may be estimated by the following empirical formula:

non-wetting phase:

\[
S_w^{min} \approx 2 \exp(-\lambda)
\]  

(2.6)

The \( S_w^{min} \) vs. \( \lambda \) curve is shown in Figure 2.1. Examples of \( k_{rw} \) and \( k_{rn} \) are shown in Figure 2.2.
Figure 2.1. Minimum Saturation in the $k_m$ Relationship Based on the Brooks and Corey (1964) Model. Points: numerical calculation; line: empirical relation Equation (2.6).

Figure 2.2. The Relative Permeability Function Based on the Brooks-Corey Retention Function. Solid line: $\lambda = 2$; dotted line: $\lambda = 3$; dashed line: $\lambda = 10$. 
2.3 van Genuchten Water Retention Model

The van Genuchten (1980) model (VG model) expresses the water retention as follows:

\[ S_w = \left[ 1 + \left( \frac{\alpha_{vg} h}{S_w} \right)^m \right]^n, \quad \text{or} \]

\[ | h | = \frac{1}{\alpha_{vg}} \left( S_w^{-1/m} - 1 \right)^{1/n} \tag{2.7a} \]

\[ S_w = \left[ 1 + \left( \frac{\alpha_{vg} h}{S_w} \right)^m \right]^n, \quad \text{or} \]

\[ | h | = \frac{1}{\alpha_{vg}} \left( S_w^{-1/m} - 1 \right)^{1/n} \tag{2.7b} \]

where \( \alpha_{vg} \) is the inverse capillary length, \( n \) is a pore-size distribution parameter, and \( m = 1 - 1/n \). After Equations (2.7) and (2.4) are substituted into Equation (2.1c), no closed-form expressions can be derived for \( k_{rw} \) and \( k_{rn} \). To determine the behavior of \( k_{rw} \) and \( k_{rn} \), we follow the same procedures as those of Niemet et al. (2002) by letting \( u = S_w^{1/m}, w = m - 1/n (= 1-2/n), \) and \( z = 1+1/n: \)

\[ \beta(0,u) = \frac{m}{\alpha} \int_0^u (1-u) z^{-1} (u^{w-1}) d\alpha, \; z > 0, \; w > 0 \tag{2.8} \]

Here \( \beta(0,1) \) has the format of the beta function of \( w \) and \( z \) (Abramowitz and Stegun 1972, p. 258); \( \beta(0,S_w) \) is the incomplete beta function; and \( \beta(0,S_w)/\beta(0,1) \) is the regularized incomplete beta function of \( w \) and \( z \) (Abramowitz and Stegun 1972, p. 944, Equation 26.5.1). Because \( w > 0 \) is required in the beta function, \( 1-2/n > 0 \) and \( n > 2 \). This indicates that any value of \( n \leq 2 \) will make Equation (2.8) not integratable. The relative permeability based on Equation (2.8) will be referred to as the van Genuchten-Embld (VG-E) model.

As in the Brooks and Corey model, there is a \( S_w^{\min} \) for the VG-model-based \( k_{rn} \) relationship. The \( S_w^{\min} \) may be estimated by the following empirical formula:

\[ S_w^{\min} \approx 7 \exp(-n) \tag{2.9} \]

The \( S_w^{\min} \) vs. \( n \) curve is shown in Figure 2.3. Examples of the \( k_{rn} \) and \( k_{rn} \) are shown in Figure 2.4.
Figure 2.3. Minimum Saturation in the $k_m$ Relationship Based on the van Genuchten (1980) Model. Points: numerical calculation; line: empirical relation Equation (2.9).

Figure 2.4. The Relative Permeability Function Based on the van Genuchten (1980) Retention Function. Solid line: $n = 2.5$; dotted line: $n = 10$; dashed line: $n = 10$.

2.4 Modified van Genuchten Water Retention Model

For an integratable relationship of relative permeability, the van Genuchten (1980) model was modified as the modified van Genuchten (MVG) model. The format of the modified relationships is
identical to the VG model except \( m = 1 + 1/n \) and \( n > 0 \) (rather than \( m = 1 - 1/n \) and \( n > 1 \) as required by the VG model). To prevent confusion, Greek letters \( \mu \) and \( \nu \) are used here in the place of \( m \) and \( n \), respectively:

\[
S_w = \left[ 1 + \left| \frac{\alpha_{mvg}}{\alpha_{mvg}} \right| h \right]^{\mu} \quad \mu = 1 + 1/\nu, \quad \nu > 0 \quad \text{or} \quad \mu = 1 + 1/\nu, \quad \nu > 0
\]

(2.10a)

| \ h | = \frac{1}{\mu_{mvg}} \left( S_w^{-1/\mu} - 1 \right)^{1/\nu} \quad \mu = 1 + 1/\nu, \quad \nu > 0

(2.10b)

where \( \alpha_{mvg} \) is the inverse capillary length, \( \nu \) is the pore-size distribution parameter, and \( \mu = 1 + 1/\nu \).

Substituting Equations (2.10) into (2.1c) produces

\[
\beta(a, b) = \frac{1}{\alpha} \left( 1 - a^{1/\mu} \right)^{\mu} - \left( 1 - b^{1/\mu} \right)^{\mu}
\]

\[
\beta(0, S_w) = \frac{1}{\alpha} \left[ 1 - \left( 1 - S_w^{1/\mu} \right)^{\mu} \right]
\]

\[
\beta(0, 1) = \frac{1}{\alpha}
\]

(2.11)

Further substitution of Equations (2.11) and (2.4) into Equations (2.1a) and (2.1b) gives

wetting phase:
\[
k_{rw} = \frac{S_w^5}{\left[ 1 + \left( 1 - S_w^{1/\mu} \right)^{\mu} \right]^2} \quad \mu = 1 + 1/\nu, \quad \nu > 0
\]

(2.12a)

non-wetting phase:
\[
k_{rn} = \frac{\left( 1 - S_w \right)^5}{\left( 1 - S_w^{1/\mu} \right)^{2\mu} \left( 1 - S_w^{1/\mu} - \mu \right)^{2\mu}} \quad S_w > S_w^{min}, \quad \mu = 1 + 1/\nu, \quad \nu > 0
\]

(2.12b)

We refer to Equation (2.12) as the Modified-van Genuchten-Embid (MVG-E) model.

As in the previous models, there is a \( S_w^{min} \) for the MVG-E-based \( k_{rn} \) relationship. The \( S_w^{min} \) may be estimated by the following empirical formula:

non-wetting phase:
\[
S_w^{min} \approx 2 \exp(-\nu)
\]

(2.13)

The format of Equation (2.13) is the same as that of Equation (2.6). The \( S_w^{min} \) vs. \( n \) curve is shown in Figure 2.5. Examples of the \( k_{rw} \) and \( k_{rn} \) are shown in Figure 2.6.
Figure 2.5. Minimum Saturation in the $k_{rn}$ Relationship. Points: numerical calculation; line: empirical relation Equation (2.13).

Figure 2.6. The Relative Permeability Function Based on the Modified van Genuchten Retention Function. Solid line: $\nu = 2.5$; dotted line: $\nu = 10$; dashed line: $\nu = 10$. 
3.0 Test of the Interfacial-Area-Based Functions

In this section, the interfacial-area-based relative-permeability relationships are examined using data from the literature. The retention parameters were fitted to the data. The fitted parameters were then used to predict the relative permeability of the corresponding soil. As a comparison, the results based on the van Genuchten (1980) retention function and the Mualem (1976) relative permeability model (VG-M) are also presented.

3.1 The van Genuchten (1980) Dataset

We digitized the soil water retention and relative permeability data for five soils in the van Genuchten (1980) paper. The saturated water content, θs, residual water content, θr, saturated hydraulic conductivity, $K_s$, and the fitted parameters for the BC, VG, and MVG water-retention models are summarized in Table 3.1. The observed and calculated hydraulic properties are shown in Figure 3.1 through Figure 3.5. For all of the soils except the Beit Netofa clay, all three models could fit the retention and predict the relative permeability reasonably well. However, for the Beit Netofa clay soil, while the fitting to the retention curve was generally acceptable, the prediction of the relatively permeability was either very poor (the BC-E, VG-M, and MVG-E models) or was not applicable (the VG-E model), because the fitted n parameter was less than 2. As pointed out by van Genuchten (1980), the quality of the dataset for the Beit Netofa clay soil may be questionable.

A comparison between the measurements and predictions of the ln($k_r$) for the first four soils in Table 3.1 is shown in Figure 3.6. The fitting MSE of the retention curve and the prediction MSE of $k_{rw}$ are given in Table 3.2. On average, the VG and the MVG retention models had the least and greatest fitting error, respectively; the VG-M and MVG-E $k_r$ models had the least and greatest prediction error of relative permeability, respectively.

Table 3.1. Soil Hydraulic Parameters for the Example Soils Used by van Genuchten (1980). θs, θr, and $K_s$ were from van Genuchten (1980); $h_e$, $\lambda$, $\alpha_{vg}$, n, $\alpha_{mvg}$, and $\nu$ are fitted parameters.

<table>
<thead>
<tr>
<th>Soil Name</th>
<th>$\theta_s$</th>
<th>$\theta_r$</th>
<th>$K_s$</th>
<th>$h_e$</th>
<th>$\lambda$</th>
<th>$\alpha_{vg}$</th>
<th>n</th>
<th>$\alpha_{mvg}$</th>
<th>$\nu$</th>
</tr>
</thead>
<tbody>
<tr>
<td>Hygiene Sandstone</td>
<td>0.250</td>
<td>0.153</td>
<td>108.0</td>
<td>102.1</td>
<td>4.30</td>
<td>0.0082</td>
<td>9.90</td>
<td>0.0079</td>
<td>9.27</td>
</tr>
<tr>
<td>Touchet Silt Loam G.E. 3</td>
<td>0.469</td>
<td>0.190</td>
<td>303.0</td>
<td>148.4</td>
<td>2.64</td>
<td>0.0051</td>
<td>7.25</td>
<td>0.0048</td>
<td>6.65</td>
</tr>
<tr>
<td>Silt Loam G.E. 3</td>
<td>0.396</td>
<td>0.131</td>
<td>4.96</td>
<td>141.2</td>
<td>0.768</td>
<td>0.00431</td>
<td>2.09</td>
<td>0.00154</td>
<td>1.438</td>
</tr>
<tr>
<td>Guelph Loam (Drying)</td>
<td>0.520</td>
<td>0.218</td>
<td>31.6</td>
<td>57.9</td>
<td>0.834</td>
<td>0.0120</td>
<td>2.08</td>
<td>0.00397</td>
<td>1.372</td>
</tr>
<tr>
<td>Beit Netofa Clay</td>
<td>0.446</td>
<td>0.0</td>
<td>0.082</td>
<td>243.9</td>
<td>0.121</td>
<td>0.00198</td>
<td>1.19</td>
<td>4.75E-6</td>
<td>0.548</td>
</tr>
</tbody>
</table>
Table 3.2. Mean Squared Error (MSE) of the Fitted Retention Curve and the Predicted Relative Permeability

<table>
<thead>
<tr>
<th>Soil Name</th>
<th>Fitting MSE$_{θ}$ (cm$^3$cm$^{-3}$)</th>
<th>Prediction MSE$_{\ln kr}$</th>
</tr>
</thead>
<tbody>
<tr>
<td></td>
<td>BC</td>
<td>VG</td>
</tr>
<tr>
<td>Hygiene Sandstone</td>
<td>0.0016</td>
<td>0.00057</td>
</tr>
<tr>
<td>Touchet Silt Loam G.E. 3</td>
<td>0.0015</td>
<td>0.00048</td>
</tr>
<tr>
<td>Silt Loam G.E. 3</td>
<td>0.0016</td>
<td>0.00012</td>
</tr>
<tr>
<td>Guelph Loam (Drying)</td>
<td>0.0020</td>
<td>0.00093</td>
</tr>
<tr>
<td>Beit Netofa Clay</td>
<td>0.0009</td>
<td>0.00042</td>
</tr>
<tr>
<td>Average</td>
<td>0.0015</td>
<td>0.00050</td>
</tr>
</tbody>
</table>

MSE$_{θ}$: MSE of Water Content; MSE$_{\ln kr}$: MSE of log-transformed $kr$

Retention Models: BC – Brooks and Corey (1964); VG – van Genuchten (1980); MVG – Modified van Genuchten
Figure 3.1. Observed and Calculated Soil Hydraulic Properties of the Hygiene Sandstone. (a) Observed and fitted water retention; (b) observed and predicted relative permeability.
Figure 3.2. Observed and Calculated Soil Hydraulic Properties of the Touchet Silt Loam G. E. 3.
(a) Observed and fitted water retention; (b) observed and predicted relative permeability.
Figure 3.3. Observed and Calculated Soil Hydraulic Properties of the Silt Loam G. E. 3. (a) Observed and fitted water retention; (b) observed and predicted relative permeability.
Figure 3.4. Observed and Calculated Soil Hydraulic Properties of the Guelph Loam. (a) Observed and fitted water retention; (b) observed and predicted relative permeability.
Figure 3.5. Observed and Calculated Soil Hydraulic Properties of the Beit Netofa Clay. (a) Observed and fitted water retention; (b) observed and predicted relative permeability.
3.2 The Rockhold et al. Dataset

Rockhold et al. (1988) measured the water retention and hydraulic conductivity of repacked soil from the 300-N lysimeter site (formerly called the Buried Waste Test Facility), and the data were summarized in their Appendix A.1. After assuming that $\theta_r$ was half of their measured minimum and $\theta$ and $\theta_s$ was the maximum $\theta$, we fitted the rest of the retention parameters to the retention data. Then, the parameters were used to predict the measured relative permeability. The pre-determined parameters and the fitted parameters are summarized in Table 3.3. The observed and calculated hydraulic properties are shown in Figure 3.7 through Figure 3.10. The models can describe the hydraulic properties of the sandy soil quite well.

A comparison between the measurements and predictions of the ln($k_r$) for the soils in Table 3.3 is shown in Figure 3.11. The fitting MSE of the retention curve and the prediction MSE of $k_w$ are given in
Table 3.4. On average, the BC and the MVG retention models had the least and greatest fitting error, respectively; the VG-E and VG-M models had the least and greatest prediction error of relative permeability, respectively.

Table 3.3. Soil Hydraulic Parameters for the Soils Studied by Rockhold et al. (1988). θs and Ks were from Rockhold et al. (1988); θr was assumed to be half of the measured minimum θ; h, λ, αvg, n, αavg, and ν are fitted parameters.

<table>
<thead>
<tr>
<th>Soil Name</th>
<th>θs (cm cm⁻³)</th>
<th>θr (cm cm⁻³)</th>
<th>Ks (cm/s)</th>
<th>h (cm)</th>
<th>λ</th>
<th>αvg cm⁻¹</th>
<th>n</th>
<th>αavg cm⁻¹</th>
<th>ν</th>
</tr>
</thead>
<tbody>
<tr>
<td>Soil E with ρb = 1.6 g cm⁻³</td>
<td>0.435</td>
<td>0.043</td>
<td>4.62E-3</td>
<td>11.7</td>
<td>1.09</td>
<td>0.067</td>
<td>2.25</td>
<td>0.025</td>
<td>1.29</td>
</tr>
<tr>
<td>Soil E with ρb = 1.7 g cm⁻³</td>
<td>0.400</td>
<td>0.042</td>
<td>7.12E-3</td>
<td>12.1</td>
<td>0.812</td>
<td>0.058</td>
<td>2.01</td>
<td>0.013</td>
<td>0.973</td>
</tr>
<tr>
<td>Soil F with ρb = 1.6 g cm⁻³</td>
<td>0.422</td>
<td>0.055</td>
<td>9.78E-3</td>
<td>12.97</td>
<td>1.19</td>
<td>0.058</td>
<td>2.49</td>
<td>0.028</td>
<td>1.59</td>
</tr>
<tr>
<td>Soil F with ρb = 1.7 g cm⁻³</td>
<td>0.386</td>
<td>0.043</td>
<td>8.91E-3</td>
<td>12.3</td>
<td>0.816</td>
<td>0.058</td>
<td>2.01</td>
<td>0.013</td>
<td>0.962</td>
</tr>
</tbody>
</table>

ρb = bulk density

Table 3.4. Mean Squared Error of the Fitted Retention Curve and the Predicted Relative Permeability

<table>
<thead>
<tr>
<th>Soil Name</th>
<th>Fitting MSEθ (cm³ cm⁻³)</th>
<th>Prediction MSElnkr</th>
</tr>
</thead>
<tbody>
<tr>
<td></td>
<td>BC</td>
<td>VG</td>
</tr>
<tr>
<td>Soil E with ρb = 1.6 g cm⁻³</td>
<td>0.0012</td>
<td>0.0021</td>
</tr>
<tr>
<td>Soil E with ρb = 1.7 g cm⁻³</td>
<td>0.0015</td>
<td>0.0022</td>
</tr>
<tr>
<td>Soil F with ρb = 1.6 g cm⁻³</td>
<td>0.0008</td>
<td>0.0018</td>
</tr>
<tr>
<td>Soil F with ρb = 1.7 g cm⁻³</td>
<td>0.0012</td>
<td>0.0021</td>
</tr>
<tr>
<td>Average</td>
<td>0.0012</td>
<td>0.0021</td>
</tr>
</tbody>
</table>

MSEθ: MSE of water content; MSElnkr: MSE of log-transformed relative permeability
Retention Models: BC – Brooks and Corey (1964); VG – van Genuchten (1980); MVG – Modified van Genuchten
Figure 3.7. Observed and Calculated Soil Hydraulic Properties of the Column E Soil with Bulk Density of 1.6 g cm$^{-3}$. (a) Observed and fitted water retention; (b) observed and predicted relative permeability.
Figure 3.8. Observed and Calculated Soil Hydraulic Properties of the Column E Soil with Bulk Density of 1.7 g cm\(^{-3}\). (a) Observed and fitted water retention; (b) observed and predicted relative permeability.
Figure 3.9. Observed and Calculated Soil Hydraulic Properties of the Column F Soil with Bulk Density of 1.6 g cm\(^{-3}\). (a) Observed and fitted water retention; (b) observed and predicted relative permeability.
Figure 3.10. Observed and Calculated Soil Hydraulic Properties of the Column F Soil with Bulk Density of 1.7 g cm$^{-3}$. (a) Observed and fitted water retention; (b) observed and predicted relative permeability.
3.3 Summary of Tests

The relative permeability models for the wetting phase were examined using two datasets from van Genuchten (1980) and Rockhold et al. (1988). The results indicate that the interfacial-area-based model can describe the relative permeability of the wetting phase reasonably well. However, the comparison of the \( k_r \) relationship of the van-Genuchten-Embid combination with that of the van-Genuchten-Mualem combination show mixed performance results. Further tests are needed with a larger dataset. Attention should be paid to low \( k_r \) values under relative dry conditions.
4.0 References


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