

# First Quarter Seismic Report for Fiscal Year 2006

Pacific Northwest National Laboratory Hanford Seismic Assessment Team

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September 2006



Prepared for the U.S. Department of Energy under Contract DE-AC05-76RL01830

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# **Summary**

Hanford Seismic Assessment (HSA) provides an uninterrupted collection of high-quality raw and processed seismic data from the Hanford Seismic Network for the U.S. Department of Energy and its contractors. The Hanford Seismic Assessment Team locates and identifies sources of seismic activity and monitors changes in the historical pattern of seismic activity at the Hanford Site. The data are compiled, archived, and published for use by the Hanford Site for waste management, Natural Phenomena Hazards assessments, and engineering design and construction. In addition, the seismic monitoring organization works with the Hanford Site Emergency Services Organization to provide assistance in the event of a significant earthquake on the Hanford Site.

The Hanford Seismic Network and the Eastern Washington Regional Network consist of 41 individual sensor sites and 15 radio relay sites maintained by the Hanford Seismic Assessment Team.

For the Hanford Seismic Network, there were 57 local earthquakes for the first quarter of fiscal year 2006. The largest earthquake for the quarter was a magnitude 2.6 event on November 10, 2005. This earthquake was a random event and occurred in the crystalline basement beneath the Horse Heaven Hills. Stratigraphically, 46 earthquakes occurred in the Columbia River basalt (approximately 0-5 km), 6 earthquakes were in the pre-basalt sediments (approximately 5-10 km), and 5 earthquakes were in the crystalline basement (approximately 10-25 km). Geographically, 50 earthquakes occurred in swarm areas, no earthquakes were on major geologic structures, and 7 earthquakes were classified as random events. The most significant event was the occurrence of 39 earthquakes at the eastern end of the Coyote Rapids Swarm Area at the 'horn' of the Columbia River.

# Acronyms

BWIP Basalt Waste Isolation Project

CDPD Cellular Digital Packet Data

CRBG Columbia River Basalt Group

DOE U.S. Department of Energy

ETNA strong motion accelerometer manufactured by Kinemetrics

EWRN Eastern Washington Regional Network

FY fiscal year

GPS Global Positioning System

HSA Hanford Seismic Assessment

HSN Hanford Seismic Network

M<sub>c</sub> Coda-Length Magnitude

M<sub>L</sub> Local Magnitude

M<sub>w</sub> Moment Magnitude

PNNL Pacific Northwest National Laboratory

RAW Rattlesnake Mountain-Wallula Alignment

SMA strong motion accelerometer

USGS United States Geological Survey

UTC Universal Time, Coordinated

UW University of Washington

WHC Westinghouse Hanford Company

# **Contents**

Sun	nmary	·······		iii
Acre	onym	s		v
1.0	Intro	oduction	n	1.1
	1.1	Missio	on	1.1
	1.2	Histor	y of Monitoring Seismic Activity at Hanford	1.1
	1.3	Docur	nentation and Reports	1.2
2.0	Net	work O <sub>j</sub>	perations	2.1
	2.1	Seism	ometer Sites	2.1
		2.1.1	Station Maintenance	2.5
		2.1.2	Data Acquisition	2.5
	2.2	Strong	Motion Accelerometer Sites	2.5
		2.2.1	Location	2.5
		2.2.2	Site Design.	2.7
		2.2.3	Strong Motion Accelerometer Operations Center	2.8
		2.2.4	Strong Motion Operational Characteristics	2.8
3.0	Eart	hquake	Catalog Description	3.1
	3.1	Coda	Length Magnitude	3.1
	3.2	Veloci	ity Model	3.1
	3.3	Qualit	y Factors	3.1
4.0	Geo	logy an	d Tectonic Analysis	4.1
	4.1	Eartho	uake Stratigraphy	4.1
	4.2	Geolo	gic Structure Beneath the Monitored Area	4.1
	4.3	Tector	nic Pattern	4.4
	4.4	Depth	of Earthquakes	4.5
	4.5	Tector	nic Activity	4.5
		4.5.1	Summary	4.5
		4.5.2	First Quarter Earthquakes of FY 2006	4.6
5.0	Stro	ng Mot	ion Accelerometer Operations	5.1
	5.1	FY 20	06 Triggers of the Hanford SMA Network	5.1
6.0	Cap	abilities	s in the Event of a Significant Earthquake	6.1
	6.1	Use of	the SMA Network in the Event of an Earthquake	6.1
7.0	Refe	erences		7.1

# **Figures**

2.1	Locations of Seismograph Stations and Strong Motion Accelerometer Sites in the Hanford Seismic Network	2.2
2.2	Locations of Seismograph Stations in the Eastern Washington Regional Network	2.4
2.3	Schematic Diagram of a Strong Motion Accelerometer Installation	2.8
4.1	Tectonic Map of Columbia Basin Showing Major Seismic Source Structures	4.2
4.2	Geologic Cross Sections Through the Columbia Basin	4.3
4.3	All Earthquakes Recorded Within the Hanford Monitoring Area during the First Quarter of FY 2006	4.7
4.4	Timing of Earthquakes in the Coyote Rapids Earthquake Swarm during the First Quarter of FY 2006.	4.8
	Tables	
2.1	Seismic Stations in the Hanford Seismic Network	2.1
2.2	Seismic Stations in the Eastern Washington Regional Network	2.3
3.1	Local Seismic Data, October 1, 2005 to December 31, 2005	3.2
3.2	Seismic Velocities for Columbia Basin Stratigraphy	3.6
4.1	Thicknesses of Stratigraphic Units in the Monitoring Area	4.4
4.2	Number of Local Earthquakes Occurring in Stratigraphic Units	4.5
4.3	Summary of Earthquake Locations for FY 2005	4.6

# 1.0 Introduction

This report covers the locations of seismicity within the Hanford monitoring region of south-central Washington for the first quarter of fiscal year (FY) 2006, and provides geologic interpretations of those earthquakes.

#### 1.1 Mission

The principal mission of Hanford Seismic Assessment (HSA) is to ensure compliance with DOE Order 420.1, "Facility Safety" and DOE Order G 420.1-1, Section 4.7, "Emergency Preparedness and Emergency Communications." DOE Order 420.1 establishes facility safety requirements related to nuclear safety design, criticality safety, fire protection, and natural phenomena hazards mitigation. For seismic assessment, this order states:

#### 4.4.5 Natural Phenomena Detection.

Facilities or sites with hazardous materials shall have instrumentation or other means to detect and record the occurrence and severity of seismic events.

The HSA supports Hanford Site emergency services organizations in complying with DOE Order G 420.1-1, Section 4.7, "Emergency Preparedness and Emergency Communications," by providing assistance in the event of an earthquake on the Hanford Site.

In addition, the HSA provides an uninterrupted collection of high-quality raw seismic data from the Hanford Seismic Network (HSN) located on and around the Hanford Site, and the Eastern Washington Regional Network (EWRN). This report provides interpretations of seismic events from the Hanford Site and vicinity. HSA locates and identifies sources of seismic activity, monitors changes in the historical pattern of seismic activity at the Hanford Site, and builds a "local" earthquake database (processed data) that is permanently archived. The focus of this report is the precise location of earthquakes proximal to or on the Hanford Site, specifically between 46 degrees and 47 degrees north latitude and between 119 degrees and 120 degrees west longitude. Data from the EWRN and other seismic networks in the northwest provide the HSA with necessary regional input for the seismic hazards analysis at the Hanford Site. These seismic data are used to support Hanford Site contractors for waste management activities, Natural Phenomena Hazards assessments, and engineering design and construction.

# 1.2 History of Monitoring Seismic Activity at Hanford

Assessing seismic activity at the Hanford Site was initiated in 1969 by the United States Geological Survey (USGS) under a contract with the U.S. Atomic Energy Commission. In 1975, the University of Washington (UW) assumed responsibility for the network and subsequently expanded it. In 1979, the Basalt Waste Isolation Project (BWIP) became responsible for collecting seismic data for the Hanford Site as part of site characterization activities. Rockwell Hanford Operations, followed by Westinghouse Hanford Company (WHC), operated the local network and were the contract technical advisors for the EWRN operated and maintained by UW. Funding ended for BWIP in December 1988. Assessment of seismic activity and responsibility for the UW contract were then transferred to WHC's Environmental Division. Maintenance responsibilities for the EWRN also were assigned to WHC who made major upgrades to EWRN sites.

Effective October 1, 1996, all seismic assessment activities were transferred to the Pacific Northwest National Laboratory (PNNL).<sup>1</sup>

The Hanford Strong Motion Accelerometer (SMA) network was constructed during 1997 and came on-line in May 1997. It operated until September 30, 1997, when it was mothballed due to lack of funding. Funding was restored on October 1, 1998, by joint agreement between the U.S. Department of Energy (DOE) and PNNL. Operation of the SMA sites resumed on November 20, 1999, and has operated continuously since that time.

# 1.3 Documentation and Reports

The HSA issues quarterly reports of local activity, an annual catalog of earthquake activity on and near the Hanford Site, and special-interest bulletins on local seismic events. The HSA also provides information and special reports to other functions as requested. Earthquake information provided in these reports is subject to revisions if new data become available. In addition, an archive of all seismic data from the HSN is maintained by PNNL.

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<sup>&</sup>lt;sup>1</sup> Pacific Northwest National Laboratory is operated by Battelle Memorial Institute for the U.S. Department of Energy.

# 2.0 Network Operations

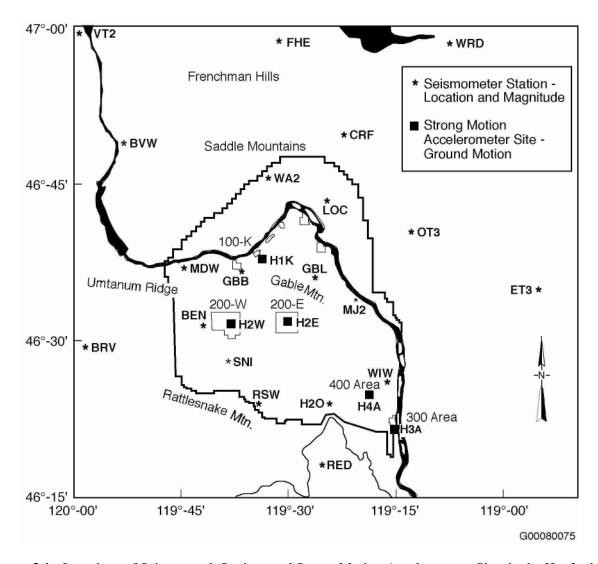
#### 2.1 Seismometer Sites

The seismic network consists of two designs of equipment and sites: seismometer sites and strong motion accelerometer (SMA) sites. Seismometer sites are designed to locate earthquakes and determine their magnitudes and hypocenter locations. SMA sites are designed to measure ground motion and are discussed in Section 2.2.

The Hanford Seismic Network (HSN) and the Eastern Washington Regional Network (EWRN) consist of 41 sensor sites. Most sites are in remote locations and require solar panels and batteries for power. The HSN uses 23 sites (Table 2.1 and Figure 2.1) and the EWRN uses 36 sites (Table 2.2 and Figure 2.2); both networks share 18 sites. The networks have 45 combined data channels because Gable Butte and Frenchman Hills East are three-component sites, each consisting of one vertical, one north-south horizontal, and one east-west horizontal data channel. Both networks use 15 additional telemetry relay sites. Data from all sites or relays are transmitted to the Sigma V building, Richland, Washington.

**Table 2.1.** Seismic Stations in the Hanford Seismic Network

The first column is the three-letter seismic station designator. The latitude and longitude, elevation above sea level in meters, and the full station name follow this. The locations of the stations are all in Washington; locations were derived from a Global Positioning System (GPS). Latitude Longitude Station Elevation (m) Deg. Min. N Deg. Min. W Station Name **BEN** 46N31.13 119W43.02 340 Benson Ranch 45N54.91 BLT 119W10.55 659 Bickelton BRV 46N49.12 119W59.47 920 Black Rock Valley BVW 119W52.99 46N48.66 670 Beverly **CRF** 46N49.50 119W23.22 189 Corfu ET3 46N34.64 118W56.25 286 Eltopia Three Frenchman Hills East \*FHE 46N57.11 119W29.82 455 \*GBB 46N36.49 119W37.62 177 Gable Butte **GBL** 46N35.92 119W27.58 330 Gable Mountain H<sub>2</sub>O 46N23.75 119W25.38 158 Water LOC 46N43.02 119W25.85 210 Locke Island MDW 119W45.66 46N36.79 330 Midway MJ2 119W21.54 46N33.45 146 May Junction Two OT3 46N40.14 119W13.98 322 Othello Three PRO 46N12.73 119W41.15 550 Prosser RED 119W26.30 46N17.92 366 Red Mountain **RSW** 46N23.67 119W35.48 1,045 Rattlesnake Mountain 119W39.60 SNI 46N27.85 312 Snively Ranch VT2 1,270 46N58.04 120W58.95 Vantage Two WA2 46N45.32 119W33.94 244 Wahluke Slope WIW 46N25.76 119W17.26 128 Wooded Island WRD 46N58.20 119W08.69 375 Warden 46N02.93 118W57.73 325 Yellepit \* Three-component station.

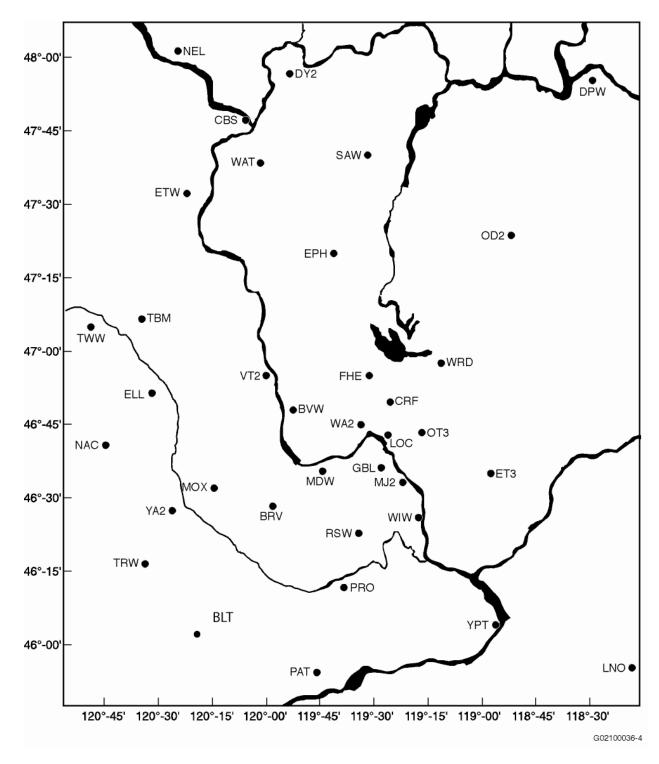


**Figure 2.1**. Locations of Seismograph Stations and Strong Motion Accelerometer Sites in the Hanford Seismic Network (see Tables 2.1 and 2.4 for description of locations). Bickelton (BLT), Prosser (PRO), Yellepit (YPT) are shown on Figure 2.2.

Table 2.2. Seismic Stations in the Eastern Washington Regional Network

The first column is the three-letter seismic station designator. The latitude and longitude, elevation above sea level in meters, and the full station name follow this. The locations of the stations are all in Washington unless otherwise indicated; locations were determined from a Global Positioning System (GPS).

wise indicate	d; locations were detern Latitude	Longitude	Sitioning System (Gr	75). 
Station	Deg. Min. N.	Deg. Min. W.	Elevation (m)	Station Name
BLT	46N54.91	119W10.55	659	Bickelton
BRV	46N29.12	119W59.47	920	Black Rock Valley
BVW	46N48.66	119W52.99	670	Beverly
CBS	47N48.26	120W02.50	1,067	Chelan Butte, South
CRF	46N49.50	119W23.22	189	Corfu
DPW	47N52.25	118W12.17	892	Davenport
DY2	47N59.11	119W46.28	890	Dyer Hill Two
ELL	46N54.58	120W33.98	789	Ellensburg
EPH	47N21.38	119W35.76	661	Ephrata
ET3	46N34.64	118W56.25	286	Eltopia Three
ETW	47N36.26	120W19.94	1,477	Entiat
*FHE	46N57.11	119W29.82	455	Frenchman Hills East
GBL	46N35.92	119W27.58	330	Gable Mountain
LNO	45N52.31	118W17.11	771	Lincton Mountain, Oregon
LOC	46N43.02	119W25.85	210	Locke Island
MDW	46N36.79	119W45.66	330	Midway
MJ2	46N33.45	119W21.54	146	May Junction Two
MOX	46N34.64	120W17.89	501	Moxee City
NAC	46N43.99	120W49.42	728	Naches
NEL	48N04.21	120W20.41	1,500	Nelson Butte
OD2	47N23.26	118W42.58	553	Odessa Two
OT3	46N40.14	119W13.98	322	Othello Three
PAT2	45N53.03	119W45.40	259	Paterson Two
PRO	46N12.73	119W41.15	550	Prosser
RSW	46N23.67	119W35.48	1,045	Rattlesnake Mountain
SAW	47N42.10	119W24.03	701	St. Andrews
TBM	47N10.20	120W35.88	1,006	Table Mountain
TRW	46N17.32	120W32.31	723	Toppenish Ridge
TWW	47N08.29	120W52.10	1,027	Teanaway
VT2	46N58.04	119W58.95	1,270	Vantage Two
WA2	46N45.32	119W33.94	244	Wahluke Slope Two
WAT	47N41.92	119W57.24	821	Waterville
WIW	46N25.76	119W17.26	128	Wooded Island
WRD	46N58.20	119W08.69	375	Warden
YA2	46N31.60	120W31.80	652	Yakima Two
YPT	46N02.93	118W57.73	325	Yellepit
* Three-com	ponent station.			



**Figure 2.2**. Locations of Seismograph Stations in the Eastern Washington Regional Network (see Table 2.2 for location descriptions)

#### 2.1.1 Station Maintenance

The HSN's maintenance records for the seismic sensor and relay sites are on file in the HSA office, Sigma V Building, Richland, Washington.

#### 2.1.2 Data Acquisition

The signals from the seismometer sites are monitored for changes in signal amplitude that are expected from earthquakes. The seismic network is subdivided into spatial groupings of stations that are monitored for nearly simultaneous amplitude changes, resulting in triggering a permanent recording of the events. The groupings and associated weighting schemes are designed to allow very small seismic events to be recorded and to minimize false triggers. Events are classified as local (south-central Washington near the Hanford Site), regional (Western U.S. and Canada), and teleseisms (from farther distances around the world). Local and regional events are usually earthquakes, but quarry and mining explosions are also recorded. Quarry and mining explosions can usually be identified from wave characteristics, time of day, and through confirmation with local government agencies and industries. Frequently, military exercises at the U.S. Army's Yakima Training Center produce a series of acoustic shocks that unavoidably trigger the recording system. Sonic booms and thunder also produce acoustic signals that trigger the recording system.

A PC-based system (Earthworm system) adapted from a USGS program and the UW system was implemented at Hanford during FY 1999. One system has been in continuous operation since January 6, 1999. A second, backup PC system was installed in mid-March 1999, and both systems have been running in parallel since that time. The hardware and software have been periodically upgraded. Data from triggers are collected on a SUN<sup>TM</sup> (registered trademark of Sun Microsystems, Santa Clara, California) workstation that is used to determine earthquake locations and magnitudes (Section 3). Although the two systems are practically identical, there is enough granularity (signal-to-noise) in the trigger timing that they sometimes record exclusive events. In nearly all cases, these exclusive triggers are "false" triggers, not earthquakes or quarry blasts (i.e., from acoustic sources). The remainders are from barely detectable, small signals from regional and teleseismic earthquakes.

# 2.2 Strong Motion Accelerometer Sites

#### 2.2.1 Location

The Hanford SMA network consists of five free-field SMA sites (see Figure 2.1) (Table 2.3). There is one free-field SMA located in each of the 200 Separations Areas, one adjacent to the K Basins in the 100-K Area, one adjacent to the 400 Area where the Fast Flux Test Reactor is located, and one at the south end of the 300 Area. With the termination of the Fast Flux Test Reactor and draining of the liquid sodium coolant taking place now, plans have been made to terminate this site and move the instrument to a new location.

The instrumentation locations were chosen based on two criteria (Moore and Reidel 1996): 1) instruments should be located in areas having the highest densities of people and 2) instruments should be located in areas having hazardous facilities. Some of the highest concentrations of employees at Hanford are 200-East and West Areas, 100-K Area, the Fast Flux Test Facility (400 Area), and the 300 Areas. The 200 Areas are where high-level radioactive waste from past processing of fuel rods is stored in single-

shell and double-shell tanks. In addition, the Canister Storage Facility that holds encapsulated spent fuel rods is in 200-East Area and the new Waste Treatment Plant, currently under construction, are also located in 200-East Area. The 100-K Area contains the K Basins where spent fuel rods from the N Reactor were stored prior to encapsulation. The Cold Vacuum Drying Facility, located in the 100-K Area, is used to encapsulate spent fuel rods from the K Basins prior to shipment to the Canister Storage Building in 200-East Area.

 Table 2.3. Free-Field Strong Motion Accelerometer Sites

Site	Site ID	Location	Latitude Longitude Elevation
100-K Area	H1K	South of K Basins outside 100 Area fence lines.	46° 38.51' 119° 35.53' 152 m
200 East Area	H2E	East of B Plant; north of 7th Street and east of Baltimore Avenue.	46° 33.58' 119° 32.00' 210 m
200 West Area	H2W	West of Plutonium Finishing Plant (PFP) and 200-West Area tree barrier.	46° 33.11' 119° 38.64' 201 m
300 Area	Н3А	South end of 300 Area inside fence lines (NE 1/4, SW 1/4, Sec. 11, T10N, R28E).	46° 21.83' 119° 16.55' 119 m
400 Area	H4A	500 ft from fence line on east side of facility and north of parking area).	46° 26.13' 119° 21.30' 171 m

### 2.2.2 Site Design

All free-field SMA sites consist of a four-panel solar array and two 30-gallon galvanized drums. Each panel has a maximum 42-watt output. The two 30-gallon drums are set in the ground such that the base of the drum is about 1 m below the surface. One drum houses only the SMA; the other drum, which is connected via a sealed conduit to the SMA drum, contains the batteries. Communication is through a General Packet Radio Service (GPRS) system, which provides a continuous radio data-link with an internet service provider. This GPRS system along with the solar power regulator is housed in a small enclosure mounted at the rear of the solar array. The enclosure serves as a junction box for all cabling between equipment inside and outside of the drums through conduit. The antenna for the GPRS is mounted on top of the enclosure. The enclosure permits quick access to check battery conditions and a connection directly to the RS-232 port of the SMA without removing the drum lids.

The SMA instruments are three-component units consisting of one vertical, one north-south horizontal, and one east-west horizontal data channel. The instruments in use are the ETNA<sup>TM</sup> system (registered trademark of Kinemetrics, Inc., Pasadena, California). Instrument specifications are summarized in Table 2.4. In addition to the three-component SMAs, each ETNA SMA unit contains a computer, and a Global Positioning System (GPS) receiver (Figure 2.3). These systems are housed in a watertight box.

The GPRS system provides the internet address connection to access the system. Stations can be monitored from any computer with appropriate access. The data can be downloaded directly at each site via a built in cable connection at the enclosure in case of communication failure.

The SMAs have an internal GPS receiver used principally to link it to the National Bureau of Standards timing system. The GPS is internally activated approximately every 4 hours and checks the "location of the instrument" and the time. Any differences between the internal clock and the GPS time are recorded and saved by the SMA. Any corrections to the internal timing are made automatically. Typically, the greatest correction recorded is approximately 4 milliseconds.

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<sup>&</sup>lt;sup>1</sup> The GPS antenna is mounted on the enclosure at the rear of the solar array.

Table 2.4. Instrument Parameters for the Kinemetrics ETNA<sup>TM</sup> System in the Hanford SMA Network

Parameter	Value or Range			
Sensor				
Туре	Tri-Axial Force Balance Accelerometer orthogonally oriented with internal standard			
Full-Scale	$\pm 2 g^{(a)}$			
Frequency Range	0–50 Hz			
Damping	Approximately 70% critical <sup>(a)</sup>			
Data Acquisition				
Number of Channels	3			
Sample Rate	18-bit resolution @ 200 samples/second			
Digital Output	Real-time, RS-232 Output Stream			
Seismic Trigger				
Filter	0.1–12.5 Hz			
Trigger level	$0.1\%g^{(b)}$			
Alarm (call-out) Threshold	Not activated			
Pre-Event Memory	10 sec			
Post-Event Time	40 sec			
(a) Setting is dependent on	instrument calibration.			
(b) See Section 2.2.4 for dis	cussion of trigger thresholds.			

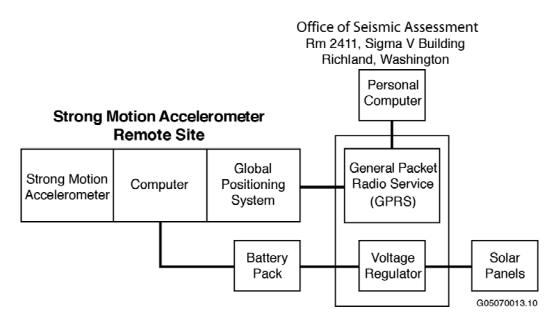


Figure 2.3. Schematic Diagram of a Strong Motion Accelerometer Installation

#### 2.2.3 Strong Motion Accelerometer Operations Center

The combined operations, data recording, data interpretation, and maintenance facility is located in the Sigma V Building and is operated by the HSA.

#### 2.2.4 Strong Motion Operational Characteristics

The signals from the three-accelerometer channels at each site are digitized with a 24-bit digitizer and temporarily stored in a memory buffer. The sampling rate of the digitizer is set to 200 samples/second. The three channels are monitored for signals that equal or exceed a programmable trigger threshold.

When one accelerometer channel is triggered, the other channels automatically record. The nominal threshold used is 0.05% of the full-scale range of 2.0 g (g is the acceleration of gravity, 9.8 m/s $^2$  or 32 ft/s $^2$ ) or 0.001 g. Threshold trigger levels are adjusted to trigger infrequently on the noise sources (e.g., vehicles, sonic booms) near each site. This provides ground motion data for smaller, non-damaging earthquakes that can be useful in estimating the ground motion expected from larger earthquakes, and to confirm correct operation of the instruments by analyzing the smaller-amplitude triggers. The recorders store information for 10 seconds before the trigger threshold is exceeded and for 40 seconds after the trigger ceases to be exceeded.

# 3.0 Earthquake Catalog Description

The Hanford Seismic Assessment staff uses an interactive program XPED developed at the University of Washington to determine earthquake locations and magnitudes. This program operates on the sections of time saved in files by the trigger algorithm of the Earthworm system. It provides the user with the ability to measure the arrival times and durations of seismic waves from earthquakes and determine the locations and magnitudes of the events. Locations of teleseismic and regional earthquakes are interpreted and saved for operational and quality review and documentation, and are not reported here. Local earthquakes near the Hanford Site (46°-47°N latitude, 119° -120°W longitude) are reported in this report (Table 3.1). Other earthquakes in southeast Washington are kept on file.

## 3.1 Coda Length Magnitude

Coda-length magnitude ( $M_c$ ), an estimate of local magnitude ( $M_L$ ) (Richter 1958), is calculated using the coda-length/magnitude relationship determined for Washington State by Crosson (1972).

This relationship is:

$$M_c = 2.82 \log (D) - 2.46$$
,

where D is the duration of the observed signal.

## 3.2 Velocity Model

The program XPED uses the velocities and layer depths given in Table 3.2. This model does not include a surficial layer for the Hanford formation or Ringold Formation sediments because most stations are located on basalt. Time corrections, which account for elevation, or local differences in the velocity model (i.e., stations on sedimentary layers), are determined empirically from sets of accurately-located earthquakes and explosions in the region.

# 3.3 Quality Factors (Q)

XPED assigns a two-letter **Quality factor** (Table 3.1) that indicates the general reliability of the solution (**A** is best quality, **D** is worst). Similar quality factors are used by the USGS for events located with the computer program HYPO71. The first letter of the quality code is a measure of the hypocenter quality based primarily on travel time residuals. For example: Quality **A** requires a root-mean-square residual (**RMS**) less than 0.15 seconds while a **RMS** of 0.5 seconds or more is **D** quality (other estimates of the location uncertainty also affect this quality parameter). The second letter of the quality code is related to the spatial distribution of stations that contribute to the event's location, including the number of stations (**NS**), the number of p-wave and s-wave phases (**NP**), the largest gap in event-station azimuth distribution (**GAP**), and the closest distance from the epicenter to a station (**DMIN**). Quality **A** requires a solution with **NP** >8, **GAP** <90°, and **DMIN** <5 km (or the hypocenter depth if it is greater than 5 km). If **NP** ≤5, **GAP** >180°, or **DMIN** >50 km, the solution is assigned Quality **D**.

Table 3.1. Local Seismic Data, October 1, 2005 to December 31, 2005

Event ID	Type	Date	Time	Latitude	Longitude	Depth	Mag	NS/NP	Gap	Dmin	RMS	Q
05100219330		2005/10/02	19:32:94.57	46N29.10	119W21.61	6.15	1.2	14/017	112	8 0.10	0.7	AB
05100603261		2005/10/06	03:26:39.65	46N29.08	119W21.29	5.32	1.2	10/013	142	8 0.11	1.2	AC
05101113302		2005/10/11	13:30:44.19	46N34.58	119W47.39	0.44	-0.2	7/012	203	4 0.17	0.5	BD
05101117422		2005/10/11	17:42:51.34	46N33.27	119W54.60	6.15	-0.2	5/007	196	13 0.30	3.8	CD
05101119433		2005/10/11	19:43:58.31	46N34.79	119W48.40	4.09	-0.2	7/011	205	5 0.14	1.4	AD
05101119484		2005/10/11	19:48:64.42	46N34.87	119W46.35	0.03*	-0.3	4/008	193	3 0.20	0.1	BD
05101203550		2005/10/12	03:55:31.02	46N37.30	119W45.42	14.18	-0.3	5/006	261	1 0.04	0.4	AD
05101209092		2005/10/12	09:09:50.61	46N33.39	119W44.32	10.09	-0.7	6/008	186	4 0.05	0.5	AD
05101504470		2005/10/15	04:46:87.52	46N42.77	119W30.03	1.06	-0.2	8/011	81	5 0.09	2.5	BB
05101615291		2005/10/16	15:29:40.09	46N34.65	119W48.58	0.04*	0.3	9/013	207	5 0.10	0.1	AD
05102021330	P	2005/10/20	21:32:91.62	46N11.39	119W13.75	0.03*		20/020	146	20 0.17	0.7	BC
05102917081		2005/10/29	17:08:30.08	46N41.57	119W39.82	7.90	-0.9	6/007	187	9 0.07	0.4	AD
05111012453		2005/11/10	12:45:59.25	46N08.30	119W55.75	11.45	2.6	34/037	71	20 0.18	0.9	BB
05111013052		2005/11/10	13:05:47.57	46N08.95	119W55.60	8.27*	1.2	14/014	113	19 0.06	0.3	AC
05111118533	P	2005/11/11	18:53:58.39	46N25.87	119W01.45	0.02*		9/009	151	17 0.28	0.3	BC
05112119580	P	2005/11/21	19:57:85.21	46N14.97	119W43.12	0.03*		9/009	212	4 0.19	1.0	BD
05120107082		2005/12/01	07:08:44.33	46N43.30	119W30.34	3.23	-0.7	7/011	89	5 0.11	1.0	AB
05120113050		2005/12/01	13:04:84.79	46N42.91	119W29.27	5.33	-0.9	3/005	225	4 0.04	0.5	AD
05120305012		2005/12/03	05:01:53.03	46N43.02	119W29.82	0.02*	0.5	8/009	77	5 0.10	0.0	AB
05120314524		2005/12/03	14:52:73.27	46N43.19	119W29.95	3.61	0.4	12/015	68	5 0.12	0.7	AB
05120314552		2005/12/03	14:55:44.73	46N43.23	119W30.13	5.99	-0.1	8/009	88	5 0.11	0.9	AA
05120319365		2005/12/03	19:36:76.35	46N43.17	119W30.02	4.39	-0.6	7/009	87	5 0.09	0.9	AB
05120323505		2005/12/03	23:50:80.94	46N43.53	119W29.88	1.71	-0.1	6/007	157	5 0.04	1.3	AC
05120404095		2005/12/04	04:09:76.06	46N43.52	119W30.04	4.83	0.5	10/010	99	5 0.12	1.1	AB

Table 3.1. (cont'd)

Event ID	Type	Date	Time	Latitude	Longitude	Depth	Mag	NS/NP	Gap	Dmin	RMS	Q
05120404482		2005/12/04	04:48:43.57	46N43.49	119W30.34	5.08	-0.2	7/009	92	5 0.09	1.0	AB
05120404551		2005/12/04	04:55:45.57	46N43.12	119W30.10	0.05*	1.2	16/017	70	5 0.09	0.1	AB
05120404591		2005/12/04	04:59:35.13	46N43.31	119W30.29	5.05	0.4	11/014	70	5 0.13	1.1	AB
05120405004		2005/12/04	04:59:22.68	46N43.27	119W29.90	0.04*	1.5	20/022	66	5 0.11	0.0	AB
05120405013		2005/12/04	04:59:53.08	46N43.59	119W30.07	0.46	-0.3	5/007	95	5 0.15	0.6	AD
05120405232		2005/12/04	05:23:50.22	46N43.20	119W30.20	3.00	0.0	8/010	87	5 0.12	1.6	AB
05120405512		2005/12/04	05:51:50.80	46N43.52	119W30.86	5.76	-0.8	7/007	91	5 0.15	1.7	AB
05120407205		2005/12/04	07:20:78.99	46N43.24	119W30.48	4.41	-0.4	7/008	87	5 0.12	1.1	AB
05120409360		2005/12/04	09:36:29.47	46N44.16	119W29.99	4.14	-0.7	6/010	178	5 0.14	1.3	AC
05120417210		2005/12/04	17:20:80.51	46N18.25	119W31.51	17.08	-0.3	9/013	233	6 0.06	0.5	AD
05120521125	P	2005/12/05	21:12:76.57	46N16.30	119W23.10	0.02*		12/012	127	4 0.09	0.2	AB
05120709163		2005/12/07	09:16:53.97	46N43.04	119W29.63	1.93	-0.3	8/011	86	4 0.13	1.5	AA
05120709552		2005/12/07	09:55:52.26	46N42.94	119W29.57	2.48	-0.4	7/009	85	4 0.10	1.1	AB
05120819204		2005/12/08	19:20:69.88	46N43.46	119W30.12	4.04	-0.2	7/010	92	5 0.12	1.0	AB
05120902400		2005/12/09	02:39:84.19	46N43.62	119W30.73	2.36	-0.2	8/009	152	5 0.16	3.3	BC
05120905403		2005/12/09	05:40:55.58	46N43.44	119W30.46	0.04*	-0.4	7/009	91	5 0.15	0.1	AB
05120911150		2005/12/09	11:14:89.95	46N43.33	119W30.18	4.60	0.3	9/012	95	5 0.11	0.9	AB
05120923194		2005/12/09	23:19:67.25	46N43.75	119W30.88	6.08	-0.6	5/009	95	4 0.14	1.6	AD
05121012090		2005/12/10	12:08:90.87	46N43.33	119W29.81	4.00	-0.2	7/008	91	5 0.07	1.0	AB
05121014510		2005/12/10	14:51:34.57	46N43.01	119W29.76	0.03*	0.6	14/015	68	4 0.13	0.1	AA
05121015262		2005/12/10	15:26:49.76	46N43.17	119W29.69	3.37	0.7	11/012	88	4 0.13	1.3	AA
05121015290		2005/12/10	15:28:85.29	46N43.67	119W30.69	6.46	-0.5	7/007	94	5 0.09	1.3	AB
05121015594		2005/12/10	15:59:67.09	46N43.25	119W30.64	4.45	-0.3	9/011	86	5 0.10	1.0	AB
05121515190		2005/12/15	15:18:88.90	46N42.93	119W29.77	3.24	0.1	10/013	84	5 0.07	0.5	AB

Table 3.1. (cont'd)

Event ID	Type	Date	Time	Latitude	Longitude	Depth	Mag	NS/NP	Gap	Dmin	RMS	Q
05121705165		2005/12/17	05:16:74.52	46N29.19	119W21.27	5.91	-0.5	6/009	174	8 0.05	0.4	AC
05121708131		2005/12/17	08:13:37.47	46N42.85	119W30.12	2.11	0.3	10/013	93	5 0.12	1.6	AB
05121802155		2005/12/18	02:15:78.01	46N43.51	119W30.79	6.20	-0.8	6/007	95	5 0.12	1.7	AC
05121803540		2005/12/18	03:54:26.31	46N43.01	119W30.30	3.57	0.0	7/009	84	5 0.05	0.3	AB
05121902281		2005/12/19	02:28:41.12	46N34.49	119W41.39	7.74	-1.1	4/006	109	6 0.03	1.0	AD
05121903221		2005/12/19	03:22:44.02	46N34.16	119W41.44	7.99	-0.8	3/006	152	5 0.02	0.2	AD
05122009324		2005/12/20	09:32:65.53	46N43.11	119W30.09	3.03	-0.2	8/012	86	5 0.11	0.9	AB
05122312412		2005/12/23	12:41:52.87	46N26.47	119W35.57	16.48	-0.3	9/011	117	5 0.06	0.8	AB
05122800143		2005/12/28	00:14:61.17	46N29.12	119W21.16	0.14	-0.4	5/007	144	7 0.16	1.5	BD
05122807352		2005/12/28	07:35:43.28	46N43.50	119W30.45	3.42	-0.4	6/010	151	5 0.17	2.3	BC
05122810593		2005/12/28	10:59:57.45	46N43.36	119W30.02	4.06	0.6	9/013	70	5 0.07	0.5	AB
05122910293		2005/12/29	10:29:60.92	46N28.78	119W22.29	4.21	1.0	11/015	104	8 0.11	1.4	AB
05123012341		2005/12/30	12:34:40.46	46N43.19	119W30.07	3.14	-0.3	6/007	90	5 0.11	2.0	AC

#### **Explanation of Table 3.1**

**Event ID**: The Earthworm Recording System creates the identification number. XPED uses the year,

month, day, and time to create a unique number for each event.

**Type**: P is Probable Blast; X is Confirmed Blast; F is Felt Earthquake; blank is local earthquake.

**Date**: The year and day of the year in Universal Time Coordinated (UTC). UTC is used

throughout this report unless otherwise indicated.

**Time**: The origin time of the earthquake given in UTC. To covert UTC to Pacific Standard Time,

subtract eight hours; to Pacific Daylight Time, subtract seven hours.

**Latitude**: North latitude, in degrees and minutes, of the earthquake epicenter.

**Longitude**: West longitude, in degrees and minutes, of the earthquake epicenter.

**Depth**: The depth of the earthquake in kilometers (km).

Mag: The magnitude is expressed as Coda-Length magnitude M<sub>c</sub>, an estimate of local magnitude

M<sub>L</sub> (Richter 1958). If magnitude is blank, a determination was not made.

**NS/NP**: Number of stations/number of phases used in the solutions.

**Gap:** Azimuthal gap. The largest angle (relative to the epicenter) containing no stations.

**DMIN**: The distance from the earthquake epicenter to the closest station.

**RMS**: The root-mean-square residual (observed arrival times minus the predicted arrival times) at

all stations used to locate the earthquake. It is only useful as a measure of quality of the solution when five or more well-distributed stations are used in the solution. Good solutions are normally characterized by RMS values of less than about 0.3 seconds.

**Q**: The Quality Factors indicate the general reliability of the solution/location (A is best

quality, D is worst). See Section 3.3 of this report, "Quality Factors."

Table 3.2. Seismic Velocities for Columbia Basin Stratigraphy (from Rohay et al. 1985)

Depth to Top of Velocity Layer (km)	Stratigraphy	Velocity (km/sec)
0.0	Saddle Mountains and Wanapum Basalts and intercalated Ellensburg Formation	3.7
0.4	Grande Ronde Basalt and pre-basalt sediments	5.2
8.5	Crystalline Basement, Layer 1	6.1
13.0	Crystalline Basement, Layer 2	6.4
23.0	Sub-basement	7.1
38.0	Mantle	7.9

# 4.0 Geology and Tectonic Analysis

The Hanford Site lies within the Columbia Basin, which is an intermontane basin between the Cascade Range and the Rocky Mountains that is filled with Cenozoic volcanic rocks and sediments. This basin forms the northern part of the Columbia Plateau physiographic province (Fenneman 1931) and the Columbia River flood-basalt province (Reidel and Hooper 1989). In the central and western parts of the Columbia Basin, the Columbia River Basalt Group (CRBG) overlies Tertiary continental sedimentary rocks and is overlain by late Tertiary and Quaternary fluvial and glaciofluvial deposits (Campbell 1989; Reidel et al. 1989, 1994; DOE 1988). In the eastern part, a thin (<100 m) sedimentary unit separates the basalt and underling crystalline basement and a thin (<10 m) veneer of eolian sediments overlies the basalt (Reidel et al. 1989, 1994).

The Columbia Basin has two structural subdivisions or subprovinces: the Yakima Fold Belt and the Palouse Slope. The Yakima Fold Belt includes the western and central parts of the Columbia Basin and is a series of anticlinal ridges and synclinal valleys with major thrust faults along the northern flanks (Figure 4.1) (Reidel and Fecht 1994a, 1994b). The Palouse Slope is the eastern part of the basin and is less deformed than the Yakima Fold Belt with only a few faults and low amplitude, long wavelength folds on an otherwise gently westward dipping paleoslope. Figure 4.2 shows north-south (B-B') and east-west (A-A') cross sections through the Columbia Basin based on surface mapping (Reidel and Fecht 1994a, 1994b), deep boreholes (Reidel et al. 1994), geophysical data (Rohay et al. 1985; DOE 1988), and magnetotelluric data obtained as part of BWIP (DOE 1988).

## 4.1 Earthquake Stratigraphy

Studies of seismicity at the Hanford Site have shown that the seismic activity is related to crustal stratigraphy (layers of rock types) (Rohay et al. 1985; DOE 1988). The main geologic units important to earthquakes at the Hanford Site and the surrounding area are:

- The Miocene Columbia River Basalt Group (CRBG)
- Pre-basalt sediments of Paleocene, Eocene, Oligocene, and Early Miocene age
- The crystalline basement consisting of two layers composed of Precambrian and Paleozoic craton
- Mesozoic accreted terranes.

# 4.2 Geologic Structure Beneath the Monitored Area

Between the late 1950s and the mid 1980s, deep boreholes were drilled for hydrocarbon exploration in the Columbia Basin. These boreholes provided accurate measurements of the physical properties of the CRBG and the pre-basalt sediments (Reidel et al. 1989, 1994), but the thickness of the pre-basalt sediments and nature of the crystalline basement are still poorly understood. The difference between the thicknesses listed in Table 4.1 and the thicknesses of the crustal layers in the velocity model in Table 3.1 reflect data specific to UW's crustal velocity model for eastern Washington. Table 4.1 is derived from Reidel et al. (1994) and was developed for the geologic interpretation in this report. The thicknesses of these units are variable across the monitored area. Table 4.1 summarizes the approximate thickness at the borders of the monitored area.

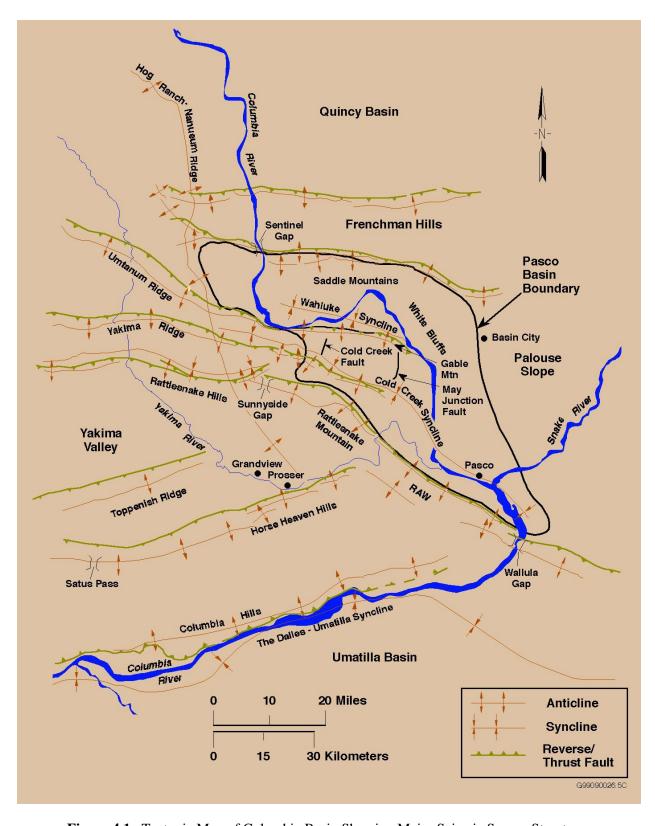


Figure 4.1. Tectonic Map of Columbia Basin Showing Major Seismic Source Structures

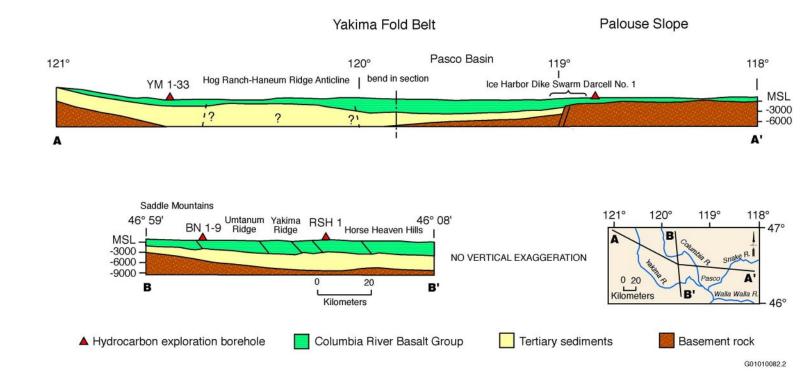


Figure 4.2. Geologic Cross Sections Through the Columbia Basin (Reidel et al. 1994)

**Table 4.1.** Thicknesses of Stratigraphic Units in the Monitoring Area (from Reidel et al. 1994)

Stratigraphy	North	South	East	West
Columbia River Basalt Group (includes suprabasalt sediments)	3.0 km	4.5 km	2.2 km	4.2 km
Pre-Basalt Sediments	3.0 km	>4.5 km	0	>6.0 km

The thickness of the basalt and the pre-basalt sediments varies as a result of different tectonic environments. The western edge of the North American craton (late Precambrian/Paleozoic continental margin and Precambrian craton) is located east of the Pasco Basin. The stratigraphy on the craton consists of CRBG overlying crystalline basement; the crystalline basement is continental crustal rock that underlies much of the western North America. The stratigraphy west of the craton consists of two parts. The exact western edge of the North American craton in Washington is not known. Between the Pasco Basin and the craton, there is about 4 to 5 km of CRBG overlying a thin section of early Cenozoic sediments. This early Cenozoic sediment lies on rocks of accreted terranes of Mesozoic age. Directly underlying the Pasco Basin is 4 to 5 km of CRBG overlying at least 6 km of pre-basalt sediments. This in turn overlies what is assumed to be accreted terranes. The area underlying the Pasco Basin was subsiding during the Eocene and Oligocene, accumulating great thickness of pre-CRBG sediments. Continued subsidence in this area during the Miocene resulted in thicker CRBG compared to that on the craton. Subsidence continues today but at a greatly reduced rate (Reidel et al. 1994).

#### 4.3 Tectonic Pattern

Studies have concluded that earthquakes can occur in the following six different tectonic environments (earthquake sources) at the Hanford Site (Geomatrix 1996).

- Major Geologic Structures. Reverse/thrust faults in the CRBG associated with major anticlinal ridges such as Rattlesnake Mountain, Yakima Ridge, and Umtanum Ridge could produce some of the largest earthquakes.
- Secondary faults. These faults are typically smaller (1 to 20 km) than the main reverse/thrust faults that occur along the major anticlinal ridges (up to 100 km). Secondary faults can be segment boundaries (tear faults) and small faults of any orientation that formed along with the main structure.
- Swarm areas. Small geographic areas not known to contain any geologic structures produce clusters of events (swarms), usually in the CRBG in synclinal valleys. These clusters consist of a series of small shocks with no outstanding principal event. Swarms occur over a period of days or months and the events may number into the hundreds and then quit, only to start again at a later date. This differs from the sequence of foreshocks, mainshock, and trailing-off aftershocks that have the same epicenter or are associated with the same fault system. In the past, swarms were thought to occur only in the CRBG. Most swarm areas are in the basalt but swarm events also appear to occur in all geologic layers. However, typically a swarm event at a specific time is usually restricted to one layer. There are seven earthquake swarm areas that are recognized in the Hanford Seismic Network area but this list will be updated as new swarm areas develop. The Saddle Mountains swarm area, Wooded Island swarm area, Wahluke swarm area, Coyote Rapids swarm area, and Horse Heaven Hills swarm area are typically active at one time or another during the year. The other earthquake swarm areas are active less frequently.

- The entire Columbia Basin. The entire basin, including the Hanford Site, could produce a "floating" earthquake. A floating earthquake is one that, for seismic design purposes, can happen anywhere in a tectonic province and is not associated with any known geologic structure. Seismic interpretation classifies it as a random event for purposes of seismic design and vibratory ground motion studies.
- Basement source structures. Studies (Geomatrix 1996) suggest that major earthquakes can
  originate in tectonic structures in the crystalline basement. Because little is known about geologic
  structures in the crystalline basement beneath the Hanford Site, earthquakes cannot be directly tied
  to a mapped fault. Earthquakes occurring in the crystalline basement without known sources are
  treated as random events.
- The Cascadia Subduction Zone. This source has been postulated to be capable of producing a magnitude 9 earthquake. Because this source is along the western boundary of Washington State and outside the HSN, the Cascadia Subduction Zone is not an earthquake source that is monitored at the Hanford Site, so subduction zone earthquakes are not reported here. Because any earthquake along the Cascadia Subduction zone can have a significant impact on the Hanford Site, or can be felt like the February 2001 Nisqually earthquake, UW monitors and reports on this earthquake source for the DOE. Ground motion from any moderate or larger Cascadia Subduction Zone earthquake is detected by Hanford SMAs and reported (see Section 5.0).

## 4.4 Depth of Earthquakes

Since records have been kept, most of the earthquakes at the Hanford Site have originated in the CRBG layer. The crystalline basement has had the next greatest amount of earthquakes followed by the pre-basalt sediments. The stratigraphic units for local earthquakes recorded during the first quarter of FY 2006 are listed in Table 4.2.

Unit	First Quarter	Second Quarter	Third Quarter	Fourth Quarter	FY 2005
Basalt	46	-	-	-	46
Pre-Basalt Sediments	6	-	-	-	6
Crystalline Basement	5	-	-	-	5
Total	57	-	-	-	57

**Table 4.2**. Number of Local Earthquakes Occurring in Stratigraphic Units

# 4.5 Tectonic Activity

### 4.5.1 Summary

During the first quarter of FY 2006, there were 57 earthquakes that occurred within the Hanford Seismic Monitoring Network (Table 4.3; Figure 4.3). Forty six (81%) events occurred in the CRBG; six (10%) occurred in the pre-basalt sediments; and 5 (9%) events occurred in the crystalline basement. No events are interpreted as occurring along major geologic structures; 50 (88%) events are interpreted as occurring in earthquake swarms; and seven (12%) events are interpreted as random events.

Table 4.3. Summary of Earthquake Locations for FY 2005

Seismic Sources		First Quarter 10/01- 12/31	Second Quarter 1/01 - 3/31	Third Quarter 4/01 - 6/30	Fourth Quarter 7/01 - 9/30	FY 2005
Geologic	Structure	-	-	-	-	-
	Frenchman Hills	-	-	-	-	-
	Saddle Mountains/ Royal Slope	-	-	-	-	-
	Wahluke Slope	1	-	-	-	1
	Coyote Rapids	39	-	-	-	39
C	Wye	4	-	-	-	4
Swarm Areas	Wooded Island	-	-	-	-	-
	Cold Creek	6	-	-	-	6
	Rattlesnake Mt.	-	-	-	-	-
	Horse Heaven Hills	-	-	-	-	-
	Total for swarm areas	50	-	-	-	50
Random Events		7	-	-	-	7
Total for	all earthquakes	57	-	-	-	57

### 4.5.2 First Quarter Earthquakes of FY 2006

Fifty seven earthquakes occurred in the Hanford area during the first quarter of FY 2006 (October 1, 2005 through December 31, 2005 [Figure 4.3]) (Table 4.3).

#### 4.5.2.1 Location and Depth of Earthquakes

### 4.5.2.1.1 Major Anticlinal Ridges

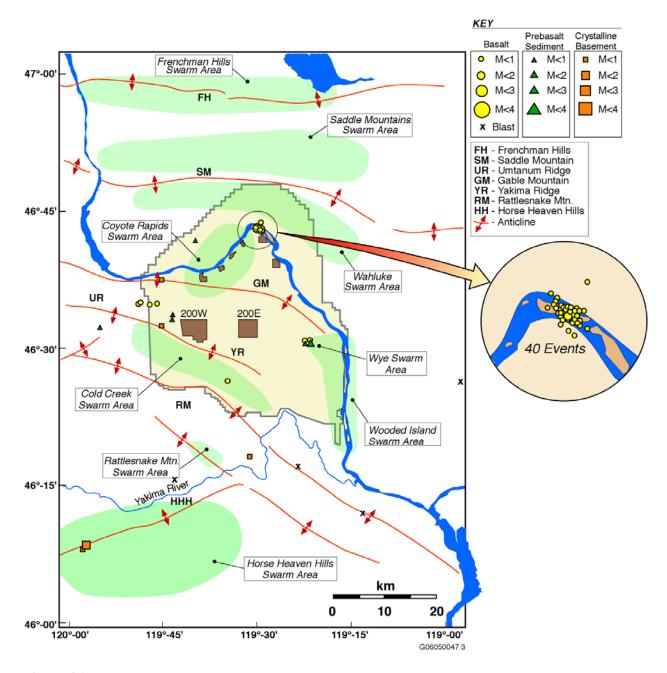
During the first quarter of FY 2006, we interpret no seismic events to have occurred on major geologic structures.

### 4.5.2.1.2 Earthquake Swarm Areas

Fifty earthquakes were located in swarm events (Figure 4.3).

#### Coyote Rapids Swarm Area

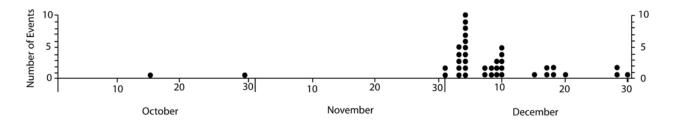
Thirty nine earthquakes occurred at the northeastern end of the Coyote Rapids swarm area starting on October 15, 2005 and continuing into 2006. All events were less than  $1.0~M_{\rm c}$  and all occurred in the basalt. The earthquakes were concentrated at the 'horn' of the Columbia River (inset Figure 4.3). Most of the earthquakes occurred in December (Figure 4.4).



**Figure 4.3**. All Earthquakes Recorded Within the Hanford Monitoring Area during the First Quarter of FY 2006 (October 1, 2005-December 31, 2005). North is top of figure.

#### Wahluke Swarm Area

One small (approximately  $0.0~M_{\rm c}$ ) earthquake occurred in the Wahluke Swarm Area during the first quarter of FY 2006. This earthquake occurred on October 29 and was in the pre-basalt sediments. This event was west of the Coyote Rapids earthquake swarm.



**Figure 4.4**. Timing of Earthquakes in the Coyote Rapids Earthquake Swarm during the First Quarter of FY 2006.

#### Wye Swarm Area

Four earthquakes occurred in the Wye Barricade Swarm Area during the first quarter of FY 2006. The first two occurred on October 2 and October 6 and had magnitudes of  $1.2\,M_c$ . These two events occurred in the pre-basalt sediments. The next earthquake occurred on December 28. This event was small (approximately  $0.0M_c$ ) and occurred in the basalt. The final earthquake to occur in this swarm area during the first quarter was on December 29. It also occurred in the basalt and had a magnitude of  $1.0\,M_c$ .

#### Cold Creek Swarm Area

Six earthquakes occurred in the western end of the Cold Creek Swarm Area during the first quarter of FY 2006. Four of the earthquakes occurred on October 11 and were in the basalt or near the base of the basalt. All four events were small (approximately  $0.0~M_c$ ). The next earthquake to occur was on October 12. This event occurred in the crystalline basement and was small (approximately  $0.0~M_c$ ). The last earthquake to occur in the swarm area during the first quarter was on October 16. This event occurred in the basalt and was small approximately  $0.0~M_c$ ).

#### 4.5.2.1.3 Random or Floating Events

During the first quarter of FY 2006, we interpret seven random events to have occurred in the monitoring area. Events that occurred in the pre-basalt sediments and in the crystalline basement are typically classified as random events. This is because there are no known geologic structures that have been identified in the rocks that occur below the Columbia River Basalt Group. However, we now recognize that some events that occur at depths that places them in the pre-basalt sediments and crystalline basement occur in patterns that fit our definition of earthquake swarms (Section 4.3). Those events are now reported in the appropriate sections on earthquake swarms.

The first random event occurred on October 12 beneath Umtanum Ridge. This was a very small event (approximately  $0.0 \text{ M}_c$ ) that occurred in the crystalline basement.

Two random events occurred on November 10 on the Horse Heaven Hills southwest of Prosser, Washington. Both were deep events (>8 km) and had magnitudes of 2.6  $M_c$  and 1.2  $M_c$ , respectively.

On December 4, a small (approximately  $0.0~M_{\rm c}$ ) earthquake occurred in the crystalline basement (>17 km depth) southeast of the crest of Rattlesnake Mountain.

Two small (approximately  $0.0\ M_{\rm c}$ ) random earthquakes occurred in the pre-basalt sediment beneath Cold Creek north of Yakima Ridge on December 19.

The last random event for the first quarter of FY 2006 occurred north of Rattlesnake Mountain. It was small (approximately 0.0 Mc) and in the basalt. Because this earthquake is north of the Rattlesnake Mountain fault and not near any mapped fault, we classify this event as random.

# **5.0 Strong Motion Accelerometer Operations**

The Hanford SMA network has been in continuous operation since November 20, 1998. The nominal threshold used in the SMA network is 0.001 g in order to provide ground motion for smaller, non-damaging earthquakes that can be useful in estimating the ground motion expected from larger earthquakes, and to confirm correct operation of the instruments by analyzing the smaller-amplitude triggers (see Section 2.2).

# 5.1 FY 2006 Triggers of the Hanford SMA Network

The Hanford SMA network did not trigger from a seismic event during the first quarter of FY 2006.

# 6.0 Capabilities in the Event of a Significant Earthquake

The SMA network was designed to provide ground motion data in areas at the Hanford Site that have high densities of people and/or facilities containing hazardous materials in order to insure the Hanford Site is in compliance with DOE Order 420.1, "Facility Safety." The network also allows the HSAP to support Hanford Site emergency services organizations in complying with DOE Order G 420.1-1, Section 4.7, "Emergency Preparedness and Emergency Communications," by providing area ground motion data in the event of an earthquake on the Hanford Site. This section summarizes the capabilities of the HSAT in the event of an earthquake at Hanford.

## 6.1 Use of the SMA Network in the Event of an Earthquake

Historically, only a few facilities at the Hanford Site had instruments to provide data on peak ground accelerations or any type of ground motion. The present SMA instruments were located so that if an earthquake occurred, ground motion data would be readily available to assess the damage at the 100-K Area, the 200-East and West Areas, and the 300 and 400 Area facilities, which have the greatest concentration of people and also contain hazardous materials (Moore and Reidel 1996).

Many facilities at the Hanford Site have undergone various degrees of seismic analysis either during design or during re-qualification. Although the seismic design of a building may be known, when an earthquake is "felt" in a facility on the Hanford Site, a determination must be made as to the extent of damage before it can be re-occupied and the systems restarted. A "felt" earthquake may not cause any significant damage to a building but, without adequate characterization of the ground motion, initial determination of the building's possibility of having damage may be impossible.

In the event of an earthquake such as the 2001 Nisqually earthquake, building managers, emergency directors, and engineers can obtain ground motion data recorded by the SMA network from the HSAT in the Sigma V Building. This is done through the Hanford Site Emergency Services Organization. Normal hours of operation for the HSAP are between 6 a.m. and 4:30 p.m., Monday through Friday. If a SMA is triggered, the HSAT will download events that were recorded and determine the peak ground accelerations. This information is then passed on to Hanford Emergency Services personnel where the facility engineers can use the data to determine if the ground motion exceeded, is equal to, or is less than the building design. This, along with assessments from trained engineers, allows the facility manager to make a rapid and cost-effective determination on whether a building is safe to re-occupy or should not be used until it has been inspected in more detail. Buildings that have designs exceeding the recorded ground motion could be put back into service very quickly; buildings with designs that are very close to or less than measured ground motion could be given priority for onsite damage inspections.

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