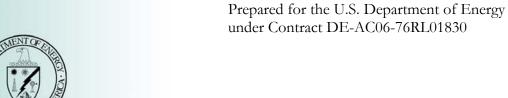


Annual Hanford Seismic Report for Fiscal Year 2002

PNNL Seismic Monitoring Team

November 2002





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Annual Hanford Seismic Report for Fiscal Year 2002

Pacific Northwest National Laboratory Seismic Monitoring Team

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November 2002

Prepared for the U.S. Department of Energy under Contract DE-AC06-76RL01830

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Summary

Hanford Seismic Monitoring provides an uninterrupted collection of high-quality raw and processed seismic data from the Hanford Seismic Network for the U.S. Department of Energy and its contractors. Hanford Seismic Monitoring also locates and identifies sources of seismic activity and monitors changes in the historical pattern of seismic activity at the Hanford Site. The data are compiled, archived, and published for use by the Hanford Site for waste management, Natural Phenomena Hazards assessments, and engineering design and construction. In addition, the seismic monitoring organization works with the Hanford Site Emergency Services Organization to provide assistance in the event of a significant earthquake on the Hanford Site.

The Hanford Seismic Network and the Eastern Washington Regional Network consist of 41 individual sensor sites and 15 radio relay sites maintained by the Hanford Seismic Monitoring staff.

For the Hanford Seismic Network, there were 1,177 triggers during fiscal year 2002. Of these triggers, 553 were earthquakes.

Two earthquakes were located in the Hanford Seismic Network area. Stratigraphically 13 occurred in the Columbia River basalt, 12 were earthquakes in the pre-basalt sediments, and 17 were earthquakes in the crystalline basement. Geographically, 13 earthquakes occurred in swarm areas, 1 earthquake was associated with major structures, and 28 were random events.

There were no earthquake triggers of the Hanford Strong Motion Accelerometers during fiscal year 2002.

Acronyms

BWIP Basalt Waste Isolation Project
CDPD Cellular Digital Packet Data
CRBG Columbia River Basalt Group

DMIN closest distance from the epicenter to a station

DOE U.S. Department of Energy

ETNA strong motion accelerometer manufactured by Kinemetrics

EWRN Eastern Washington Regional Network

FY fiscal year

GAP largest gap in event-station azimuth distribution

GPS Global Positioning System
HSN Hanford Seismic Network
Mc Coda Length Magnitude

 M_L Local Magnitude M_w Moment Magnitude

NP number of p-wave and s-wave phases

NS number of stations

PNNL Pacific Northwest National Laboratory
RAW Rattlesnake Mountain-Wallula Alignment

RMS root-mean-square residual
SMA strong motion accelerometer
USGS United States Geological Survey
UTC Universal Time, Coordinated
UW University of Washington

WHC Westinghouse Hanford Company

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1.0 Introduction

This report is the annual Hanford seismic activity report for fiscal year (FY) 2002. The report includes earthquake activity that occurred on the Hanford Site and vicinity between October 1, 2001 and September 30, 2002 and the geologic interpretation of the sources of the earthquakes.

1.1 Mission

The principal mission of Hanford Seismic Monitoring at the Hanford Site is to insure compliance with DOE Order 420.1, "Facility Safety" and DOE Order G 420.1-1, Section 4.7, "Emergency Preparedness and Emergency Communications." DOE Order 420.1 establishes facility safety requirements related to nuclear safety design, criticality safety, fire protection, and natural phenomena hazards mitigation. For seismic monitoring, this order states:

4.4.5 Natural Phenomena Detection.

Facilities or sites with hazardous materials shall have instrumentation or other means to detect and record the occurrence and severity of seismic events.

The Seismic Monitoring Project supports Hanford Site emergency services organizations in complying with DOE Order G 420.1-1, Section 4.7, "Emergency Preparedness and Emergency Communications," by providing assistance in the event of an earthquake on the Hanford Site.

In addition, seismic monitoring provides an uninterrupted collection of high-quality raw seismic data from the Hanford Seismic Network (HSN) located on and around the Hanford Site, and provides interpretations of seismic events from the Hanford Site and vicinity. Hanford Seismic Monitoring locates and identifies sources of seismic activity, monitors changes in the historical pattern of seismic activity at the Hanford Site, and builds a "local" earthquake database (processed data) that is permanently archived. The focus of this report is the precise location of earthquakes proximal to or on the Hanford Site, specifically between 46 degrees and 47 degrees north latitude and between 119 degrees and 120 degrees west longitude. Data from the Eastern Washington Regional Network (EWRN) and other seismic networks in the northwest provide the Seismic Monitoring Project with necessary regional input for the seismic hazards analysis at the Hanford Site. These seismic data are used to support Hanford Site contractors for waste management activities, Natural Phenomena Hazards assessments, and engineering design and construction.

1.2 History of Seismic Monitoring at Hanford

Seismic monitoring at the Hanford Site was established in 1969 by the United States Geological Survey (USGS) under a contract with the U.S. Atomic Energy Commission. In 1975, the University of Washington (UW) assumed responsibility for the network and subsequently expanded it. In 1979, the Basalt Waste Isolation Project (BWIP) became responsible for collecting seismic data for the Hanford Site as part of site characterization activities. Rockwell Hanford Operations, followed by Westinghouse Hanford

Company (WHC), operated the local network and were the contract technical advisors for the EWRN operated and maintained by UW. Funding ended for BWIP in December 1988. Seismic monitoring and responsibility for the UW contract were then transferred to WHC's Environmental Division. Maintenance responsibilities for the EWRN also were assigned to WHC who made major upgrades to EWRN sites.

Effective October 1, 1996, seismic monitoring was transferred to the Pacific Northwest National Laboratory (PNNL). (a) Seismic monitoring is part of PNNL's Applied Geology and Geochemistry Group, Environmental Technology Division.

The Hanford Strong Motion Accelerometer (SMA) network was constructed during 1997 and came on line in May 1997. It operated continuously until September 30, 1997 when it was mothballed due to lack of funding. Funding was restored on October 1, 1998 by joint agreement between the U.S. Department of Energy (DOE) and PNNL. Operation of the free-field sites resumed on November 20, 1999 and has operated continuously since that time.

1.3 Documentation and Reports

The Seismic Monitoring Project generally issues quarterly reports of local activity, an annual catalog of earthquake activity on and near the Hanford Site, and special-interest bulletins on local seismic events. The annual catalog includes the fourth quarter report for the fiscal year. During FY 2002 only this annual report is issued; no quarterly reports were issued. Hanford Seismic Monitoring also provides information and special reports to other functions as requested. Earthquake information provided in these reports is subject to revisions if new data become available. In addition, an archive of all seismic data from the HSN is maintained by PNNL.

1.2

⁽a) Pacific Northwest National Laboratory is operated by Battelle Memorial Institute for the U.S. Department of Energy.

2.0 Network Operations

2.1 Seismometer Sites

The seismic monitoring network consists of two designs of equipment and sites: seismometer sites and strong motion accelerometer (SMA) sites. Seismometer sites are designed to locate earthquakes and determine their magnitude and hypocenter location. SMA sites are designed to measure ground motion.

The HSN and the EWRN consist of 41 sensor sites. Most sites are in remote locations and require solar panels and batteries for power. The HSN uses 21 sites (Table 2.1 and Figure 2.1) and the EWRN uses 35 sites (Table 2.2 and Figure 2.2); both networks share 16 sites. The networks have 45 combined data channels because Gable Butte and Frenchman Hills East are three-component sites, each consisting of one vertical, one north-south horizontal, and one east-west horizontal data channel. Both networks use 15 additional telemetry relay sites. Data from all sites or relays are transmitted to the Sigma V building, Richland, Washington.

2.1.1 Station Maintenance

The HSN's maintenance records for the seismic sensor and relay sites are on file in the Hanford Seismic Monitoring office, Sigma V Building, Richland Washington.

2.1.2 Data Acquisition

The signals from the seismometer sites are monitored for changes in signal amplitude that are expected from earthquakes. The seismic network is subdivided into spatial groupings of stations that are monitored for nearly simultaneous amplitude changes, resulting in triggering a permanent recording of the events. The groupings and associated weighting schemes are designed to allow very small seismic events to be recorded and to minimize false triggers. Events are classified as local (south-central Washington near the Hanford Site), regional (Western U.S. and Canada), and teleseisms (from farther distances around the world). Local and regional events are usually earthquakes, but quarry and mining explosions are also recorded. Quarry and mining explosions can usually be identified from wave characteristics, time of day, and through confirmation with local government agencies and industries. Frequently, military exercises at the U.S. Army's Yakima Training Center produce a series of acoustic shocks that unavoidably trigger the recording system. Sonic booms and thunder also produce acoustic signals that trigger the recording system.

A PC-based system adapted from a USGS program and the UW system was implemented at Hanford during FY 1999. One system has been in continuous operation since January 6, 1999. A second, backup PC system was installed in mid-March 1999, and both systems have been running in parallel since that time. The hardware and software have been periodically upgraded. Data from triggers are collected on a SUNTM (registered trademark of Sun Microsystems, Santa Clara, California) workstation that is used to

Table 2.1. Seismic Stations in the Hanford Seismic Network

The first column is the three-letter seismic station designator. The latitude and longitude, elevation above sea level in meters, and the full station name follow this. The locations of the stations are all in Washington; locations were derived from a Global Positioning System (GPS).

	Latitude	Longitude		
Station	Deg. Min. N	Deg. Min. W	Elevation (m)	Station Name
BEN	46N31.13	119W43.02	340	Benson Ranch
BRV	46N49.12	119W59.47	920	Black Rock Valley
BVW	46N48.66	119W52.99	670	Beverly
CRF	46N49.50	119W23.22	189	Corfu
ET3	46N34.64	118W56.25	286	Eltopia Three
*FHE	46N57.11	119W29.82	455	Frenchman Hills East
*GBB	46N36.49	119W37.62	177	Gable Butte
GBL	46N35.92	119W27.58	330	Gable Mountain
H2O	46N23.75	119W25.38	158	Water
LOC	46N43.02	119W25.85	210	Locke Island
MDW	46N36.79	119W45.66	330	Midway
MJ2	46N33.45	119W21.54	146	May Junction Two
OT3	46N40.14	119W13.98	322	Othello Three
PRO	46N12.73	119W41.15	550	Prosser
RED	46N17.92	119W26.30	366	Red Mountain
RSW	46N23.67	119W35.48	1,045	Rattlesnake Mountain
SNI	46N27.85	119W39.60	312	Snively Ranch
WA2	46N45.32	119W33.94	244	Wahluke Slope
WIW	46N25.76	119W17.26	128	Wooded Island
WRD	46N58.20	119W08.69	375	Warden
YPT	46N02.93	118W57.73	325	Yellepit
* Three-comp	ponent station.			

determine earthquake locations and magnitudes (Section 3). Although the two systems are practically identical, there is enough granularity in the trigger timing that they sometimes record exclusive events. In nearly all cases, these exclusive triggers are "false" triggers, not earthquakes or quarry blasts (i.e., from acoustic sources). The remainders are from barely detectable, small signals from regional and teleseismic earthquakes.

The types and numbers of triggers recorded in FY 2002 by the seismic acquisition system are summarized in Table 2.3.

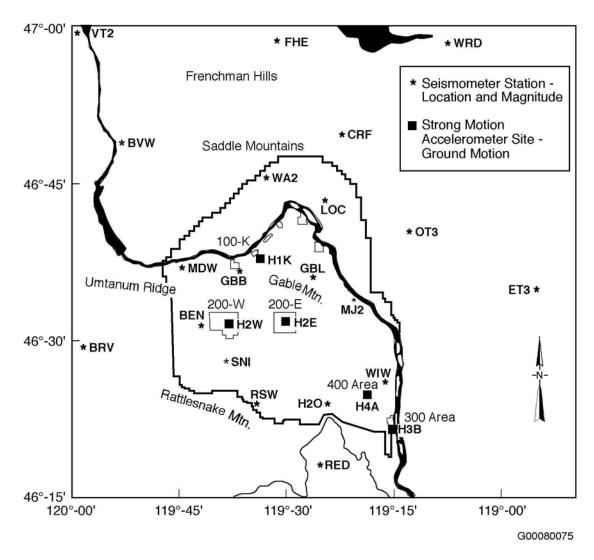


Figure 2.1. Locations of Seismograph Stations and Strong Motion Accelerometer Sites in the Hanford Seismic Network (see Table 2.1 for description of locations). Locations for Prosser (PRO) and Yellepit (YPT) are not shown. See Figure 2.2 for the locations of those sites.

2.2 Strong Motion Accelerometer Sites

2.2.1 Location

The Hanford SMA network consists of five free-field SMA sites (see Figure 2.1) (Table 2.4). There is one free-field SMA located in each of the 200 Separations Areas, one adjacent to the K Basins in the 100-K Area, one adjacent to the 400 Area where the Fast Flux Test Reactor is located, and one at the south end of the 300 Area.

Table 2.2. Seismic Stations in the Eastern Washington Regional Network

The first column is the three-letter seismic station designator. The latitude and longitude, elevation above sea level in meters, and the full station name follow this. The locations of the stations are all in Washington unless otherwise indicated; locations were determined from a Global Positioning System (GPS).

C4 - 4*	Latitude	Longitude	Floor 4' ()	CALA LOS NICOS
Station	Deg. Min. N.	Deg. Min. W.	Elevation (m)	Station Name
BRV	46N29.12	119W59.47	920	Black Rock Valley
BVW	46N48.66	119W52.99	670	Beverly
CBS	47N48.26	120W02.50	1,067	Chelan Butte, South
CRF	46N49.50	119W23.22	189	Corfu
DPW	47N52.25	118W12.17	892	Davenport
DY2	47N59.11	119W46.28	890	Dyer Hill Two
ELL	46N54.58	120W33.98	789	Ellensburg
EPH	47N21.38	119W35.76	661	Ephrata
ET3	46N34.64	118W56.25	286	Eltopia Three
ETW	47N36.26	120W19.94	1,477	Entiat
*FHE	46N57.11	119W29.82	455	Frenchman Hills East
*GBL	46N35.92	119W27.58	330	Gable Mountain
LNO	45N52.31	118W17.11	771	Lincton Mountain, Oregon
LOC	46N43.02	119W25.85	210	Locke Island
MDW	46N36.79	119W45.66	330	Midway
MJ2	46N33.45	119W21.54	146	May Junction Two
MOX	46N34.64	120W17.89	501	Moxee City
NAC	46N43.99	120W49.42	728	Naches
NEL	48N04.21	120W20.41	1,500	Nelson Butte
OD2	47N23.26	118W42.58	553	Odessa Two
OT3	46N40.14	119W13.98	322	Othello Three
PAT	45N52.92	119W45.14	262	Paterson
PRO	46N12.73	119W41.15	550	Prosser
RSW	46N23.67	119W35.48	1,045	Rattlesnake Mountain
SAW	47N42.10	119W24.03	701	St. Andrews
TBM	47N10.20	120W35.88	1,006	Table Mountain
TRW	46N17.32	120W32.31	723	Toppenish Ridge
TWW	47N08.29	120W52.10	1,027	Teanaway
VT2	46N58.04	119W58.95	1,270	Vantage Two
WA2	46N45.32	119W33.94	244	Wahluke Slope Two
WAT	47N41.92	119W57.24	821	Waterville
WIW	46N25.76	119W17.26	128	Wooded Island
WRD	46N58.20	119W08.69	375	Warden
YA2	46N31.60	120W31.80	652	Yakima Two
YPT	46N02.93	118W57.73	325	Yellepit
* Three-comp	onent station.	•	•	-

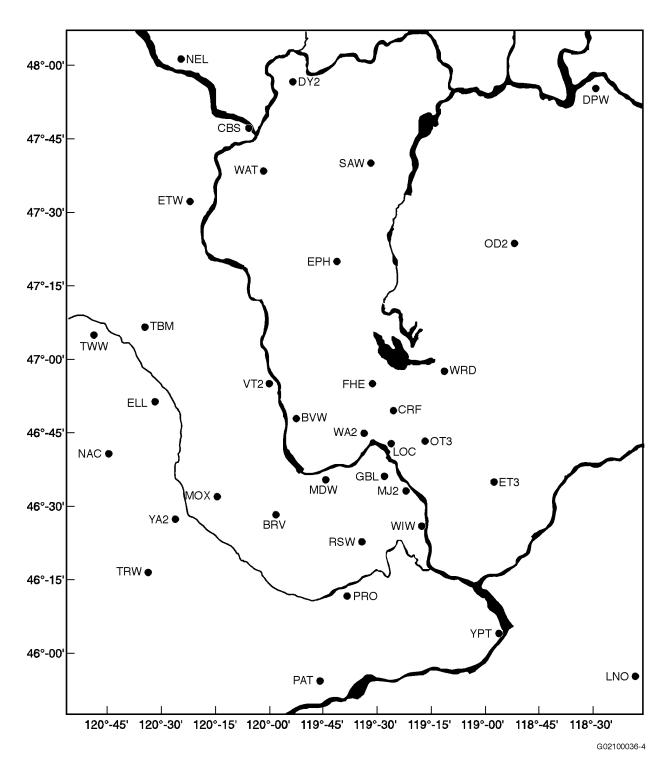


Figure 2.2. Locations of Seismograph Stations in the Eastern Washington Regional Network (see Table 2.2 for location descriptions).

 Table 2.3.
 Acquisition System Recorded Triggers

Event Type	First Quarter	Second Quarter	Third Quarter	Fourth Quarter	Total	Description
South-Central Washington	18	17	23	25	83	Seismic events in south-central Washington and north-central Oregon that triggered the HSN.
Regional	55	40	72	70	237	Seismic events in the Western United States and Canada.
Teleseism	64	56	59	54	233	Seismic events at farther distances from around the world.
Total Earthquake Events	137	113	154	149	553	Total number of earthquake triggers.
Total Triggers on Primary System	383	242	330	222	1,177	Total number of triggers examined. Includes all sources of triggers.
Local Explosions	3	6	4	3	16	Quarry blasts, typically, within the 46-47 degrees north latitude and 119-120 degrees west longitude.
Local Earthquakes	10	5	14	13	42	Seismic events within the 46-47 degrees north latitude and 119-120 degrees west longitude.

 Table 2.4.
 Free-Field Strong Motion Accelerometer Sites

Site	Site ID	Location	Latitude Longitude Elevation
100-K Area	H1K	South of K Basins outside 100 Area fence lines.	46° 38.51' 119° 35.53' 152 m
200 East Area	Н2Е	East of B Plant; north of 7 th Street and east of Baltimore Avenue.	46° 33.58' 119° 32.00' 210 m
200 West Area	H2W	Northeast of Plutonium Finishing Plant (PFP); north of 19 th street and east of Camden Avenue.	46° 33.23' 119° 37.51' 206 m
300 Area	Н3А	South end of 300 Area inside fence lines. (NE 1/4, SW 1/4, Sec. 11, T10N, R28E).	46° 21.83' 119° 16.55' 119 m
400 Area	H4A	500 feet from fence line on east side of facility and north of parking area).	46° 26.13' 119° 21.30' 171 m

The instrumentation locations were chosen based on two criteria (Moore and Reidel 1996): 1) instruments should be located in areas having the highest densities of people and 2) instruments should be located in areas having hazardous facilities. Some of the highest concentrations of employees at Hanford are 200 East and West Areas, 100-K Area, the Fast Flux Test Facility (400 Area), and the 300 Areas. The 200 Areas are where high-level radioactive waste from past processing of fuel rods is stored in single-shell and double-shell tanks. In addition, the Canister Storage Facility that holds encapsulated spent fuel rods is in 200 East Area. The 100-K Area contains the K Basins where spent fuel rods from the N Reactor are stored prior to encapsulation. The Cold Vacuum Drying Facility, located in the 100-K Area, is used to encapsulate spent fuel rods from the K Basins prior to shipment to the Canister Storage Building in 200 East Area.

2.2.2 Site Design

All free-field SMA sites consist of a four panel solar array and two 30-gallon galvanized drums. Each panel has a maximum 42 watt output. The two 30-gallon drums are set in the ground such that the base of the drum is about 1 m below the surface. One drum houses only the SMA; the other drum, which is connected via a sealed conduit to the SMA drum, previously housed the batteries and a cellular telephone. During FY 2002 a change was made to the configuration. The second drum now contains only the batteries. The cellular telephone was replaced with a Cellular Digital Packet Data (CDPD) system which provides a continuous radio data-link with the AT&T internet service. This CDPD system along with the solar regulator is housed in a small enclosure mounted at the rear of the solar array. The enclosure serves as an junction box for all cabling between in the drums and outside the drums through conduit. The antenna for the CDPD is mounted on top of the enclosure. The enclosure permits quick access to check battery conditions and a connection directly to the RS-232 port of the SMA without removing the drum lids.

The SMA instruments are three-component units consisting of one vertical, one north-south horizontal, and one east-west horizontal data channel. The instruments in use are the ETNATM system (registered trademark of Kinemetrics, Inc., Pasadena, California). Instrument specifications are summarized in Table 2.5. In addition to the three-component SMAs, each ETNA SMA unit contains a computer, Global Positioning System (GPS) receiver (Figure 2.3). These systems are housed in a watertight box.

The CDPD system provides the internet address connection to access the system. Stations can be monitored from any computer, and data can be downloaded to a dedicated computer in the seismic monitoring laboratory. The data can also be downloaded directly at each site via a built in cable connection at the enclosure in case of communication failure.

The SMAs have an internal GPS receiver used principally to link it to the National Bureau of Standards timing system. (a) The GPS is internally activated approximately every 4 hours and checks the "location of the instrument" and the time. Any differences between the internal clock and the GPS time

⁽a) The GPS antenna is mounted on the enclosure at the rear of the solar array.

Table 2.5. Instrument Parameters for the Kinemetrics ETNATM System in the Hanford SMA Network

Parameter	Value or Range						
Sensor							
Туре	Tri-axial Force Balance Accelerometer orthogonally oriented with internal standard						
Full-Scale	±2 g ^(a)						
Frequency Range	0-50 Hz						
Damping	Approximately 70% critical (a)						
Data Acquisition							
Number of Channels	3						
Sample Rate	18-bit resolution @ 200 samples/second						
Digital Output	Real-time, RS-232 Output Stream						
Seismic Trigger							
Filter	0.1 - 12.5 Hz						
Trigger level	0.05% - 0.20% g ^(b)						
Alarm (call-out) Threshold	4.00% g						
Pre-event Memory	10 sec						
Post-event Time	40 sec						
(a) Setting is dependent on instrument calibration.(b) See Section 2.2.4 for discussion of trigger thresholds.							

Office of Seismic Monitoring Rm 2411, Sigma V Building Richland, Washington AT&T Personal **CDPD** Computer **Strong Motion Accelerometer Remote Site** Cellular Global Digital Strong Motion Computer Positioning Accelerometer Packet Data System System Battery Voltage Solar Pack Regulator **Panels** G02100036-5

Figure 2.3. Schematic Diagram of a Strong Motion Accelerometer Installation

are recorded and saved by the SMA. Any corrections to the internal timing are made automatically. Typically, the greatest correction recorded is approximately 4 milliseconds.

2.2.3 Strong Motion Accelerometer Operations Center

The combined operations, data recording, data interpretation, and maintenance facility is located in the Sigma V building and is operated by the PNNL Seismic Monitoring Team.

2.2.4 Strong Motion Operational Characteristics

The signals from the three-accelerometer channels at each site are digitized with a 24-bit digitizer and temporarily stored in a memory buffer. The sampling rate of the digitizer is set to 200 samples/second. The three channels are monitored for signals that equal or exceed a programmable trigger threshold. When one accelerometer channel is triggered, the other channels automatically record. The nominal threshold used is 0.05% of the full-scale range of 2.0 g (g is the acceleration of gravity, 9.8 m/s² or 32 ft/s²) or 0.001 g. Threshold trigger levels are being adjusted to trigger infrequently on the noise sources (e.g., vehicles, sonic booms) near each site. This will provide ground motion data for smaller, non-damaging earthquakes that can be useful in estimating the ground motion expected from larger earthquakes, and to confirm correct operation of the instruments by analyzing the smaller-amplitude triggers. The recorders store information for 10 seconds before the trigger threshold is exceeded and for 40 seconds after the trigger ceases to be exceeded.

3.0 Earthquake Catalog Description

3.1 Coda Length Magnitude

Coda-length magnitude (M_c), an estimate of local magnitude (M_L) (Richter 1958), is calculated using the coda-length/magnitude relationship determined for Washington State by Crosson (1972).

3.2 Velocity Model

The program XPED uses the velocities and layer depths given in Table 3.1. XPED was developed at UW and the velocity model used in XPED is based on Rohay et al. (1985). XPED is an interactive earthquake seismogram display program used to analyze seismic events.

3.3 Quality Factors (Q)

XPED assigns a two-letter **Quality factor** (Table 3.2) that indicates the general reliability of the solution (**A** is best quality, **D** is worst). Similar quality factors are used by the USGS for events located with the computer program HYPO71. The first letter of the quality code is a measure of the hypocenter quality based primarily on travel time residuals. For example: Quality **A** requires a root-mean-square residual (**RMS**) less than 0.15 seconds while a **RMS** of 0.5 seconds or more is **D** quality (other estimates of the location uncertainty also affect this quality parameter). The second letter of the quality code is related to the spatial distribution of stations that contribute to the event's location, including the number of stations (**NS**), the number of p-wave and s-wave phases (**NP**), the largest gap in event-station azimuth distribution (**GAP**), and the closest distance from the epicenter to a station (**DMIN**). Quality **A** requires a solution with **NP** >8, **GAP** <90°, and **DMIN** <5 km (or the hypocenter depth if it is greater than 5 km). If **NP** ≤5, **GAP** >180°, or **DMIN** >50 km, the solution is assigned Quality **D**.

Table 3.1. Seismic Velocities for Columbia Basin Stratigraphy (from Rohay et al. 1985)

Depth to Top of Velocity Layer (km)	Stratigraphy	Velocity (km/sec)
0.0	Saddle Mountains and Wanapum Basalts and intercalated Ellensburg Formation	3.7
0.4	Grande Ronde Basalt and pre-basalt sediments	5.2
8.5	Crystalline Basement, Layer 1	6.1
13.0	Crystalline Basement, Layer 2	6.4
23.0	Sub-basement Sub-basement	7.1
38.0	Mantle	7.9

Table 3.2. Local Earthquake Data, October 1, 2001 to September 30, 2002

Event ID	Type	Date	Time	Latitude	Longitude	Depth	Mag	NS/NP	Gap	DMIN	RMS	Q	Location
01101700254		01/10/17	00:26:02.61	46N21.97	119W37.05	4.50	0.3	7/08	138	3	0.29	BC	21.2 km NNE of Prosser
01102206001		01/10/22	06:00:34.90	46N42.55	119W28.75	1.82	0.7	8/11	129	3	0.04	AB	11.7 km NE of 100-K Area
01102206004		01/10/22	06:01:18.21	46N42.68	119W27.98	7.64	-0.3	4/08	134	2	0.31	CD	12.6 km NE of 100-K Area
01110101030		01/11/01	01:03:24.23	46N16.86	119W23.05	0.02	1.0	7/08	279	4	0.52	DD	7.1 km W of Richland
01110113325		01/11/01	13:33:19.34	46N42.25	119W29.01	2.21	0.8	6/08	100	4	0.07	AC	11.1 km NE of 100-K Area
01110617153		01/11/06	17:16:00.52	46N42.30	119W28.95	2.14	0.9	14/20	78	4	0.07	AA	11.2 km NE of 100-K Area
01110820001	P	01/11/08	20:00:41.49	46N18.45	119W40.15	0.53		12/16	121	10	0.06	AC	13.5 km NE of Prosser
01120804035		01/12/08	04:04:04.67	46N09.53	119W04.83	9.23	0.7	7/10	104	15	0.27	BB	5.9 km SE of Kennewick
01121222090	P	01/12/12	22:09:22.58	46N10.72	119W44.97	0.03		6/07	309	6	0.23	BD	3.5 km SSE of Prosser
01121504353		01/12/15	04:35:54.75	46N19.70	119W50.73	17.96	0.7	17/23	145	20	0.18	BC	14.8 km NNW of Prosser
01121609500		01/12/16	09:50:28.11	46N32.67	119W35.03	17.76	0.9	24/29	41	7	0.15	AA	4.1 km WSW of 200 East
01121719541	P	01/12/17	19:54:21.92	46N56.27	119W05.58	5.37		5/07	247	5	0.32	CD	13.6 km NNE of Othello
01122913593		01/12/29	14:00:01.07	46N41.50	119W42.34	14.57	0.5	13/18	98	9	0.07	AB	10.0 km WNW of 100-K Area
02011722132	P	02/01/17	22:13:44.85	46N55.85	119W54.73	2.42		17/17	85	7	0.14	AB	6.2 km ESE of Vantage
02011918522	P	02/01/19	18:52:48.01	46N08.21	119W27.97	0.98		11/11	273	17	0.08	BD	21.1 km SW of Richland
02020517492	P	02/02/05	17:49:47.13	46N07.48	119W02.31	2.45		9/10	136	10	0.17	BC	10.9 km SE of Kennewick
02020521422	X	02/02/05	21:42:45.49	46N14.42	119W43.45	0.02		16/17	64	4	0.11	AA	5.0 km NE of Prosser
02020601005	X	02/02/06	01:01:22.32	46N07.61	119W28.22	0.66		17/19	111	19	0.15	BC	22.2 km SW of Richland
02021422303		02/02/14	22:30:53.46	46N13.79	119W45.56	7.41	1.3	21/34	211	6	0.14	AD	2.6 km NNE of Prosser
02021422595		02/02/14	23:00:15.65	46N14.01	119W45.03	6.01	0.5	9/11	211	5	0.13	AD	3.2 km NNE of Prosser
02021613123		02/02/16	13:12:50.85	46N50.47	119W44.78	1.68	0.7	11/18	140	10	0.20	BC	22.3 km SE of Vantage
02021910095		02/02/19	10:10:14.92	46N13.05	119W52.06	14.23	0.9	9/14	146	14	0.22	BC	7.8 km W of Prosser
02031102482		02/03/11	02:48:45.91	46N50.73	119W44.52	3.51	2.5	25/27	54	11	0.14	AC	22.3 km SE of Vantage
02031200293	P	02/03/12	00:29:55.53	46N10.47	119W13.94	0.04		9/09	283	20	0.08	AD	9.0 km WSW of Kennewick

Table 3.2. (contd)

Event ID	Type	Date	Time	Latitude	Longitude	Depth	Mag	NS/NP	Gap	DMIN	RMS	Q	Location
02040620452		02/04/06	20:45:45.10	46N14.04	119W46.35	6.02	0.3	9/17	251	7	0.10	AD	3.0 km N of Prosser
02041019252		02/04/10	19:25:50.90	46N23.43	119W39.01	13.41	0.1	9/16	218	4	0.06	AD	18.7 km S of 200 West
02041410582		02/04/14	10:58:46.48	46N35.17	119W20.68	15.79	0.3	13/19	74	3	0.14	AA	14.9 km ENE of 200 East
02041917210		02/04/19	17:21:27.58	46N05.51	119W38.31	11.56	0.5	5/06	166	13	0.06	AD	16.3 km SE of Prosser
02042620265	P	02/04/26	20:27:15.66	46N04.81	119W34.37	5.21		5/06	188	17	0.39	CD	20.7 km SE of Prosser
02050615093		02/05/06	15:09:57.61	46N36.19	119W47.82	7.55	0.1	7/10	259	2	0.05	AD	13.2 km WNW of 200 West
02051321165	P	02/05/13	21:17:19.01	46N04.10	119W33.87	7.32		4/07	194	18	0.13	AD	22.0 km SE of Prosser
02051720554	P	02/05/17	20:55:57.29	46N04.80	119W36.16	0.02		7/07	179	16	0.08	AC	19.0 km SE of Prosser
02052407023		02/05/24	07:02:55.60	46N12.98	119W17.94	17.60	0.9	17/20	130	13	0.14	AB	7.5 km S of Richland
02052418534	P	02/05/24	18:54:04.74	46N52.71	119W59.60	0.61		18/18	62	9	0.14	AB	8.9 km S of Vantage
02060106040		02/06/01	06:04:33.12	46N16.57	119W32.27	7.98	0.5	12/16	165	8	0.06	AC	19.0 km W of Richland
02060106373		02/06/01	06:37:55.36	46N16.56	119W32.23	8.21	0.3	11/12	260	8	0.07	AD	18.9 km W of Richland
02060107530		02/06/01	07:52:30.74	46N24.74	119W16.03	0.02	0.4	6/11	249	2	0.20	BD	6.1 km N of 300 Area
02060213162		02/06/02	13:16:46.77	46N18.02	119W32.72	15.55	0.6	17/18	135	8	0.14	AB	19.6 km W of Richland
02060221442		02/06/02	21:44:50.81	46N30.14	119W15.90	0.50	0.4	10/12	120	8	0.06	AB	10.4 km NE of 400 Area
02061609254		02/06/16	09:26:06.89	46N26.63	119W37.45	16.06	0.1	8/11	154	3	0.06	AC	12.7 km S of 200 West
02062300422		02/06/23	00:42:50.68	46N29.32	119W13.19	14.75	1.7	19/27	92	8	0.11	AB	12.2 km ENE of 400 Area
02062504113		02/06/25	04:12:01.73	46N36.74	119W46.32	15.47	0.1	7/12	181	0	0.07	AD	11.9 km WNW of 200 West
02072600292		02/07/26	00:29:55.85	46N34.81	119W41.13	0.13	-0.5	5/09	111	5	0.14	AD	4.4 km NW of 200 West
02080717000		02/08/07	17:00:23.95	46N42.83	119W28.09	0.70	-0.1	4/04	141	2	0.00	AD	12.7 km NE of 100-K Area
02080909180		02/08/09	09:18:24.63	46N07.57	119W48.71	0.02	0.0	7/07	199	13	0.09	AD	9.6 km SSW of Prosser
02081106062		02/08/11	06:06:51.59	46N13.69	119W37.00	7.36	0.3	9/10	140	5	0.08	AC	11.9 km ENE of Prosser
02081121492		02/08/11	21:49:50.19	46N11.95	119W36.44	9.99	0.4	10/12	233	6	0.07	AD	12.4 km E of Prosser
02081309064		02/08/13	09:07:06.67	46N13.40	119W37.22	7.77	0.4	7/09	208	5	0.07	AD	11.5 km E of Prosser
02082201111		02/08/22	01:11:32.07	46N51.51	119W32.54	14.33	-0.3	7/09	185	10	0.03	AD	24.4 km N of 100-K Area

Table 3.2. (cont)

Event ID	Type	Date	Time	Latitude	Longitude	Depth	Mag	NS/NP	Gap	DMIN	RMS	Q	Location
02082201214		02/08/22	01:21:58.19	46N51.49	119W32.38	15.05	0.0	7/11	165	10	0.06	AC	24.4 km N of 100-K Area
02082800424	P	02/08/28	00:43:12.06	46N15.35	119W28.20	0.41		6/07	225	5	0.02	AD	14.1 km WSW of Richland
02090418423		02/09/04	18:43:00.33	46N26.20	119W36.28	16.47	0.0	9/14	137	4	0.05	AC	13.7 km S of 200 West
02090818241		02/09/08	18:24:32.22	46N33.55	119W32.68	14.95	-0.3	7/08	120	7	0.05	AB	0.7 km W of 200 East
02091107473		02/09/11	07:47:53.25	46N01.44	119W42.23	8.54	0.6	9/09	165	16	0.05	AC	21.0 km SSE of Prosser
02092021232	P	02/09/20	21:23:50.38	46N14.35	119W26.88	5.19		7/07	253	6	0.22	BD	13.0 km WSW of Richland
02092415094		02/09/24	15:10:01.43	46N07.89	119W26.43	7.34	0.1	4/05	335	18	0.06	AD	20.4 km SW of Richland
02092509220		02/09/25	09:22:27.19	46N36.65	119W18.92	13.96	0.2	16/27	76	6	0.10	AA	17.8 km ENE of 200 East
02092823595	P	02/09/29	00:00:14.56	46N14.63	119W26.92	0.02		7/07	250	6	0.11	AD	12.9 km WSW of Richland

Explanation of Table 3.2

EVENT ID: The Earthworm Recording System creates the identification number. XPED uses the

year, month, day and time to create a unique number for each event.

TYPE: P is Probable Blast; X is Confirmed Blast; F is Felt Earthquake; S is surficial event

(rockslide, avalanche) and not a explosion or tectonic earthquake; blank is local

earthquake.

DATE: The year and day of the year in Universal Time Coordinated (UTC). UTC is used

throughout this report unless otherwise indicated.

TIME: The origin time of the earthquake given in UTC. To covert UTC to Pacific Standard

Time, subtract eight hours; to Pacific Daylight Time, subtract seven hours.

LATITUDE: North latitude, in degrees and minutes, of the earthquake epicenter.

LONGITUDE: West longitude, in degrees and minutes, of the earthquake epicenter.

DEPTH: The depth of the earthquake in kilometers (km).

MAG: The magnitude is expressed as Coda-Length magnitude M_c, an estimate of local mag-

nitude M_L (Richter 1958). If magnitude is blank a determination was not made.

NS/NP: Number of stations/number of phases used in the solutions.

GAP: Azimuthal gap. The largest angle (relative to the epicenter) containing no stations.

DMIN: The distance from the earthquake epicenter to the closest station

RMS: The root-mean-square residual (observed arrival times minus the predicted arrival

times) at all stations used to locate the earthquake. It is only useful as a measure of quality of the solution when five or more well-distributed stations are used in the solution. Good solutions are normally characterized by RMS values of less than about

0.3 seconds.

Q: The Quality Factors indicate the general reliability of the solution/location (A is best

quality, D is worst). See Section 3.3 of this report: Quality Factors.

4.0 Geology and Tectonic Analysis

The Hanford Site lies within the Columbia Basin, which is an intermontane basin between the Cascade Range and the Rocky Mountains that is filled with Cenozoic volcanic rocks and sediments. This basin forms the northern part of the Columbia Plateau physiographic province (Fenneman 1931) and the Columbia River flood-basalt province (Reidel and Hooper 1989). In the central and western parts of the Columbia Basin, the Columbia River Basalt Group (CRBG) overlies Tertiary continental sedimentary rocks and is overlain by late Tertiary and Quaternary fluvial and glaciofluvial deposits (Campbell 1989; Reidel et al. 1989, 1994; DOE 1988). In the eastern part, a thin (<100 m) sedimentary unit separates the basalt and underling crystalline basement and a thin (<10 m) veneer of eolian sediments overlies the basalt (Reidel et al. 1989, 1994).

The Columbia Basin has two structural subdivisions or subprovinces: the Yakima Fold Belt and the Palouse Slope. The Yakima Fold Belt includes the western and central parts of the Columbia Basin and is a series of anticlinal ridges and synclinal valleys with major thrust faults along the northern flanks (Figure 4.1) (Reidel and Fecht 1994a, 1994b). The Palouse Slope is the eastern part of the basin and is less deformed than the Yakima Fold Belt with only a few faults and low amplitude, long wavelength folds on an otherwise gently westward dipping paleoslope. Figure 4.2 shows north-south and east-west cross sections through the Columbia Basin based on surface mapping (Reidel and Fecht 1994a, 1994b), deep boreholes (Reidel et al. 1994), geophysical data (Rohay et al. 1985; DOE 1988), and magnetotelluric data obtained as part of BWIP (DOE 1988).

4.1 Earthquake Stratigraphy

Studies of seismicity at the Hanford Site have shown that the seismic activity is related to crustal stratigraphy (layers of rock types) (Rohay et al. 1985; DOE 1988). The main geologic units important to earthquakes at Hanford and the surrounding area are:

- The Miocene Columbia River Basalt Group (CRBG)
- Pre-basalt sediments of Paleocene, Eocene, and Oligocene age
- The crystalline basement consisting of 2 layers composed of Precambrian and Paleozoic craton and Mesozoic accreted terranes.

4.2 Geologic Structure Beneath the Monitored Area

Between the late 1950s and the early 1980s, deep boreholes were drilled for hydrocarbon exploration in the Columbia Basin. These boreholes provided accurate measurements of the physical properties of the CRBG and the pre-basalt sediments (Reidel et al. 1989, 1994), but the thickness of the pre-basalt sediments and nature of the crystalline basement are still poorly understood. The difference between the thicknesses listed in Table 4.1 and the thicknesses of the crustal layers in the velocity model in Table 3.1 reflect data specific to UW's crustal velocity model for eastern Washington. Table 4.1 is derived from

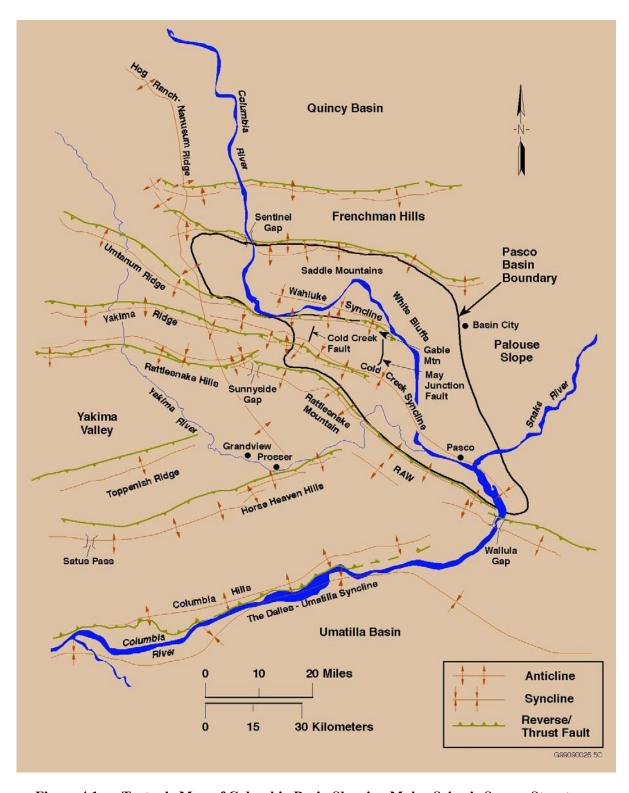


Figure 4.1. Tectonic Map of Columbia Basin Showing Major Seismic Source Structures

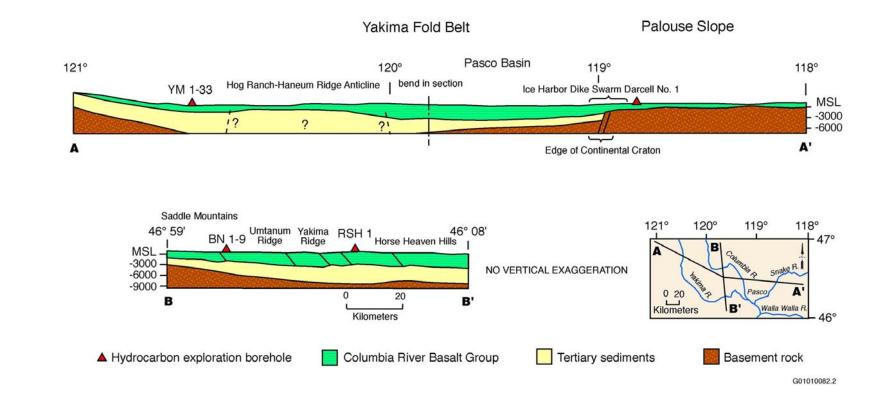


Figure 4.2. Geologic Cross Sections Through the Columbia Basin (Reidel et al. 1994)

Table 4.1. Thicknesses of Stratigraphic Units in the Monitoring Area

Stratigraphy	North	South	East	West
Columbia River Basalt Group (includes suprabasalt sediments)	3.0 km	4.5 km	2.2 km	4.2 km
Pre-basalt Sediments	3.0 km	>4.5 km	0	>6.0 km

Figure 4.2 and was developed for the geologic interpretation in this report. The thicknesses of these units are variable across the monitored area. Table 4.1 summarizes the approximate thickness at the borders of the monitored area.

The thickness of the basalt and the pre-basalt sediments varies as a result of different tectonic environments. The western edge of the North American craton (late Precambrian/Paleozoic continental margin and Precambrian craton) is located in the eastern portion of the monitored area. The stratigraphy on the craton consists of CRBG overlying crystalline basement; the crystalline basement is continental crustal rock that underlies much of the western North America. The stratigraphy west of the craton consists of 4 to 5 km of CRBG overlying greater than 6 km of pre-basalt sediments. This in turn overlies accreted terranes of Mesozoic age. The area west of the craton was subsiding during the Eocene and Oligocene, accumulating great thickness of pre-CRBG sediments. Continued subsidence in this area during the Miocene resulted in thicker CRBG compared to that on the craton. Subsidence continues today but at a greatly reduced rate (Reidel et al. 1994).

4.3 Depth of Earthquakes

Since records have been kept, most of the earthquakes at the Hanford Site have originated in the CRBG layer. The crystalline basement has had the next greatest amount of earthquakes followed by the pre-basalt sediments. The stratigraphic units for local earthquakes recorded during FY 2002 are listed in Table 4.2.

Table 4.2. Number of Local Earthquakes Occurring in Stratigraphic Units

Unit	First Quarter	Second Quarter	Third Quarter	Fourth Quarter	FY 2002
Basalt	6	2	2	3	13 (30.9%)
Pre-basalt Sediments	1	2	4	5	12 (28.6%)
Crystalline Basement	3	1	8	5	17 (40.5%)
Total	10	5	18	16	42

4.4 Tectonic Pattern

Studies have concluded that earthquakes can occur in the following six different tectonic environments (earthquake sources) at the Hanford Site (Geomatrix 1996).

- Reverse/thrust faults. Reverse/thrust faults in the CRBG associated with major anticlinal ridges such as Rattlesnake Mountain, Yakima Ridge, and Umtanum Ridge could produce some of the largest earthquakes.
- Secondary faults. These are associated with the major anticlinal ridges.
- Swarm areas. Small geographic areas not known to contain any geologic structures produce clusters of events (swarms), usually in the CRBG in synclinal valleys. These clusters consist of a series of small shocks with no outstanding principal event. Swarms occur over a period of days or months and the events may number into the hundreds and then quit, only to start again at a later date. This differs from the sequence of foreshocks, mainshock, and trailing-off aftershocks that have the same epicenter or are associated with the same fault system. In the past, swarms were thought to occur only in the Columbia River Basalt Group. Most swarm areas are in the basalt but swarm events also appear to occur in all geologic layers. However, typically a swarm event at a specific time is usually restricted to one layer. There are seven earthquake swarm areas that we recognize in the monitoring area (Figure 4.3) but this list will be updated as new swarm areas develop. The Saddle Mountains Swarm Area, Wooded Island Swarm Area and Coyote Rapids Swarm area are typically active at one time or another during each year. The other earthquake swarm areas are active less frequently.
- The entire Columbia Basin. The entire basin, including the Hanford Site, could produce a "floating" earthquake. A floating earthquake is one that, for seismic design purposes, can happen anywhere in a tectonic province and is not associated with any known geologic structure. Seismic Monitoring classifies it as a random event for purposes of seismic design and vibratory ground motion studies.
- **Basement source structures**. Studies (Geomatrix 1996) suggest that major earthquakes can originate in tectonic structures in the crystalline basement. Because little is known about geologic structures in the crystalline basement beneath the Hanford Site, earthquakes cannot be directly tied to a mapped fault. Earthquakes occurring in the crystalline basement without known sources are treated as random events.
- The Cascadia Subduction Zone. This source has been postulated to be capable of producing a magnitude 9 earthquake. Because this source is along the western boundary of Washington State and outside the HSN, the Cascadia Subduction Zone is not an earthquake source that is monitored at the Hanford Site, so subduction zone earthquakes are not reported here. Because any earthquake along the Cascadia Subduction zone can have a significant impact on the Hanford Site or can be felt like the February 2001 Nisqually earthquake, UW monitors and reports on this earthquake source for DOE. Ground motion from any moderate or larger Cascadia Subduction Zone earthquake is detected by Hanford SMAs and reported (see Section 5.0).

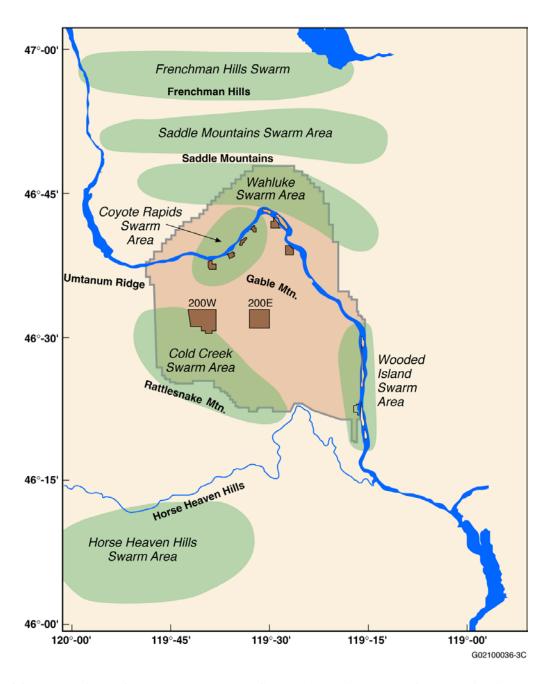


Figure 4.3. Locations of Known Earthquake Swarm Areas in the Hanford Monitoring Network Area

4.5 Tectonic Activity

4.5.1 Annual Summary

Forty-two earthquakes occurred in the Hanford monitoring area during FY 2002 (October 1, 2001 through September 2002) (Table 4.3) (Figure 4.4). This section summarizes the earthquake activity for the fiscal year. More detailed descriptions of this activity for each quarter are given in later sections.

Table 4.3. Summary of Earthquake Locations for FY 2002

Seis	emic Sources	First Quarter 10/01- 12/30	Second Quarter 1/01 - 3/31	Third Quarter 4/01 - 6/30	Fourth Quarter 7/01 - 9/30	FY 2002
Geologic Structure		1	0	0	0	1 (2.3%)
	Saddle Mountains/ Royal Slope	0	2	0	0	2 (4.8%)
	Coyote Rapids	0	0	0	0	0
Swarm	Wooded Island	0	0	2	0	2 (4.8%)
Areas	Wahluke Slope	4	0	0	1	5 (11.9 %)
	Cold Creek	0	0	0	0	0
	Horse Heaven Hills	0	0	0	4	4 (9.5%)
	Total for swarms	4	2	2	5	13 (31%)
Random Events		5	3	12	8	28 (66.7%)
Total for all earthquakes		10	5	14	13	42

4.5.1.1 Depth of Earthquakes

During FY 2002, 31% of the earthquakes occurred in the Columbia River Basalt Group, 29% of the earthquakes occurred in the underlying pre-basalt sediments, and 40% of the earthquakes occurred in the crystalline basement. Typically, more earthquakes occur in the basalt but earthquakes in the crystalline basement dominated this year.

4.5.1.2 Location of Earthquakes

During FY 2002, thirteen events (31% of the earthquakes) were classified as swarm events. The Saddle Mountains, Wooded Island, Horse Heaven Hills and Wahluke swarm areas were active at one time or another during FY 2002. The Wahluke and Horse Heaven Hills were the most active having five and four events each, respectively.

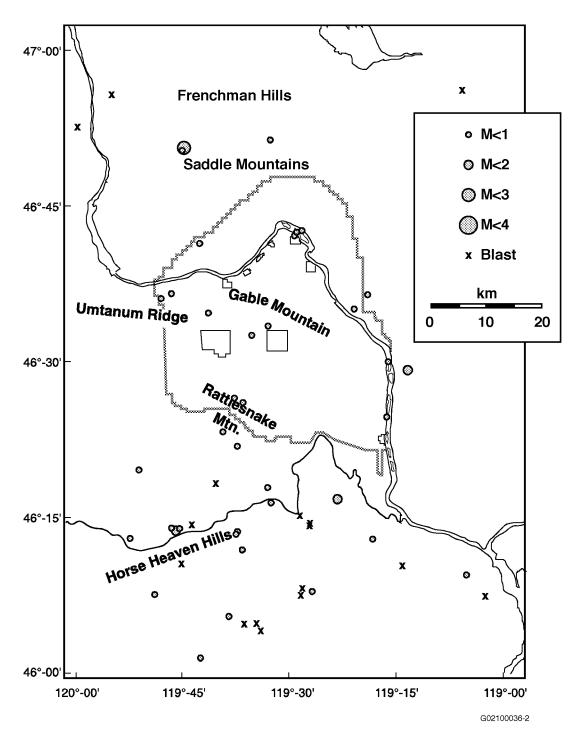


Figure 4.4. All Earthquakes Recorded in the Hanford Monitoring Area Between October 1, 2001 and September 30, 2002 (Coda Length Magnitude (M_c) scale is shown at the side of the map)

Only one event during FY 2002 was classified as having some association with major geologic structure. An earthquake occurred along the Rattlesnake-Wallula structural trend (RAW Figure 4.1) between Red Mountain and Candy Mountain. These are the first and second doubly plunging anticlines, respectively, southeast of Rattlesnake Mountain. This event occurred in the basalt.

Twenty-eight events (66.7% of the earthquakes) were classified as random events. Earthquakes typically are classified as random if they occur below the Columbia River Basalt Group. Very little is known about geologic structures in the pre-basalt sediments and crystalline basement so any interpretations are speculative at this time. An earthquake also can be classified as a random event if it occurs in the basalt but is not located near any known geologic structure. During FY 2002, seventeen random events occurred in the crystalline basement. Three random events occurred in the basalt and were not near any known structure. Eight random events occurred in the pre-basalt sediments.

4.5.2 First Quarter of FY 2002

The locations of all earthquakes that occurred between October 1, 2001 and December 31, 2001 are shown on Figure 4.5.

4.5.2.1 Major Anticlinal Ridges

During the first quarter of FY 2002, we interpret one seismic event to have occurred on a major geologic structure. On October 10th, a small (1.0 M_c), shallow earthquake occurred along the Rattlesnake-Wallula structural trend (RAW Figure 4.1) between Red Mountain and Candy Mountain. Red Mountain and Candy Mountain are the first and second doubly plunging anticlines, respectively, southeast of the main Rattlesnake Mountain structure. This event occurred in the CRBG.

4.5.2.2 Earthquake Swarm Areas

During the first quarter of FY 2002, we interpret 4 seismic events to have occurred in swarm areas (Figure 4.5). All four events occurred in the Wahluke swarm area. The other swarm areas were not active.

4.5.2.2.1 Wahluke Slope Swarm Area

During the first quarter of FY 2002, four events occurred in the Wahluke swarm area (Figure 4.5). All were small ($M_c < 1.0$) and occurred at the southwest end of the swarm that was active during most of FY 2001. Two of the events occurred on October 22^{nd} followed by one each on November 1^{st} and 6^{th} . All events probably occurred in the basalt although the smallest (approximately $0.0 M_c$) gave a depth in the pre-basalt sediments. Because of its small size, this depth solution could be in error.

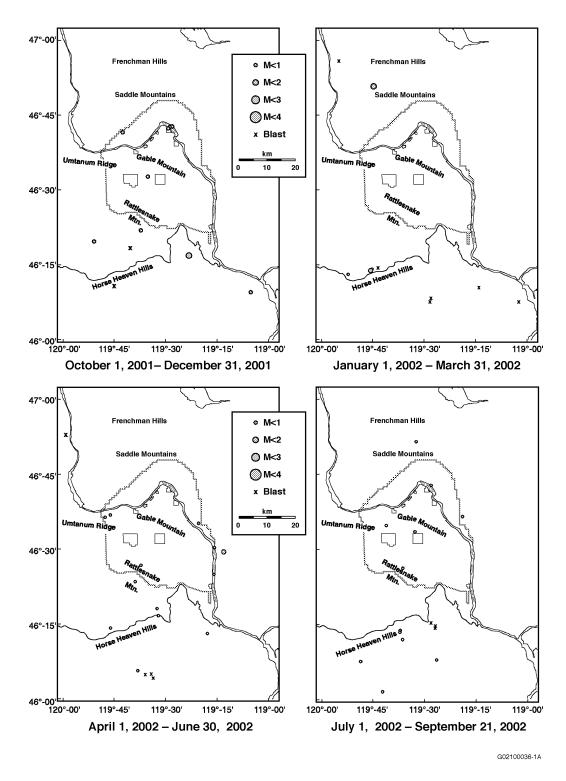


Figure 4.5. Locations of All Events Between October 1, 2001 and September 30, 2002 Separated by FY Quarter (Coda Length Magnitude (M_c) scale is shown at the side)

4.5.2.3 Random or Floating Events

During the first quarter of FY 2002, we interpret five small ($M_c < 1.0$), random events to have occurred. One was in the basalt, one was in pre-basalt sediments and three were in the crystalline basement.

Two events occurred under the south slope of Rattlesnake Mountain (Figure 4.1). The first one occurred on October 17th in the basalt and was about 3 km from the crest of the ridge and not near any known geologic structure. The second event occurred on December 15th about 10 km south of the Rattlesnake Hills. This event occurred in the crystalline basement.

The three other events were scattered across the monitoring area. On December 8th an event occurred in the pre-basalt sediments 1 km north of The Butte, which is one of the doubly plunging anticlines along the Rattlesnake-Wallula structural trend (RAW Figure 4.1). On December 16th an event occurred in the crystalline basement under the 200 West Area. The final event of the quarter occurred on December 29th in the crystalline basement under the Wahluke Slope.

4.5.3 Second Quarter of FY 2002

The locations of all mapped earthquakes that occurred between January 1, 2002 and March 31, 2002 are shown on Figure 4.5.

4.5.3.1 Major Anticlinal Ridges

There were no events that we interpreted to have occurred on a major geologic structure during this quarter.

4.5.3.2 Earthquake Swarm Areas

During the second quarter of FY 2002, we interpret 2 seismic events to have occurred in swarm areas (Figure 4.5). Both events occurred in the Saddle Mountains swarm area (Figure 4.3). The other swarm areas were not active.

4.5.3.2.1 Saddle Mountain Swarm Area

Two events occurred at the same location in the basalt in the Saddle Mountain swarm area. The first event $(0.7~M_c)$ occurred on February 16^{th} and second $(2.5~M_c)$ occurred on March 11^{th} . Both events were in the basalt.

4.5.3.3 Random or Floating Events

During the second quarter, three events were classified as random events. Two events occurred in the pre-basalt sediments and one occurred in the crystalline basement. On February 14th, two small events (1.3 and 0.5 M_c) occurred in the pre-basalt sediments on the extreme south slope of Rattlesnake Mountain

2 km north of Prosser and the Horse Heaven Hills. The third event occurred on February 19th 3 km north of the Horse Heaven Hills. This event occurred in the crystalline basement and was 6 km west of the two previous random events of this quarter.

4.5.4 Third Quarter of FY 2002

The locations of all mapped earthquakes that occurred between April 1, 2002 and June 30, 2002 are shown on Figure 4.5.

4.5.4.1 Major Anticlinal Ridges

During the third quarter of FY 2002, we interpret no seismic event to have occurred on major geologic structures.

4.5.4.2 Earthquake Swarm Areas

Two earthquakes occurred in swarm areas during the third quarter of FY 2002. Both occurred in the Wooded Island swarm area (Figure 4.1).

4.5.4.2.1 Wooded Island Swarm Area.

One earthquake occurred on June 1st within the Wooded Island swarm area south of Johnson Island. This event was small (0.4 M_c) and was in the CRBG. A second event occurred at the north end of the swarm area on June 2nd. It also was small (0.4 M_c) and occurred in the basalt.

4.5.4.3 Random or Floating Events

During the third quarter, twelve events were classified as random. Four occurred in the pre-basalt sediments and eight occurred in the crystalline basement. The events were scattered across the monitored area.

On April 6^{th} , a small $(0.3~M_c)$ event occurred was just north of the Horse Heaven Hills. This event occurred at the same location as two previous events of February 14, 2002 and all three were in the prebasalt sediments. The next random seismic event occurred on April 10^{th} and was in the crystalline basement. It was small $(0.1~M_c)$ and occurred on the south slope of Rattlesnake Mountain. On April 14^{th} , a small event $(0.3~M_c)$ occurred in the crystalline basement beneath the Horse Heaven Hills southeast of Prosser, Washington. The last event for the month of April was small $(0.5~M_c)$ and occurred on April 19^{th} in the crystalline basement on the south flank of the Horse Heaven Hills.

Two events occurred in May. On May 6^{th} a small (0.1 M_c) event occurred under Umtanum Ridge in the pre-basalt sediments. This was followed on May 24^{th} by a small (0.9 Mc) event in the crystalline basement near the RAW (Figures 4.1 and 4.5).

Six events occurred in June. Two small events occurred in the pre-basalt sediments and four in the crystalline basement. The first two events (0.5 and 0.3 M_c respectively) occurred on June 1st near the intersection of the northwest and northeast trends of the Horse Heaven Hills (Figure 4.1) in the basement sediments. A third small event (0.6 M_c) occurred on June 2nd in the crystalline basement and was 3 km north of the two June 1st events. A small event (0.1 Mc) occurred in the crystalline basement on June 16th north of Rattlesnake Mountain and east of Snively Basin. This was followed on June 23rd by a 1.7 Mc earthquake in the crystalline basement east of Gable Mountain and centered under the Columbia River. The final earthquake for the quarter (0.1 Mc) occurred under Umtanum Ridge at the same location and depth as the event of May 6th.

4.5.5 Fourth Quarter of FY 2002

The locations of all mapped earthquakes that occurred between July 1, 2002 and September 30, 2002 are shown on Figure 4.5.

4.5.5.1 Major Anticlinal Ridges

No earthquakes occurred along major anticlinal structures during the fourth quarter of FY 2002.

4.5.5.2 Earthquake Swarm Areas

A total of five earthquakes occurred in known swarm areas during the fourth quarter of FY 2002. Four occurred in the Horse Heaven Hills swarm area, and one in the Wahluke swarm area (see Figure 4.5).

4.5.5.2.1 Wahluke Swarm Area

The Wahluke swarm area that was active during the first quarter of FY 2002 and most of FY 2001 had another seismic event on August 7^{th} . The event was small (approximately $0.0~M_c$) and like the previously events, it occurred in the basalt.

4.5.5.2.2 Horse Heaven Hills Swarm Area

Four earthquakes occurred in the Horse Heaven Hills swarm area during the fourth quarter. All four earthquakes occurred in the pre-basalt sediments at depths of 7-10 km. Although we normally do not classify earthquakes below the basalt as swarms, the Horse Heaven Hills earthquakes are occurring like swarm events in the basalt so we are classifying these deeper events as swarm events. Two small (0.3 and 0.4 Mc) events occurred on August 11th. This was followed on August 13th by a third event (0.4 M_c). All three events were at the same location. They occurred at the same location as the dense cluster of earthquakes had during December 2001 and January 2002. A fourth event occurred at the south end of the Horse Heaven Hills earthquake swarm on September 11th. This event was also small (0.6 Mc) and it too occurred in the pre-basalt sediments.

4.5.5.3 Random or Floating Events

Eight events were classified as random during the fourth quarter. Two occurred in the basalt; one occurred in the pre-basalt sediments; and five occurred in the crystalline basement. The first event occurred in the Cold Creek depression northwest of 200 West Area. It was small (approximately $0.0~M_{\rm c}$) and occurred in the basalt about 2 km east of the Cold Creek fault. On August 9^{th} a small earthquake occurred in the basalt on the south flank of the Horse Heaven Hills. This event was approximately $0.0~M_{\rm c}$ and occurred about 3 km from the crest of the Horse Heaven Hills. On August 22^{nd} two events of nearly $0.0~M_{\rm c}$ occurred in the crystalline basement 3 km north of the Saddle Mountains in the Smyrna Bench area.

Four small events (0.0-0.1 Mc) occurred in the crystalline basement during the month of September. The first one occurred on September 4th under the Snively Basin portion of the Rattlesnake Mountain structure. This event occurred about 2 km west of the June 16th event and at a similar depth. The next event occurred on September 8th on the west side of the 200 East Area. The third event occurred in the pre-basalt sediments on the south flank of the Horse Heaven Hills. This earthquake was east of the Horse Heaven Hills swarm area but at a similar depth. The final event of the quarter occurred east of Gable Mountain in the same area as the August 9th event.

5.0 Strong Motion Accelerometer Operations

The Hanford SMA network has been in continuous operation since November 20, 1998. The nominal threshold used in the SMA network is 0.001 g in order to provide ground motion for smaller, non-damaging earthquakes that can be useful in estimating the ground motion expected from larger earthquakes, and to confirm correct operation of the instruments by analyzing the smaller-amplitude triggers (see Section 2.2).

5.1 FY 2002 Triggers of the Hanford SMA Network

During FY 2002 the Hanford SMA network did not trigger on any earthquakes.

6.0 Capabilities in the Event of a Significant Earthquake

The SMA network was designed to provide ground motion in areas at the Hanford Site that have high densities of people and/or facilities containing hazardous materials in order to insure the Hanford Site is in compliance with DOE Order 420.1, "Facility Safety." The network also allows Hanford Seismic Monitoring to support Hanford Site emergency services organizations in complying with DOE Order G 420.1-1, Section 4.7, "Emergency Preparedness and Emergency Communications," by providing area ground motion data in the event of an earthquake on the Hanford Site. This section summarizes the capabilities of the Seismic Monitoring Team in the event of an earthquake at Hanford.

6.1 Use of the SMA Network in the Event of an Earthquake

Historically, only a few facilities at the Hanford Site had instruments to provide data on peak ground accelerations or any type of ground motion. The present SMA instruments were located so that if an earthquake occurred, ground motion data would be readily available to assess the damage at the 100-K Area, the 200 East and West Areas, the 300 and 400 Area facilities, which have the greatest concentration of people and also contain hazardous materials.

Many facilities at the Hanford Site have undergone various degrees of seismic analysis either during design or during re-qualification. Although the seismic design of a building may be known, when an earthquake is felt, a determination must be made as to the extent of damage before it can be reoccupied and the systems restarted. A felt earthquake may not cause any significant damage to a building but, without adequate characterization of the ground motion, initial determination of the buildings possibility of having damage may be impossible.

In the event of an earthquake such as the 2001 Nisqually earthquake, building managers, emergency directors, and engineers can obtain ground motion data recorded by the SMA network from the Seismic Monitoring Team in the Sigma V Building. Normal hours of operation are between 6 a.m. and 4:30 p.m., Monday through Friday. If a SMA is triggered, the Seismic Monitoring Team will download events that were recorded and determine the peak ground accelerations. This information is then passed on to Hanford Emergency Services personnel where the facility engineers can use the data to determine if the ground motion exceeded, is equal to, or is less than the building design. This, along with assessments from trained engineers, allows the facility manager to make a rapid and cost-effective determination on whether a building is safe to re-occupy or should not be used until it has been inspected in more detail. Buildings that have designs exceeding the recorded ground motion could be put back into service very quickly; buildings with designs that are very close to or less than measured ground motion could be given priority for onsite damage inspections.

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