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Pacific Northwest National Laboratory
Operated by Battelle for the U.S. Department of Energy

Second and Third Quarters Hanford Seismic Report for Fiscal Year 1999

D. C. Hartshorn
S. P. Reidel
A. C. Rohay

October 1999

Prepared for the U.S. Department of Energy
under Contract DE-AC06-76RLO 1830

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Richland, Washington 99352

Summary

Hanford Seismic Monitoring provides an uninterrupted collection of high-quality raw and processed seismic data from the Hanford Seismic Network (HSN) for the U.S. Department of Energy and its contractors. Hanford Seismic Monitoring also locates and identifies sources of seismic activity and monitors changes in the historical pattern of seismic activity at the Hanford Site. The data are compiled, archived, and published for use by the Hanford Site for waste management, Natural Phenomena Hazards assessments, and engineering design and construction. In addition, the seismic monitoring organization works with the Hanford Site Emergency Services Organization to provide assistance in the event of a significant earthquake on the Hanford Site.

The HSN and the Eastern Washington Regional Network (EWRN) consist of 42 individual sensor sites and 15 radio relay sites maintained by the Hanford Seismic Monitoring staff. For the HSN, there were 270 triggers during the second quarter of fiscal year (FY) 1999 and 229 triggers during the third quarter on the primary recording system. During the second quarter, 22 seismic events were located; 11 were earthquakes in the Columbia River Basalt Group, 6 were earthquakes in the crystalline basement, and 5 were quarry blasts. Two earthquakes appear to be related to major geologic structures, eight earthquakes occurred in known swarm areas, and seven earthquakes were random occurrences. During the third quarter, 23 seismic events were located; 11 were earthquakes in the Columbia River Basalt Group, 4 were earthquakes in the pre-basalt sediments, 4 were earthquakes in the crystalline basement, and 4 were quarry blasts. Five earthquakes occurred in known swarm areas, six earthquakes formed a new swarm near the Horse Heavens Hills and Prosser, Washington, and eight earthquakes were random occurrences.

No earthquakes triggered the Hanford Strong Motion Accelerometers during the second or third quarters of FY 1999.

Acronyms

BWIP	Basalt Waste Isolation Project
CRBG	Columbia River Basalt Group
DMIN	closest distance from the epicenter to a station
DOE	U.S. Department of Energy
ETNA	strong motion accelerometer manufactured by Kinemetrics
EWRN	Eastern Washington Regional Network
FY	fiscal year
GAP	largest gap in event-station azimuth distribution
GPS	Global Positioning System
HSN	Hanford Seismic Network
M_c	Coda Length Magnitude
M_L	Local Magnitude
NP	number of p-wave and s-wave phases
NS	number of stations
PNNL	Pacific Northwest National Laboratory
RAW	Rattlesnake Mountain-Wallula Alignment
RMS	root-mean-square residual
SMA	strong motion accelerometer
USGS	United States Geological Survey
UTC	Universal Time, Coordinated
UW	University of Washington
WHC	Westinghouse Hanford Company
YFB	Yakima Fold Belt

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1.0 Introduction

1.1 Mission

The principal mission of seismic monitoring at the Hanford Site is to insure compliance with DOE Order 420.1, Facility Safety. This order establishes facility safety requirements related to nuclear safety design, criticality safety, fire protection, and natural phenomena hazards mitigation. With respect to seismic monitoring, the order states:

> 4.4.5 Natural Phenomena Detection.
> Facilities or sites with hazardous materials shall have instrumentation or other means to detect and record the occurrence and severity of seismic events.

In addition, seismic monitoring provides an uninterrupted collection of high-quality raw seismic data from the Hanford Seismic Network (HSN) located on and around the Hanford Site, and provides interpretations of seismic events from the Hanford Site and vicinity. Hanford Seismic Monitoring locates and identifies sources of seismic activity, monitors changes in the historical pattern of seismic activity at the Hanford Site, and builds a "local" earthquake database (processed data) that is permanently archived. The focus this report is the precise location of earthquakes proximal to or on the Hanford Site, specifically between 46 degrees and 47 degrees north latitude and between 119 degrees and 120 degrees west longitude. Data from the Eastern Washington Regional Network (EWRN) and other seismic networks in the northwest provide the Seismic Monitoring Project with necessary regional input for the seismic hazards analysis at the Hanford Site.

The seismic data are used by the Hanford Site contractors for waste management activities, Natural Phenomena Hazards assessments, and engineering design and construction. In addition, the Seismic Monitoring Project works with Hanford Site Emergency Services Organization to provide assistance in the event of an earthquake on the Hanford Site.

1.2 History of Seismic Monitoring at Hanford

Seismic monitoring at the Hanford Site was established in 1969 by the United States Geological Survey (USGS) under a contract with the U.S. Atomic Energy Commission. In 1975, the University of Washington (UW) assumed responsibility for the network and subsequently expanded it. In 1979, the Basalt Waste Isolation Project (BWIP) became responsible for collecting seismic data for the Hanford Site as part of site characterization activities. Rockwell Hanford Operations followed by Westinghouse Hanford Company (WHC), operated the local network and were the contract technical advisors for the EWRN operated and maintained by the UW. Funding ended for BWIP in December 1988. Seismic monitoring and responsibility for the UW contract were then transferred to WHC's Environmental Division. Maintenance responsibilities for the EWRN were also assigned to WHC who made major upgrades to EWRN sites.

Effective October 1, 1996, seismic monitoring was transferred to the Pacific Northwest National Laboratory (PNNL).[1] Seismic monitoring is part of PNNL's Applied Geology and Geochemistry Group, Energy Technology Division.

The Hanford Strong Motion Accelerometer (SMA) network was constructed during 1997 and came on line in May 1997. It operated continuously until September 30, 1997 when it was mothballed due to lack of funding. Funding was restored on October 1, 1998 by joint agreement between the U.S. Department of Energy (DOE) and PNNL. Operation of the free-field sites resumed on November 20, 1999.

1.3 Documentation

The Seismic Monitoring Project issues quarterly reports of local activity, an annual catalog of earthquake activity on and near the Hanford Site, and special-interest bulletins on local seismic events. The annual catalog includes the fourth quarter report for the fiscal year. Hanford Seismic Monitoring also provides information and special reports to other functions as requested. Earthquake information provided in these reports is subject to revisions if new data become available. In addition, an archive of all seismic data from the HSN is maintained by PNNL.

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[1] Pacific Northwest National Laboratory is operated by Battelle Memorial Institute for the U.S. Department of Energy.

2.0 Network Operations

2.1 Seismometer Sites

The seismic monitoring network consists of two designs of equipment and sites: seismometer sites and strong motion accelerometer (SMA) sites. Seismometer sites are designed to locate earthquakes and determine their magnitude and hypocenter location. SMA sites are designed to measure ground motion.

The HSN and the EWRN consists of 42 sensor sites. Most sites are in remote locations and require solar panels and batteries for power. The HSN uses 21 sites (Table 2.1 and Figure 2.1) and the EWRN uses 36 sites (Table 2.2 and Figure 2.2); both networks share 16 sites. The networks have 46 combined data channels because Gable Butte and Frenchman Hills East are three-component sites, each consisting of one vertical, one north-south horizontal, and one east-west horizontal data channel. The Frenchman Hills site is a new site that is under construction and will be operational by the end of the calendar year. Both networks use 15 additional telemetry relay sites. All sites or relays are transmitted to the Sigma V building.

During the period of time for this report, all sites and relays were transmitted to the 337 building in the 300 Area. This telemetry system used 15 additional relays and a series of telephone datalines. The central recording system was moved to the Sigma V building at the end of fiscal year (FY) 1999, and all sites and relays are now transmitted directly to Sigma V via radio telemetry.

2.1.1 Station Maintenance

The HSN's maintenance records for the seismic sensor and relay sites are filed in the Hanford Seismic Monitoring office.

2.1.2 Data Acquisition

The signals from the seismometer sites are monitored for changes in signal amplitude that are expected from earthquakes. The seismic network is subdivided into spatial groupings of stations that are monitored for nearly simultaneous amplitude changes, resulting in triggering a permanent recording of the events. The groupings and associated weighting schemes are designed to allow very small seismic events to be recorded and to minimize false triggers. Events are classified as locals (southcentral Washington near the Hanford Site), regionals (Western U.S. and Canada), and teleseisms (from farther distances around the world). Local and regional events are usually earthquakes, but mining explosions are also recorded. The latter can usually be identified from wave characteristics, time of day, and through confirmation with local government agencies and industries. Frequently, military exercises at the Yakima Training Center produce a series of acoustic shocks that unavoidably trigger the recording system. Sonic booms and thunder also produce acoustic signals that trigger the recording system.

Beginning in the second quarter, there was a significant change in the data acquisition systems. For many years, a single computer system recorded the triggered periods and wrote them to disk. A new

Table 2.1. Seismic Stations in the Hanford Seismic Network

The first column is the three-letter seismic station designator. The latitude and longitude, elevation above sea level in meters; and the full station name follow this. An asterisk before the three-letter designator means it is a three-component station. The locations of the stations are all in Washington; locations were derived from a Global Positioning System (GPS).

Station	Latitude Deg.Min.N	Longitude Deg.Min.W	Elevation (m)	Station Name
BEN	46N31.13	119W43.02	340	Benson Ranch
BRV	46N49.12	119W59.47	920	Black Rock Valley
BVW	46N48.66	119W52.99	670	Beverly
CRF	46N49.50	119W23.22	189	Corfu
ET3	46N34.64	118W56.25	286	Eltopia Three
*FHE	46N57.13	119W29.92	463	Frenchman Hills East
*GBB	46N36.49	119W37.62	177	Gable Butte
GBL	46N35.92	119W27.58	330	Gable Mountain
H2O	46N23.75	119W25.38	158	Water
LOC	46N43.02	119W25.85	210	Locke Island
MDW	46N36.79	119W45.66	330	Midway
MJ2	46N33.45	119W21.54	146	May Junction Two
OT3	46N40.14	119W13.98	322	Othello Three
PRO	46N12.73	119W41.15	550	Prosser
RED	46N17.92	119W26.30	366	Red Mountain
RSW	46N23.67	119W35.48	1,045	Rattlesnake Mountain
SNI	46N27.85	119W39.60	312	Snively Ranch
WA2	46N45.32	119W33.94	244	Wahluke Slope
WG4	46N01.85	118W51.34	511	Wallula Gap Four
WIW	46N25.76	119W17.26	128	Wooded Island
WRD	46N58.20	119W08.69	375	Warden

PC-based system was adapted from a USGS program and the UW system and was implemented at Hanford. One new system has been in continuous operation since January 6, 1999, initially running concurrently with the old system. The older system irrecoverably failed in early March, so that data could only be recovered through February 10. A second, backup PC system was installed in mid-March, and both new systems have been running in parallel since that time. Although the two new systems are practically identical, there is enough granularity in the trigger timing that they sometimes record exclusive events. In nearly all cases, these exclusive triggers are “false” triggers, not earthquakes or quarry blasts (i.e., from acoustic sources). The remainder are from barely detectable, small signals from regional and teleseismic earthquakes. By comparing the period from January 6 to February 10, it is clear that the new system was much more sensitive to seismic events and acoustic events than the older system. Some periods of down time have occurred on one or another system, but there were no times of simultaneous down time and no known earthquakes have been missed. Table 2.3 shows the number of each type of trigger for each quarter on the two new systems.

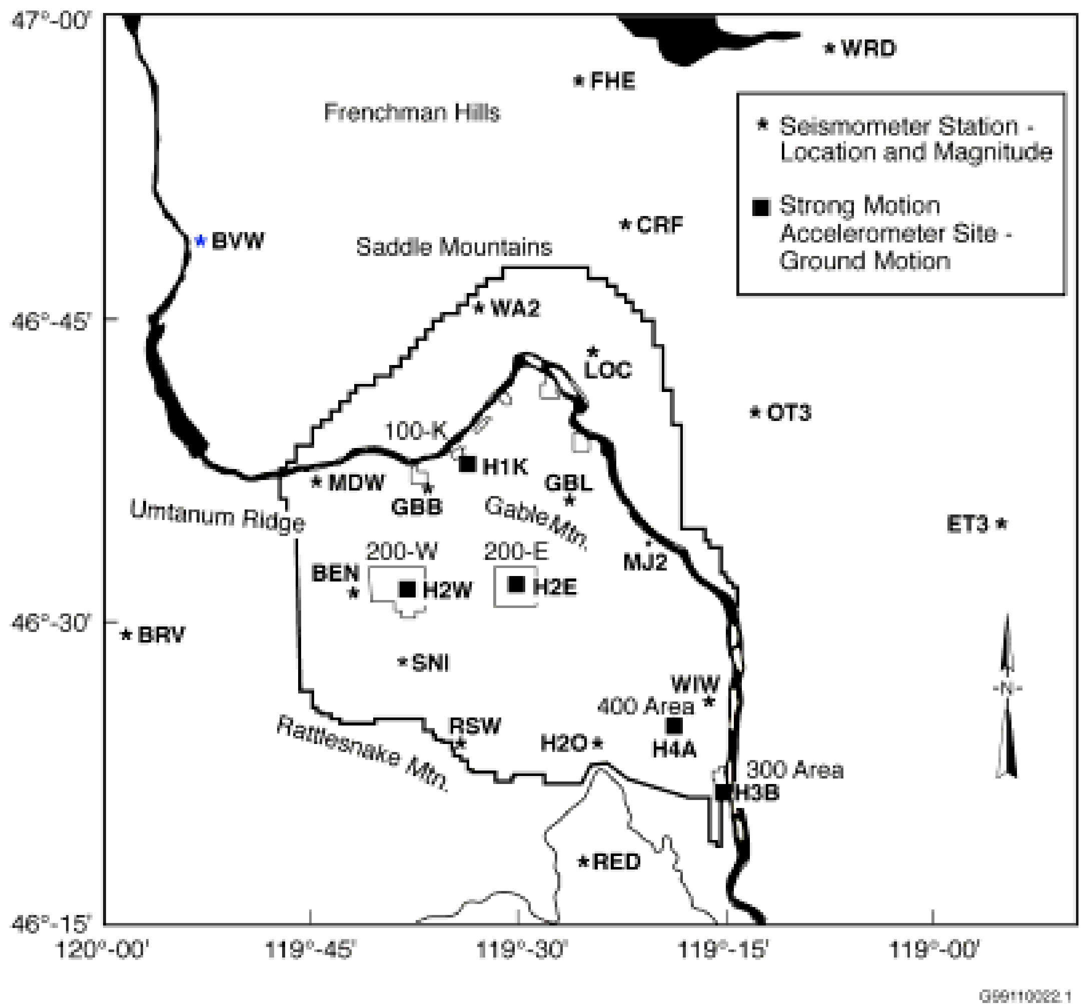

Figure 2.1. Locations of Seismograph Stations and Strong Motion Accelerometer Sites in the Hanford Seismic Network (see Table 2.1 for description of locations)

2.2 Strong Motion Accelerometer Sites

2.2.1 Location

The Hanford SMA network consists of five free-field SMA sites (Figure 2.1) (Table 2.4). There is one free-field SMA located in each of the 200 Separations Areas, one adjacent to the K-Basins in 100-K Area, one adjacent to the 400 Area where the Fast Flux Test Reactor is located, and one at the south end of the 300 Area.

The instrumentation locations were chosen based on two criteria (Moore and Reidel 1996): 1) instruments should be located in areas having the highest densities of people; and 2) instruments should be located in areas having hazardous facilities. Some of the highest concentrations of employees at Hanford are 200 East and West Areas, 100-K Area, the Fast Flux Test Facility (400 Area), and the 300 Areas. The

Table 2.2. Seismic Stations in the Eastern Washington Regional Network

The first column is the three-letter seismic station designator. The latitude and longitude, elevation above sea level in meters; and the full station name follow this. An asterisk before the three-letter designator means it is a three-component station. The locations of the stations are all in Washington unless otherwise indicated; locations were determined from a Global Positioning System (GPS).

Station	Latitude Deg.Min.N.	Longitude Deg.Min.W.	Elevation (m)	Station Name
BRV	46N29.12	119W59.47	920	Black Rock Valley
BVW	46N48.66	119W52.99	670	Beverly Washington
CBS	47N48.26	120W02.50	1,067	Chelan Butte, South
CRF	46N49.50	119W23.22	189	Corfu
DPW	47N52.25	118W12.17	892	Davenport
DY2	47N59.11	119W46.28	890	Dyer Hill Two
ELL	46N54.58	120W33.98	789	Ellensburg
EPH	47N21.38	119W35.76	661	Ephrata
ET3	46N34.64	118W56.25	286	Eltopia Three
ETW	47N36.26	120W19.94	1,477	Entiat
*FHE	46N57.13	119W29.92	463	Frenchman Hills East
*GBL	46N35.92	119W27.58	330	Gable Mountain
LNO	45N52.31	118W17.11	771	Lincton Mountain, Oregon
LOC	46N43.02	119W25.85	210	Locke Island
MDW	46N36.79	119W45.66	330	Midway
MJ2	46N33.45	119W21.54	146	May Junction Two
MOX	46N34.64	120W17.89	501	Moxee City
NAC	46N43.99	120W49.42	728	Naches
NEL	48N04.21	120W20.41	1,500	Nelson Butte
OD2	47N23.26	118W42.58	553	Odessa Two
OT3	46N40.14	119W13.98	322	Othello Three
PAT	45N52.92	119W45.14	262	Paterson
PRO	46N12.73	119W41.15	550	Prosser
RSW	46N23.67	119W35.48	1,045	Rattlesnake Mountain
SAW	47N42.10	119W24.03	701	St. Andrews
SNI	46N27.85	119W39.60	312	Snively Ranch
TBM	47N10.20	120W35.88	1,006	Table Mountain
TRW	46N17.32	120W32.31	723	Toppenish Ridge
TWW	47N08.29	120W52.10	1,027	Teanaway
VT2	46N58.04	119W58.95	1,270	Vantage Two
WA2	46N45.32	119W33.94	244	Wahluke Slope Two
WAT	47N41.92	119W57.24	821	Waterville
WG4	46N01.85	118W51.34	511	Wallula Gap Four
WIW	46N25.76	119W17.26	128	Wooded Island
WRD	46N58.20	119W08.69	375	Warden
YA2	46N31.60	120W31.80	652	Yakima Two

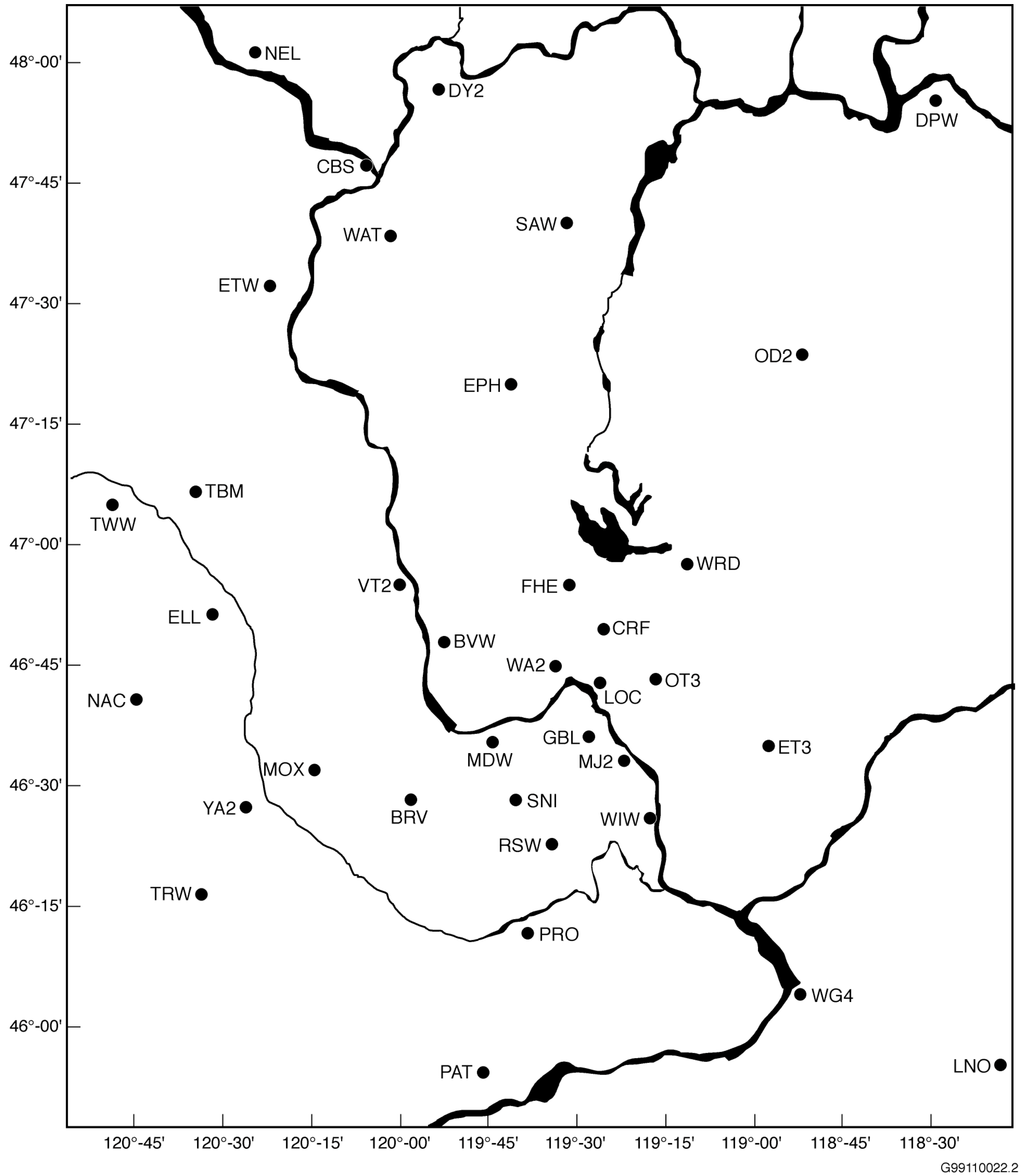

Figure 2.2. Locations of Seismograph Stations in the Eastern Washington Regional Network (see Table 2.2 for location descriptions)

Table 2.3. Acquisition System Recorded Triggers

Event Type	First Quarter	Second Quarter A	Second Quarter B	Third Quarter A	Third Quarter B	Fourth Quarter A	Fourth Quarter B	Remarks
Southcentral Washington	NA	30	7	41	37	-	-	Seismic events in southcentral Washington and northcentral Oregon that triggered the HSN
Regional	39	33	3	47	39	-	-	Seismic events in the Western United States and Canada
Teleseism	43	58	5	73	64	-	-	Seismic events at farther distances from around the world
Acoustic	12	122	33	21	20	-	-	Acoustical shock waves from artillery, thunder, and high-performance aircraft sonic booms
Noise	10	27	3	47	35	-	-	Triggers caused by data line circuits, lightning, maintenance triggers during system testing, high winds, coincidental noise at multiple sites within a trigger subnet, etc.
Total Triggers	121	270	51	229	195	-	-	
Local	14	22		23		-	-	Seismic events within the 46-47 degrees north latitude and 119-120 degrees west longitude
NOTE: A and B are the two new triggering computer systems.								

200 Areas are where all high-level radioactive waste from past processing of fuel rods has been stored in single-shell and double-shell tanks. In addition, the Canister Storage Facility that will hold encapsulated spent fuel rods is being constructed in 200 East Area. The 100-K Area contains the K Basins where all spent fuel rods from the N Reactor are stored prior to encapsulation. The Cold Vacuum Drying Facility, located in the 100-K Area, will be used to encapsulate spent fuel rods from the K Basins prior to shipment to the Canister Storage Building in 200 East Area.

2.2.2 Site Design

This section describes the designs of the SMA sites as of October 1, 1999. All sites were upgraded during FY 1999 to insure continuous operation and low maintenance.

All free-field SMA sites consist of two 30-gallon drums set in the ground such that the base of the drum is about 1 meter below the surface. One drum houses only the SMA; the other drum houses the electronics and communications equipment. A distance of 1 to 2.16 m (40 to 85 inches) separates the drum containing the electronics and communications equipment from the SMA drum; a sealed conduit connects the two drums.

The SMA instruments are three-component units consisting of one vertical, one north-south horizontal, and one east-west horizontal data channel. The instruments in use are the ETNA™ system (registered trademark of Kinemetrics, Inc.). Instrument specifications are summarized in Table 2.5. In

Table 2.4. Free-Field Strong Motion Accelerometer Sites

Site	Site ID	Location	Latitude Longitude Elevation	Design
100-K Area	H1K	South of K Basins outside 100 Area fence lines.	46° 38.51' 119° 35.53' 152 m	One free-field Kinematrics ETNA Model SMA housed in a ground vault.
200 East Area	H2E	East of B Plant; north of 7th street and east of Baltimore Ave.	46° 33.58' 119° 32.00' 210 m	One free-field Kinematrics ETNA Model SMA housed in ground vault.
200 West Area	H2W	Northeast of Plutonium Finishing Plant (PFP); north of 19th street and east of Camden Ave.	46° 33.23' 119° 37.51' 206 m	One free-field Kinematrics ETNA Model SMA housed in ground vault.
300 Area	H3A	South end of 300 Area inside fence lines. (NE 1/4, SW 1/4, Sec. 11, T10N, R28E).	46° 21.83' 119° 16.55' 119 m	One free-field Kinematrics ETNA Model SMA housed in ground vault.
400 Area	H4A	500 feet from fence line on east side of facility and north of parking area).	46° 26.13' 119° 21.30' 171 m	One free-field Kinematrics ETNA Model SMA housed in ground vault.

addition to the three-component SMA's, each ETNA SMA unit contains a computer, Global Positioning System (GPS) receiver and a modem (Figure 2.3). These systems are housed in a watertight box.

Two 100 amp-hour batteries that are housed in the equipment and communications drum (Figure 2.3) power the SMAs. The batteries currently are charged by four solar panels; a regulator is located between the solar panels and the batteries.

The communication link between the SMAs and the data analysis computer system housed in the Sigma V building is a cellular telephone/modem connection. The built-in modem in the SMA allows the system to use a cellular telephone to call an accelerometer or for the accelerometer to call out in the event it is triggered.

The SMAs have an internal GPS receiver used principally to link it to the National Bureau of Standards timing system. The GPS is internally activated approximately every 4 hours and checks the "location of the instrument" and the time. Any differences between the internal clock and the GPS time are recorded and saved by the SMA. Any corrections to the internal timing are made automatically. After the first 6 months of operation in 1997, the greatest difference recorded is approximately 4 milliseconds.

2.2.3 Strong Motion Accelerometer Operations Center

The combined operations, data recording, data interpretation, and maintenance facility is located in the Sigma V building and is operated by the PNNL Seismic Monitoring Team.

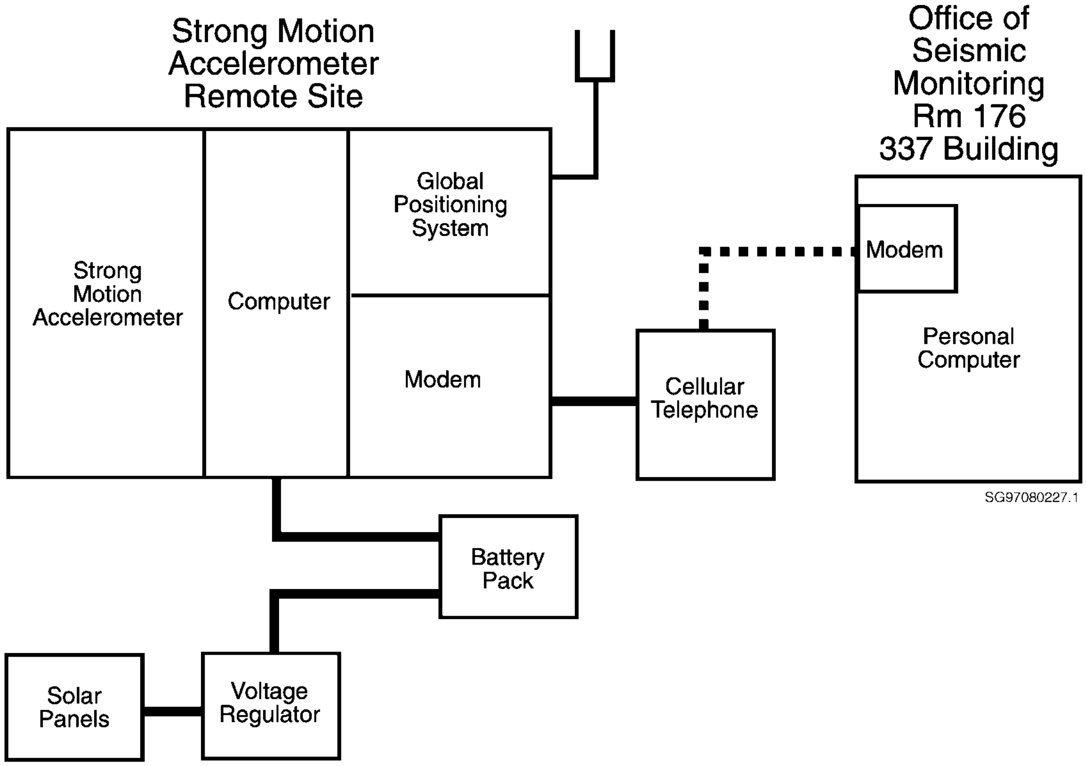

Figure 2.3. Schematic Diagram of a Strong Motion Accelerometer Installation

2.2.4 Strong Motion Operational Characteristics

The signals from the three accelerometer channels at each site are digitized with a 24-bit digitizer and temporarily stored in a memory buffer. The sampling rate of the digitizer is set to 200 Hz. The three channels are monitored for signals that equal or exceed a programmable trigger threshold. When one accelerometer channel is triggered, the other channels automatically record. The nominal threshold used is 0.05% of the full-scale range of 2.0 g (g is the acceleration of gravity, 9.8 m/s^2 or 32 ft/s^2) or 0.001 g. Threshold trigger levels are being adjusted to trigger infrequently on the noise sources (e.g., vehicles, sonic booms) near each site. This will provide ground motion data for smaller, non-damaging earthquakes that can be useful in estimating the ground motion expected from larger earthquakes, and to confirm correct operation of the instruments by analyzing the smaller-amplitude triggers. The recorders store information for 10 seconds before the trigger threshold is exceeded and for 40 seconds after the trigger ceases to be exceeded.

The SMA network is designed to transmit the data to the Hanford Seismic Recording Center at the Sigma V building or to be remotely accessed with a PC and modem. In addition, all SMAs can be accessed in the field where the data can be downloaded and interpreted.

Table 2.5. Instrument Parameters for the Kinematrics ETNA System in the Hanford SMA Network

Parameter	Value or Range
Sensor	
Type	Tri-axial Force Balance Accelerometer orthogonally oriented with internal standard
Full-Scale	2-2.5 g[a]
Natural Frequency	50 ± Hz range
Damping	approximately 70% critical[a]
Data Acquisition	
Number of Channels	3
Sample Rate	18-bit resolution @ 200 samples/second
Digital Output	Real-time, RS-232 Output Stream
Seismic Trigger	
Filter	0.1 - 12.5 Hz
Trigger level	0.10% - 0.20% g[b]
Alarm (call-out) Threshold	4.00% g
Pre-event Memory	10 sec
Post-event Time	40 sec

(a) Setting is dependent on instrument calibration.
(b) See Section 2.2.4 for discussion of trigger thresholds.

3.0 Magnitude, Velocity Models, and Quality Factors

3.1 Coda Length Magnitude

Coda-length magnitude (M_c), an estimate of local magnitude (M_L) (Richter 1958), is calculated using the coda-length/magnitude relationship determined for Washington State by Crosson (1972).

3.2 Velocity Model

The program XPED uses the velocities and layer depths given in Table 3.1. XPED was developed at the UW and the velocity model used in XPED is based on Rohay et al. (1985). XPED is an interactive earthquake seismogram display program used to analyze seismic events.

3.3 Quality Factors (Q)

XPED assigns a two-letter **Quality factor** (Table 3.2) that indicates the general reliability of the solution (**A** is best quality, **D** is worst). Similar quality factors are used by the USGS for events located with the computer program HYPO71. The first letter of the quality code is a measure of the hypocenter quality based primarily on travel time residuals. For example: Quality **A** requires a root-mean-square residual (**RMS**) less than 0.15 seconds while a **RMS** of 0.5 seconds or more is **D** quality (other estimates of the location uncertainty also affect this quality parameter). The second letter of the quality code is related to the spatial distribution of stations that contribute to the event's location, including the number of stations (**NS**), the number of p-wave and s-wave phases (**NP**), the largest gap in event-station azimuth distribution (**GAP**), and the closest distance from the epicenter to a station (**DMIN**). Quality **A** requires a solution with **NP** >8, **GAP** <90^o, and **DMIN** <5 km (or the hypocenter depth if it is greater than 5 km). If **NP** $\leq$5, **GAP** >180^o, or **DMIN** >50 km, the solution is assigned Quality **D**.

Table 3.1. Seismic Velocities for Columbia Basin Stratigraphy (from Rohay et al. 1985)

Depth to Top of Velocity Layer (km)	Stratigraphy	Velocity (km/sec)
0.0	Saddle Mountains and Wanapum Basalts and intercalated Ellensburg Formation	3.7
0.4	Grande Ronde Basalt and pre-basalt sediments	5.2
8.5	Crystalline Basement, Layer 1	6.1
13.0	Crystalline Basement, Layer 2	6.4
23.0	Sub-basement	7.1
38.0	Mantle	7.9

Table 3.2. Local Earthquake Data, October 1, 1998 to June 30, 1999

Event ID	Type	Date	Time	Latitude	Longitude	Depth	Mag	NS/NP	Gap	DMIN	RMS	Q	Location
98100916092	X	98/10/09	16:09:26.94	46N15.16	119W21.95	0.04	2.3	14/19	217	7	0.08	AD	6.7 km WSW of Richland
98101601453		98/10/16	01:47:42.76	46N21.14	119W23.19	1.61	0.1	4/07	171	5	0.09	AD	10.5 km SSW of FFTF
98103021542		98/10/30	21:54:21.39	46N28.19	119W05.34	11.61	0.6	10/14	190	15	0.08	AD	21 km east of FFTF
98111623560		98/11/16	23:56:01.77	46N12.11	119W48.35	19.19	0.9	13/18	266	27	0.06	AD	39 km SSW of 200 West
98111700013		98/11/17	00:01:33.71	46N13.27	119W47.43	21.49	1.2	9/14	273	24	0.06	AD	37 km SSW of 200 West
98111702504		98/11/17	02:50:49.13	46N09.96	119W49.60	17.78	1.4	15/20	275	12	0.06	AD	43.5 km of SSW of 200 West
98111702561		98/11/17	02:56:09.49	46N11.21	119W47.43	18.19	1.0	11/17	292	8	0.06	AD	40 km SSW of 200 West
98112203194		98/11/22	03:19:43.49	46N41.17	119W17.22	4.37	0.6	13/19	95	4	0.04	AB	19 km east of 100-N
98112203344		98/11/22	03:34:46.78	46N40.96	119W17.46	4.62	0.8	15/22	89	4	0.08	AA	20 km E of 100-N
98112203351		98/11/22	03:34:72.40	46N41.01	119W17.21	3.72	0.3	14/17	137	4	0.06	AC	21 km east of 100-N
98112203384		98/11/22	03:38:45.90	46N40.95	119W16.81	4.68	0.8	15/21	94	3	0.08	AB	21 km east of 100-N
98112203392		98/11/22	03:38:84.47	46N41.34	119W16.99	6.61	-0.6	5/09	147	4	0.02	AD	20 km east of 100-N
98112413312		98/11/24	13:31:19.90	46N13.92	119W48.70	12.76	0.0	4/05	329	24	0.06	BD	36 km SSW of 200 West
98112518342		98/11/25	18:34:21.75	46N40.71	119W23.76	24.99	0.5	5/08	262	15	0.02	AD	12 km east of 100-N
98120312550		98/12/03	12:55:07.92	46N50.21	119W17.70	1.57	0.9	15/24	138	7	0.04	AC	27 km NE of 100 N
99010703240		99/01/07	03:24:22.41	46N29.74	119W38.50	22.45	0.1	05/06	163	11	0.03	AD	6.9 km S of 200 West
99010709345		99/01/07	09:35:25.08	46N43.13	119W30.12	1.50	1.6	16/17	94	5	0.07	AB	11.2 km NE of 100-K
99010709371		99/01/07	09:37:29.10	46N43.58	119W29.89	4.01	0.9	10/10	159	5	0.06	AC	12.1 km NE of 100-K
99010801025		99/01/08	01:03:19.18	46N12.00	119W27.96	8.81	0.8	11/17	303	11	0.08	AD	16.3 km SW of Richland
99011018163		99/01/10	18:16:58.45	46N34.57	119W22.89	0.29	1.4	05/05	124	2	0.56	DD	11.9 km E of 200 East
99011118562		99/01/11	18:56:51.94	46N22.24	119W24.93	0.41	-0.1	04/06	156	2	0.26	BD	8.4 km SSW of 400 Area
99012502032		99/01/25	02:03:48.51	46N41.44	119W28.28	3.26	0.6	09/13	128	4	0.06	AB	11.0 km ENE of 100-K
99020407021		99/02/04	07:02:09.55	46N28.08	119W46.16	17.16	0.3	08/13	241	6	0.05	AD	14.2 km SW of 200 West
99020422351	P	99/02/04	22:35:23.53	46N15.95	119W37.08	0.46	1.1	07/08	175	7	0.57	DC	13.3 km ENE of Prosser
99021421375		99/02/14	21:38:17.60	46N57.89	119W32.07	0.02	1.0	09/11	266	19	0.13	AD	26.3 km SW of Moses Lake
99021721232	P	99/02/17	21:23:49.13	46N10.40	119W22.51	0.02	1.4	09/11	243	14	0.16	BD	13.8 km SSW of Richland
99021917565	P	99/02/19	17:57:25.42	46N31.11	119W53.41	2.37	1.5	15/21	106	8	0.38	CB	19.8 km WSW of 200 West

Table 3.2. (contd)

Event ID	Type	Date	Time	Latitude	Longitude	Depth	Mag	NS/NP	Gap	DMIN	RMS	Q	Location
99030104271		99/03/01	04:27:41.92	46N27.75	119W30.50	17.05	0.7	17/25	50	9	0.12	AA	11.0 km S of 200 East
99030122204	X	99/03/01	22:21:17.72	46N07.07	119W20.27	1.16	1.6	12/14	260	21	0.35	CD	18.7 km S of Richland
99030421433	X	99/03/04	21:43:53.88	46N11.33	119W21.41	0.03	1.3	07/07	264	13	0.11	AD	11.6 km SSW of Richland
99031813004		99/03/18	13:01:11.70	46N37.42	119W38.55	2.79	0.4	09/11	119	2	0.05	AB	4.0 km WSW of 100-K
99032106360		99/03/21	06:36:35.21	46N37.48	119W38.41	2.88	1.5	19/30	57	2	0.08	AA	3.8 km WSW of 100-K
99032403395		99/03/24	03:40:17.64	46N39.81	119W35.46	0.44	0.7	07/08	169	6	0.13	AC	2.4 km N of 100-K
99032420173		99/03/24	20:18:04.20	46N37.47	119W38.46	3.50	0.5	08/13	119	2	0.04	AB	3.9 km WSW of 100-K
99032513594		99/03/25	14:00:08.62	46N12.40	119W27.96	9.92	1.2	13/17	243	10	0.05	AD	15.9 km WSW of Richland
99032719241		99/03/27	19:24:41.12	46N12.76	119W28.54	8.42	0.3	08/11	252	9	0.09	AD	16.2 km WSW of Richland
99033004034		99/03/30	04:04:05.00	46N43.58	119W29.63	3.62	0.6	07/11	96	4	0.09	AB	12.3 km NE of 100-K
99040403411		99/04/04	03:41:36.22	46N22.82	119W15.91	3.98	1.0	08/10	186	5	0.07	AD	2.7 km NNE of 300 Area
99040403413		99/04/04	03:42:45.54	46N22.99	119W16.54	0.24	1.8	15/17	182	5	0.23	BD	2.8 km N of 300 Area
99040722360	P	99/04/07	22:36:22.48	46N54.97	119W54.44	2.27	2.0	17/18	99	8	0.19	BB	7.4 km SE of Vantage
99040800150	P	99/04/08	00:15:32.73	46N09.18	119W42.02	20.95	1.6	07/10	313	6	0.20	BD	8.0 km SE of Prosser
99040905545		99/04/09	05:55:18.79	46N22.73	119W22.44	0.43	1.4	13/19	158	4	0.23	BC	6.4 km S of 400 Area
99040906452		99/04/09	06:45:50.57	46N40.44	119W33.48	5.75	1.1	06/07	152	9	0.17	BC	4.6 km NE of 100-K
99041003584		99/04/10	03:58:11.23	46N36.13	119W48.22	7.87	0.8	07/09	200	3	0.08	AD	13.6 km WNW of 200 West
99041113390		99/04/11	13:39:27.93	46N42.60	119W33.27	4.10	0.6	06/10	159	5	0.07	AC	8.2 km NNE of 100-K
99041319271	P	99/04/13	19:27:45.25	46N13.82	119W36.27	0.02	2.0	14/16	192	6	0.34	CD	12.9 km ENE of Prosser
99041501354		99/04/15	01:36:03.10	46N48.22	119W42.72	14.09	1.0	10/14	202	12	0.04	AD	20.0 km NNW of 100-K
99041716583		99/04/17	16:58:59.22	46N12.32	119W30.50	6.48	0.3	05/07	245	11	0.12	AD	18.8 km WSW of Richland
99041806575		99/04/18	06:58:19.17	46N48.03	119W33.80	15.60	0.9	10/12	222	5	0.06	AD	17.8 km N of 100-K
99042200594		99/04/22	01:00:10.74	46N22.57	119W15.81	0.83	0.7	05/07	254	6	0.06	BD	2.3 km NNE of 300 Area
99042723160		99/04/27	23:16:27.77	46N27.91	119W37.82	17.64	0.8	09/14	107	2	0.09	AB	10.3 km S of 200 West
99051420084	P	99/05/14	20:09:07.18	46N46.80	119W58.36	0.34	2.1	07/07	208	7	0.20	BD	19.9 km S of Vantage

Table 3.2. (contd)

Event ID	Type	Date	Time	Latitude	Longitude	Depth	Mag	NS/NP	Gap	DMIN	RMS	Q	Location
99051513385		99/05/15	13:39:17.49	46N48.20	119W34.18	17.01	0.4	07/09	190	5	0.11	AD	18.1 km N of 100-K
99051709541		99/05/17	09:54:40.35	46N36.24	119W47.56	7.78	1.5	15/20	114	2	0.14	AB	12.9 km WNW of 200 West
99052401220		99/05/24	01:22:26.46	46N09.92	119W59.80	1.04	1.3	06/06	302	24	0.18	DD	18.3 km WSW of Prosser
99052401310		99/05/24	01:31:28.86	46N10.26	119W57.96	5.77	1.7	08/08	252	22	0.07	BD	15.8 km WSW of Prosser
99052420535		99/05/24	20:54:10.96	46N13.51	119W54.56	4.25	1.1	07/07	289	17	0.01	AD	11.1 km W of Prosser
99060810312		99/06/08	10:31:48.61	46N05.95	119W58.89	8.40	1.2	05/06	322	26	0.19	BD	20.4 km SW of Prosser
99060820095		99/06/08	20:10:15.62	46N12.16	119W58.27	0.41	0.8	05/06	301	22	0.21	CD	15.7 km W of Prosser
99061314485		99/06/13	14:49:19.04	46N04.32	119W54.70	0.77	0.7	07/09	227	23	0.08	AD	18.7 km SW of Prosser

Explanation of Table 3.2

EVENT ID: The Earthworm Recording System creates the identification number. XPED uses the year, month, day and time to create a unique number for each event.

DATE: The year and day of the year in Universal Time Coordinated (UTC). UTC is used throughout this report unless otherwise indicated.

TIME: The origin time of the earthquake given in UTC. To covert UTC to Pacific Standard Time, subtract eight hours; to Pacific Daylight Time, subtract seven hours.

LATITUDE: North latitude, in degrees and minutes, of the earthquake epicenter.

LONGITUDE: West longitude, in degrees and minutes, of the earthquake epicenter.

DEPTH: The depth of the earthquake in kilometers (km).

MAG: The magnitude is expressed as Coda-Length magnitude M_c, an estimate of local magnitude M_L (Richter 1958). If Magnitude is blank no determination could be made.

NS/NP: Number of stations/number of phases used in the solutions.

GAP: Azimuthal gap. The largest angle (relative to the epicenter) containing no stations.

DMIN: The distance from the earthquake epicenter to the closest station

RMS: The root-mean-square residual (observed arrival times minus the predicted arrival times) at all stations used to locate the earthquake. It is only useful as a measure of quality of the solution when five or more well-distributed stations are used in the solution. Good solutions are normally characterized by RMS values of less than about 0.3 seconds.

Q: The Quality Factors indicate the general reliability of the solution/location (**A** is best quality, **D** is worst). See Section 3.3 of this report: Quality Factors.

TYPE: P is Probable Blast; X is Confirmed Blast; F is Felt Earthquake; H is hand picked from helicorder; S is surficial event (rockslide, avalanche) and not a explosion or tectonic earthquake; blank is local earthquake**.**

4.0 Geology and Tectonic Analysis

The Hanford Site lies within the Columbia Basin, which is an intermontane basin between the Cascade Range and the Rocky Mountains that is filled with Cenozoic volcanic rocks and sediments. This basin forms the northern part of the Columbia Plateau physiographic province (Fenneman 1931) and the Columbia River flood-basalt province (Reidel and Hooper 1989). In the central and western parts of the Columbia Basin, the CRBG overlies Tertiary continental sedimentary rocks and is overlain by late Tertiary and Quaternary fluvial and glaciofluvial deposits (Campbell 1989; Reidel and others 1989; DOE 1988). In the eastern part, a thin (<100 m) sedimentary unit separates the basalt and underling crystalline basement and a thin (<10 m) veneer of eolian sediments overlies the basalt (Reidel and others 1989).

The Columbia Basin has two structural subdivisions or subprovinces: the Yakima Fold Belt (YFB) and the Palouse Slope. The YFB includes the western and central parts of the Columbia Basin and is a series of anticlinal ridges and synclinal valleys with major thrust faults along the northern flanks (Figure 4.1). The Palouse Slope is the eastern part of the basin and is the least deformed subprovince with only a few faults and low amplitude, long wavelength folds on an otherwise gently westward dipping paleoslope. Figure 4.2 shows north-south and east-west cross sections through the Columbia Basin based on surface mapping, deep boreholes, geophysical data (including the work of Rohay et al. [1985]), and magnetotelluric data obtained as part of BWIP (DOE 1988).

4.1 Earthquake Stratigraphy

Studies of seismicity at the Hanford Site have shown that the seismicity is related to crustal stratigraphy (layers of rock types) (Rohay et al. 1985; DOE 1988). The main geologic units important to earthquakes at Hanford and the surrounding area are:

- The Miocene CRBG
- The Paleocene, Eocene, and Oligocene sediments
- The crystalline basement (Precambrian and Paleozoic craton; Mesozoic accreted terranes).

4.2 Geologic Structure Beneath the Monitored Area

Between the late 1950s and the early 1980s, deep boreholes were drilled for hydrocarbon exploration in the Columbia Basin. These boreholes provided accurate measurements of the physical properties of the CRBG and the pre-basalt sediments (Reidel et al. 1994, 1998), but the thickness of the pre-basalt sediments and nature of the crystalline basement are still poorly understood. The difference between the thicknesses listed in Table 4.1 and the thicknesses of the crustal layers in the velocity model in Table 3.1 reflect data specific to the UW's crustal velocity model for eastern Washington. Table 4.2 is derived

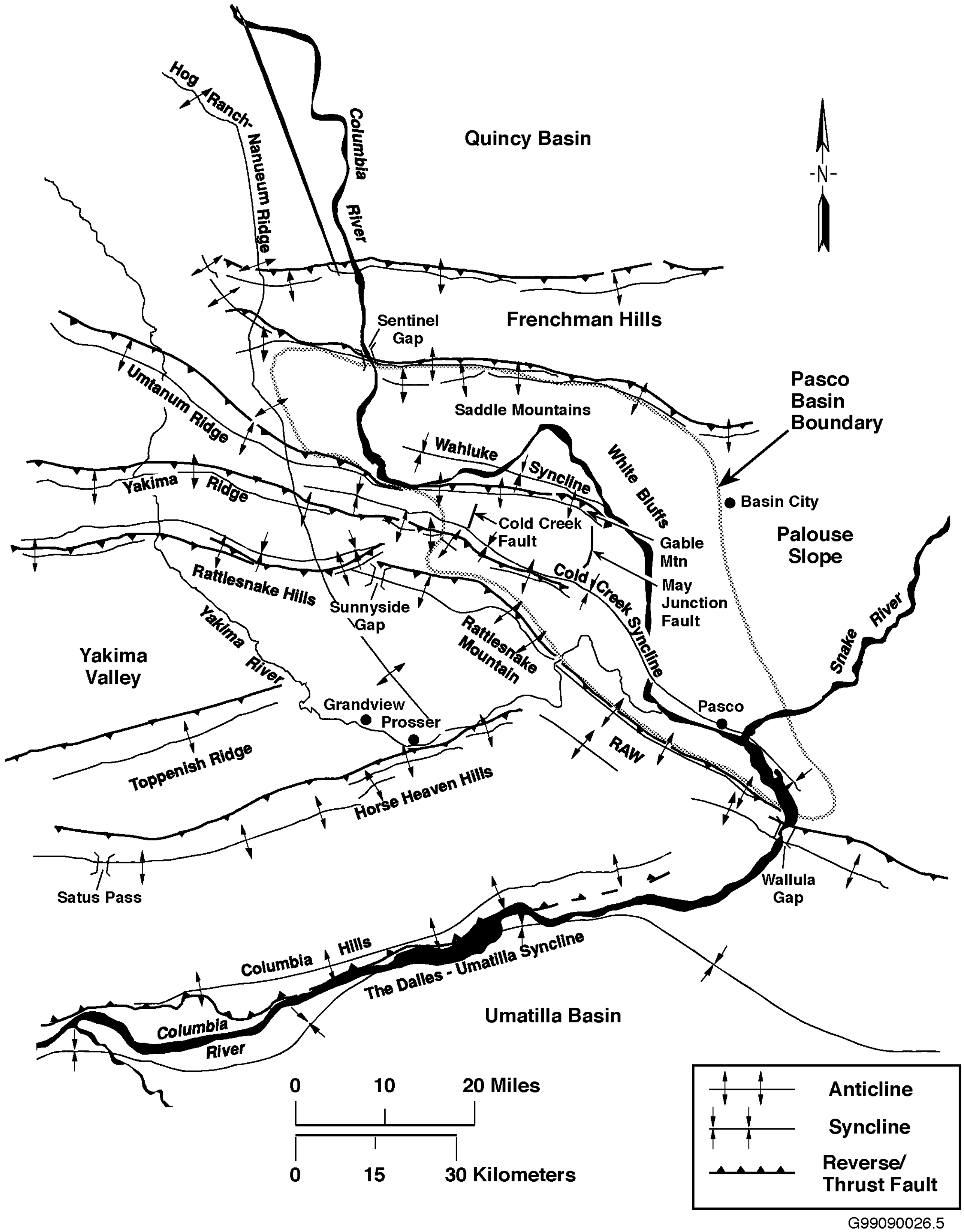

Figure 4.1. Structural and Tectonic Map of Columbia Basin Showing Major Seismic Source Structures

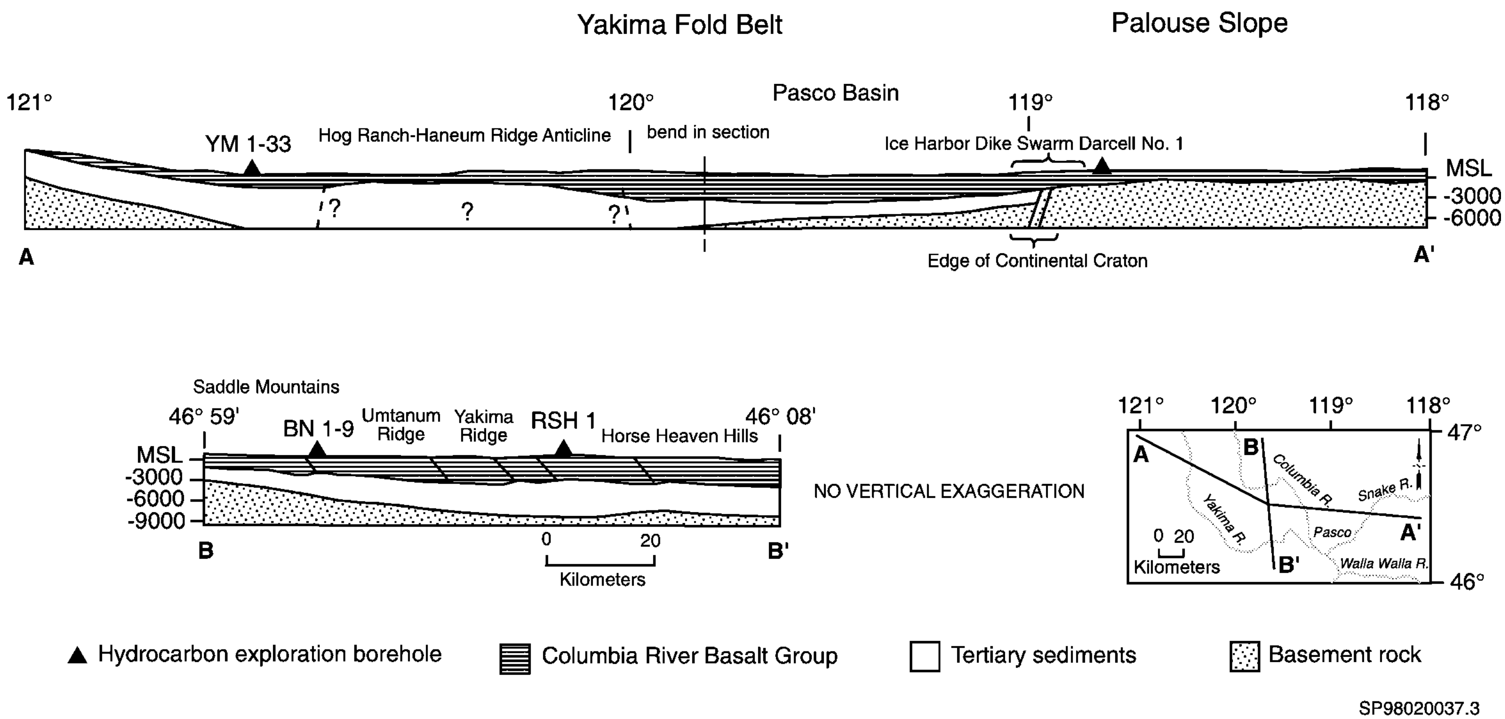

Figure 4.2. Geologic Cross Sections Through the Columbia Basin

from Figure 4.2 and was developed for the geologic interpretation in this report. The thicknesses of these units are variable across the monitored area. Table 4.1 summarizes the approximate thickness at the borders of the monitored area.

Table 4.1. Thicknesses of Stratigraphic Units in the Monitoring Area

Stratigraphy	North	South	East	West
Columbia River Basalt Group (includes suprabasalt sediments)	3.0 km	4.5 km	2.2 km	4.2 km
Pre-basalt Sediments	3.0 km	>4.5 km	0	>6.0 km

The thickness of the basalt and the pre-basalt sediments varies as a result of different tectonic environments. The western edge of the North American craton (late Precambrian/Paleozoic continental margin and Precambrian craton) is located in the eastern portion of the monitored area. The stratigraphy on the craton consists of CRBG overlying crystalline basement; the crystalline basement is continental crustal rocks that underlie much of the western North America. The stratigraphy west of the craton consists of 4-5 km of CRBG overlying greater than 6 km of pre-basalt sediments. This in turn overlies accreted terranes of Mesozoic age. The area west of the craton was subsiding during the Eocene and Oligocene, accumulating great thickness of pre-CRBG sediments. Continued subsidence in this area during the Miocene resulted in thicker CRBG compared to that on the craton. Subsidence continues today but at a greatly reduced rate (Reidel et al. 1994).

4.3 Depth of Earthquakes

Since records have been kept, about 75% of the earthquakes at the Hanford Site have originated in the CRBG layer. The pre-basalt sediments have had about 7% of the events and the crystalline basement has had 18%. Local earthquakes recorded for FY 1999 are listed in Table 4.2.

Table 4.2. Number of Local Earthquakes Occurring in Stratigraphic Units

Unit	First Quarter	Second Quarter	Third Quarter	Fourth Quarter	FY 1999
Basalt	7	11	11	-	29
Pre-basalt Sediments	0	0	4	-	4
Crystalline Basement	7	6	4	-	17
Total	14	17	19	-	50

4.4 Tectonic Pattern

Studies have concluded that earthquakes can occur in the following six different tectonic environments (earthquake sources) at the Hanford Site (Geomatrix 1996).

- **Reverse/thrust faults**. Reverse/thrust faults in the CRBG associated with major anticlinal ridges such as Rattlesnake Mountain, Yakima Ridge, and Umtanum Ridge could produce some of the largest earthquakes.

- **Secondary faults**. These are associated with the major anticlinal ridges.

- **Swarm areas**. Small geographic areas of unknown geologic structure produce clusters of events (swarms), usually in the CRBG in synclinal valleys. These clusters consist of a series of small shocks with no outstanding principal event. Swarms occur over a period of days or months and the events may number into the hundreds and then quit, only to start again at a later date. This differs from the sequence of foreshocks, mainshock, and trailing-off aftershocks that have the same epicenter or are associated with the same fault system. Three principal swarm areas are known at the Hanford Site. One is the Wooded Island Swarm Area along the Columbia River near the 300 Area. The second area, the Coyote Rapids Swarm Area, extends from the vicinity of the 100-K Area north-northeast along the Columbia River Horn to the vicinity of the 100-N Area. The third major swarm area is along the Saddle Mountains on the northern boundary of the Hanford Site. Other earthquake swarm areas are present, but activity is less frequent.

- **The entire Columbia Basin**. The entire basin, including the Hanford Site, could produce a "floating" earthquake. A floating earthquake is one that, for seismic design purposes, can happen anywhere in a tectonic province and is not associated with any known geologic structure. It is classified as a random event by Seismic Monitoring for purposes of seismic design and vibratory ground motion studies.

- **Basement source structures**. Studies (Geomatrix 1996) suggest that major earthquakes can originate in tectonic structures in the crystalline basement. Because little is known about geologic structures in the crystalline basement beneath the Hanford Site, earthquakes cannot be directly tied to a mapped fault. Earthquakes occurring in the crystalline basement without known sources are treated as random events for seismic hazards analysis and seismic design.

- **The Cascadia Subduction Zone**. This source recently has been postulated to be capable of producing a magnitude 9 earthquake. Because this source is along the western boundary of Washington State and outside the HSN, the Cascadia Subduction Zone is not an earthquake source that is monitored at the Hanford Site, so subduction zone earthquakes are not reported here. Because any earthquake along the Cascadia Subduction zone can have a significant impact on the Hanford Site (Geomatrix 1996), the UW monitors and reports on this earthquake source for DOE. Ground motion from any moderate or larger Cascadia Subduction Zone earthquake is detected by seismometers in the HSN.

4.5 Tectonic Activity

The locations for earthquakes that occurred in the first, second, and third quarters of FY 1999 are summarized in Tables 4.2 and 4.3. Earthquakes that occurred in the second and third quarters are described in the following sections.

Table 4.3. Summary of Earthquake Locations

		First Quarter	Second Quarter	Third Quarter	Fourth Quarter	FY 1999
Major Structures along anticlinal Ridges		1	2	0	-	3
Swarm Areas	Saddle Mountains	1	0	0	-	1
	Coyote Rapids	0	7	2	-	9
	Wooded Island	0	1	3	-	4
	Wahluke Slope	6	0	0		6
	Horse Heaven Hills	5	0	6	-	11
Random Events		6	7	8	-	21

4.5.1 Second Quarter of FY 1999

The locations of all located earthquakes that occurred between January 1, 1999 and March 31, 1999 are shown on Figure 4.3.

4.5.2 Major Anticlinal Ridges

During the second quarter, we interpret two seismic events to have occurred on major ridges. On January 11, a small (0.1 M_c), near-surface earthquake occurred along the Horn Rapids anticline near the location of an earthquake that occurred on October 16, 1998. On February 14, another small (1.0 M_c), near-surface event occurred along the north flank of the Frenchman Hills. Both events were located where the frontal fault zones of the respective ridges have been mapped.

4.5.3 Secondary Geologic Structures

We interpret no events to have occurred on a secondary geologic structure during the second quarter.

4.5.4 Swarm Area Activity

Eight earthquakes (47%) occurred in swarm areas during the second quarter of FY 1999 (Table 4.3). Seven events occurred in the Coyote Rapids swarm area and one occurred in the Wooded Island swarm area.

4.5.4.1 Wooded Island Swarm Area

A 1.4 M_c earthquake occurred on January 10 near Savage Island. This is on the northern extent of the Wooded Island swarm and also near the Othello swarm area (DOE 1989).

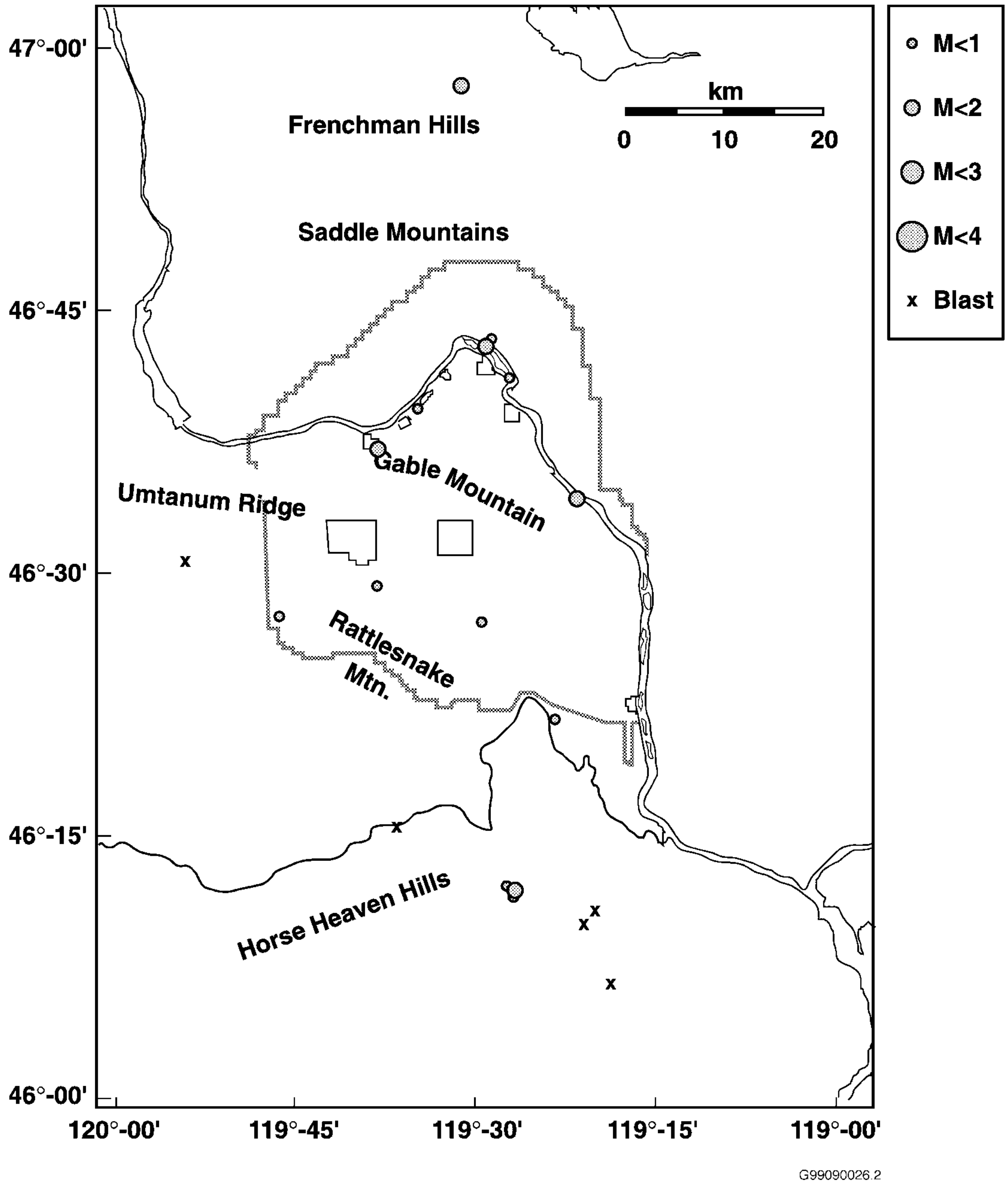

Figure 4.3. Locations of All Events Between January 1, 1999 and March 31, 1999 (Coda Length Magnitude (M_c) scale is shown at the side of the map)

4.5.4.2 Coyote Rapids Swarm Area

Seven events occurred in the Coyote Rapids swarm area during the second quarter of FY 1999. Two events (1.6 and 0.9 M_c respectively) occurred on January 7 near the "horn" of the Columbia River north of 100-D Area. A third event occurred near Lock Island 4 km southeast of the January 7 earthquakes.

Between March 21 and March 24, three earthquakes (0.4, 1.5, and 0.5 M_c, respectively) occurred within 2 km of 100-K Area. The seventh event occurred on March 24 near 100-N Area.

4.5.5 Random or Floating Events

There were 7 events (41%) classified as random events this quarter because they did not occur in a known earthquake swarm area or along a known geologic structure. One event occurred in the CRBG, three of the events were in the prebasalt sediments and four events were in the crystalline basement. No faults or folds have been identified below the basalt so events in the prebasalt sediments and crystalline basement are classified as random events.

Three small (>1.2 M_c) events occurred near Webber Canyon south of Kennewick, Washington on January 8, March 25, and March 27. All were in the sediments underlying the CRBG along the Horse Heaven Hills.

Two events occurred in the Cold Creek syncline. The first event (0.1 Mc) occurred on January 7, 5 km south of 200-West Area near Benson Ranch. The second event (0.7 M_c) occurred on March 1, 10 km southeast of the first event. Both events occurred in the crystalline basement.

On February 4, a small (0.3 M_c) earthquake occurred on the Arid Lands Ecology Reserve near the southwest corner of the Hanford Site. This event also occurred in the crystalline basement.

On March 30, a small (0.6 Mc) event occurred on the Wahluke Slope, 3 km south of the Saddle Mountains. This event occurred in the CRBG but it is not along any known fault.

4.6 Third Quarter of FY 1999

The locations of all located earthquakes that occurred between April 1, 1999 and June 30, 1999 are shown on Figure 4.4.

4.6.1 Major Anticlinal Ridges

No earthquakes occurred on major anticlinal ridges during this time period.

4.6.2 Swarm Area Activity

Eleven earthquakes (58%) occurred in swarm areas during the third quarter of FY 1999 (Table 4.3). Two events occurred in the Coyote Rapids swarm area, three occurred in the Wooded Island swarm area, and six occurred in a new swarm near Prosser, Washington and the Horse Heaven Hills.

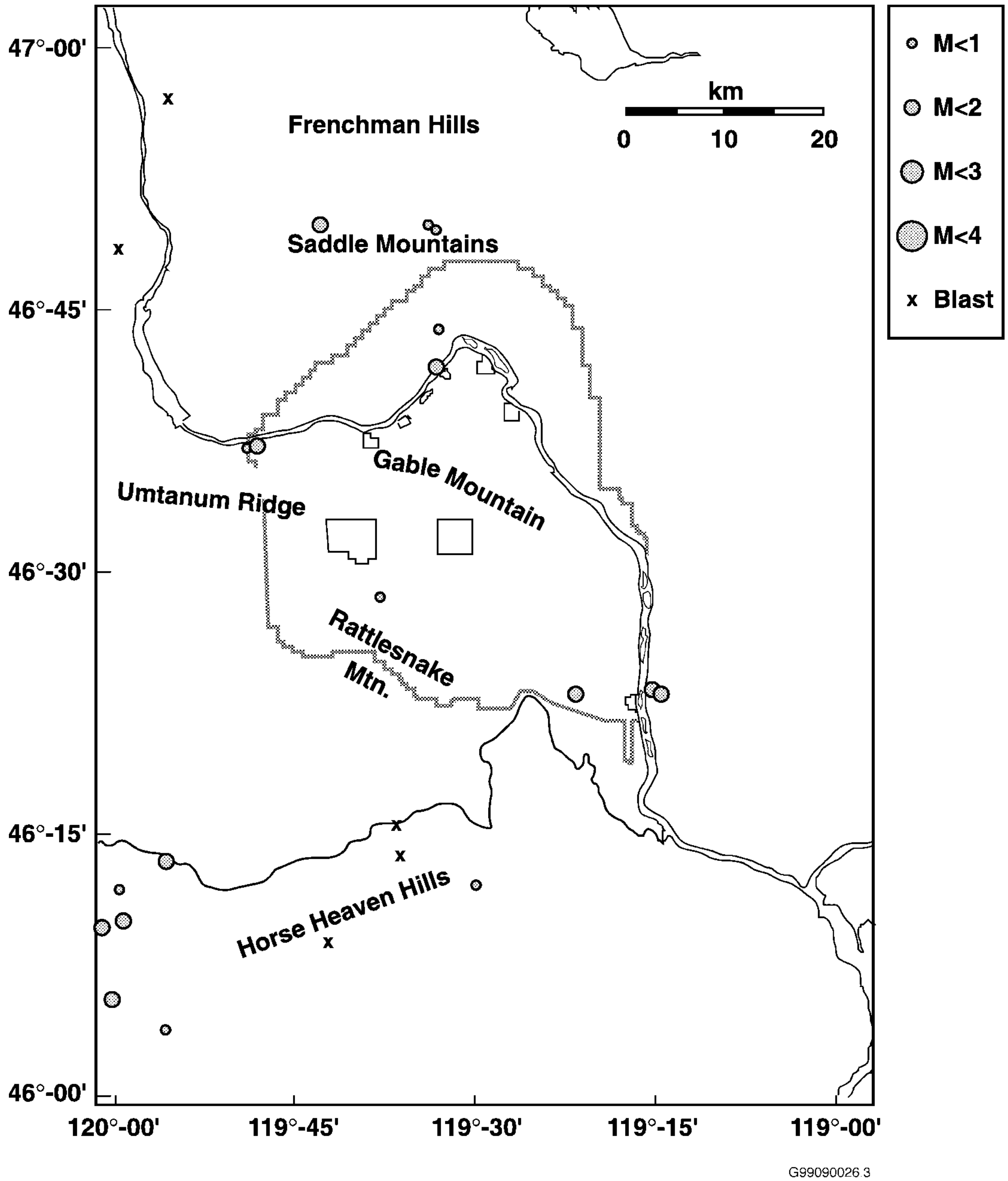

Figure 4.4. Locations of All Events Between April 1, 1999 and June 30, 1999 (Coda Length Magnitude (M_c) scale is shown at the side of the map)

4.6.2.1 Wooded Island Swarm Area

Two events (1.0 and 1.8 M_c, respectively) occurred on April 4 near Johnson Island. A third event occurred at the same location on April 22. All events were in the CRBG.

4.6.2.2 Coyote Rapids Swarm Area

Two events occurred in the Coyote Rapids swarm area near 100-N Area. The first event was 1.1 M_c on April 9 and the second was 0.6 M_c on April 11. An event occurred during the first quarter near the same location.

4.6.2.3 Horse Heaven Hills Swarm Area

Six earthquakes occurred in the Horse Heaven Hills area west of Prosser between May 24 and June 13. The largest event was 1.7M_c. This is a new swarm area. The first three events occurred on May 24 and were located north of the frontal fault of the Horse Heaven Hills. Each event was progressively farther north than the previous event. The next two events occurred on June 8 and the last event occurred on June 13. The locations and depths for the six events are less well constrained than for events closer to Hanford. The closest seismometer sites are Prosser (PRO, Figure 2), which provides coverage on the east side and Black Rock Valley (BRV) to the north.

4.6.3 Random or Floating Events

There were 8 events (42%) classified as random events because they did not occur in a known earthquake swarm area or along a known geologic structure. One event occurred in the CRBG, three of the events were in the sediments underlying the basalt and three events were in the crystalline basement. No faults or folds have been identified below the basalt so events that deep are classified as random events.

On April 9, a small (1.4 M_c), near surface event occurred near the Horn Rapids anticline. Earthquakes occurred 2 km to the west along the Horn Rapids anticline on January 11 and October 16. The April event was shallow and east of the earlier events. This event is tentatively classified as a random event even though it appears to be similar to the previous events.

Two events (1.4 and 1.5 M_c, respectively) occurred on the western boundary of the Hanford Site near Umtanum Ridge. The first event occurred on April 11 and the second on May 15. Both were in the sediments underlying the CRBG.

Three small events (M_c 1.0 or less) occurred in the crystalline basement beneath the Saddle Mountains. Two events (April 18 and May 15) were located near Wahatis Peak (0.6 M_c) in the Smyrna Bench segment and the third (April 17) was located 8 km to the east in the Sentinel Gap segment.

One small (0.3 M_c) event occurred on April 17 near Webber Canyon south of Kennewick, Washington. This event followed three other events (>1.2 M_c) that occurred there during the second quarter. All four earthquakes were in the sediments underlying the CRBG along the Horse Heaven Hills.

On April 27, a small (0.8 M_c) event occurred in the Cold Creek syncline north of the intersection of Rattlesnake Mountain and the Rattlesnake Hills in the crystalline basement. This earthquake occurred 3 km south of an event during the first quarter.

5.0 Strong Motion Accelerometer Operations

The Hanford SMA network was restarted November 20, 1999 after a one year hiatus. During the 1 month of operating during the first quarter and during the second and third quarters, there were no earthquake triggers. The SMA network had several triggers resulting from noise. During FY 1999 and FY 2000, the number of triggers resulting from noise and normal human activity will be monitored to determine the optimal settings for the triggering system. An optimum balance between having minimal triggers caused by noise and detection of the smallest possible earthquake is our objective.

6.0 Capabilities in the Event of a Significant Earthquake

The SMA network was designed to provide ground motion in areas at Hanford that have high densities of people and/or have hazardous facilities. This section summarizes the capabilities of the Seismic Monitoring Team in the event of an earthquake at Hanford.

6.1 Use of the SMA Network in the Event of an Earthquake

Historically, only a few facilities at the Hanford Site had instruments to provide data on peak ground accelerations or any type of ground motion. The present SMA instruments were located so that if an earthquake occurred, ground motion data would be readily available to assess the damage at the 100-K Area, the 200 East and West Areas, the 300 and 400 Area facilities, which have the greatest concentration of people, and all the hazardous materials.

Many facilities at the Hanford Site have undergone various degrees of seismic analysis either during design or during re-qualification. Although the seismic design of a building may be known, when an earthquake is felt, a determination must be made as to the extent of damage before it can be reoccupied and the systems restarted. A felt earthquake may not cause any damage to a building but without adequate characterization of the ground motion, initial determination of damage may be impossible.

In the event of an earthquake, building managers, emergency directors, and engineers can obtain ground motion data recorded by the SMA network from the Seismic Monitoring Team in Sigma V. If a SMA is triggered, the Seismic Monitoring Team will download events that were recorded and determine the peak ground accelerations and the spectral response curves. This information can then be used by the facility engineers to determine if the ground motion exceeded, is equal to, or is less than the building design. This, together with assessments from trained engineers, allows the facility manager to make a rapid and cost effective determination on whether a building is safe to reoccupy or should be not be used until it has been inspected in more detail. Buildings that have designs exceeding the recorded ground motion could be put back into service very quickly; buildings with designs that are very close to or less than measured ground motion could be given priority for onsite damage inspections.

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